

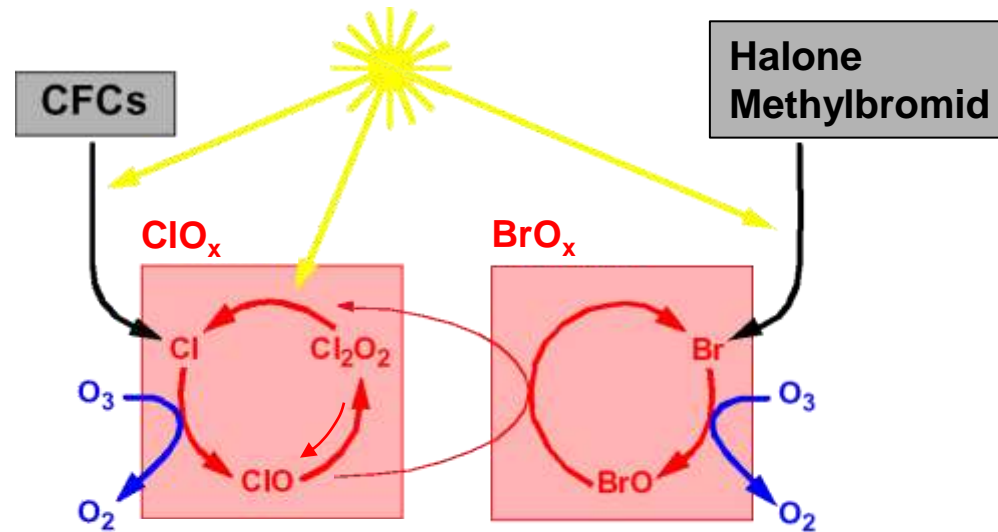
Polar Ozone Loss and the Tropical Tropopause Layer

Research based on meteorological fields
from data assimilation systems

Markus Rex

Alfred Wegener Institute for Polar and Marine Research
Potsdam, Germany

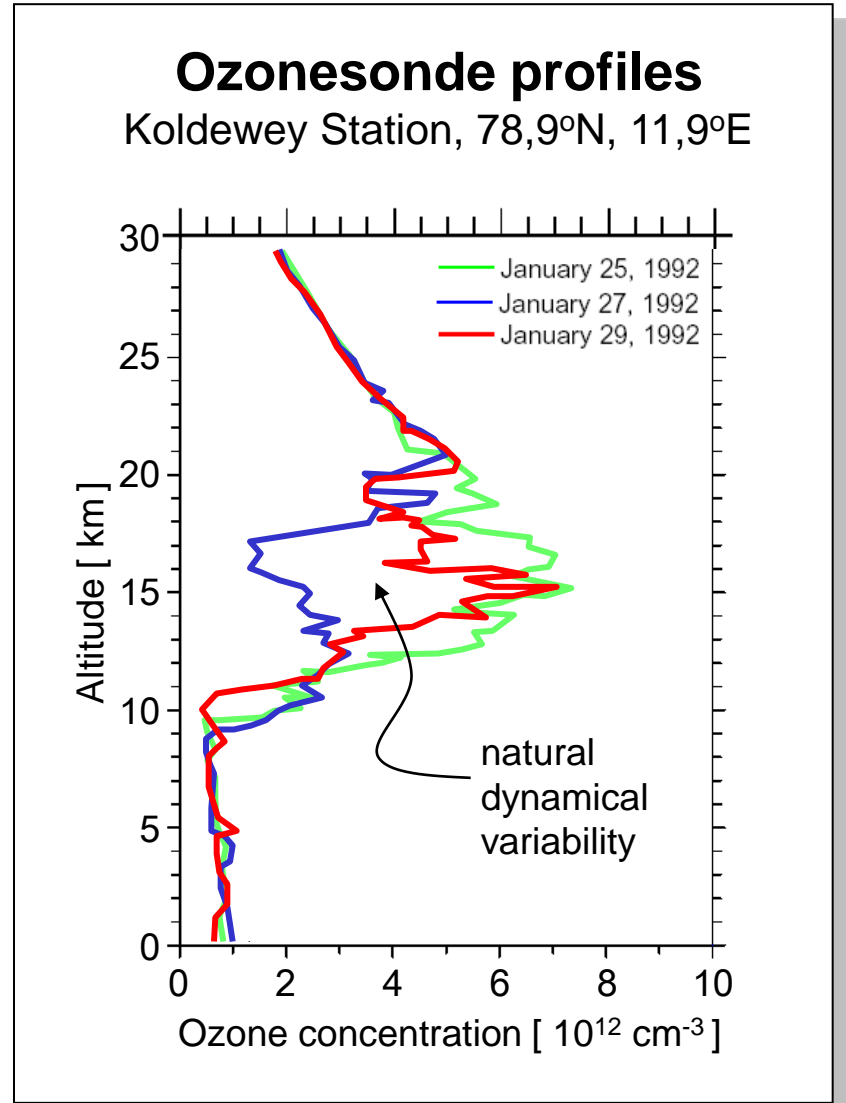
Polar ozone loss process



Ozonesondes

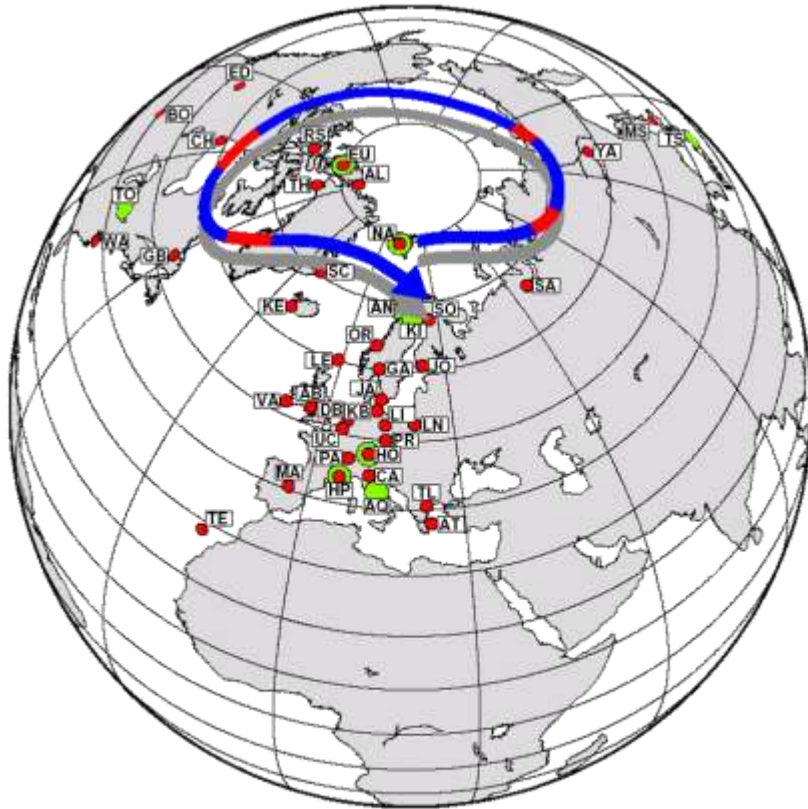


Ozonesonde launch in the Arctic



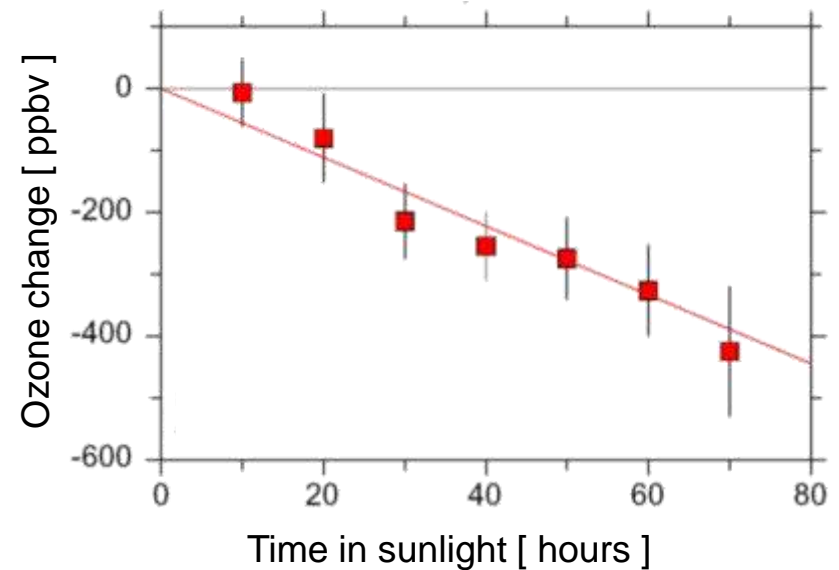
Rex et al., JGR, 1998

The Match project: Lagrangian measurements of chemical ozone loss rates

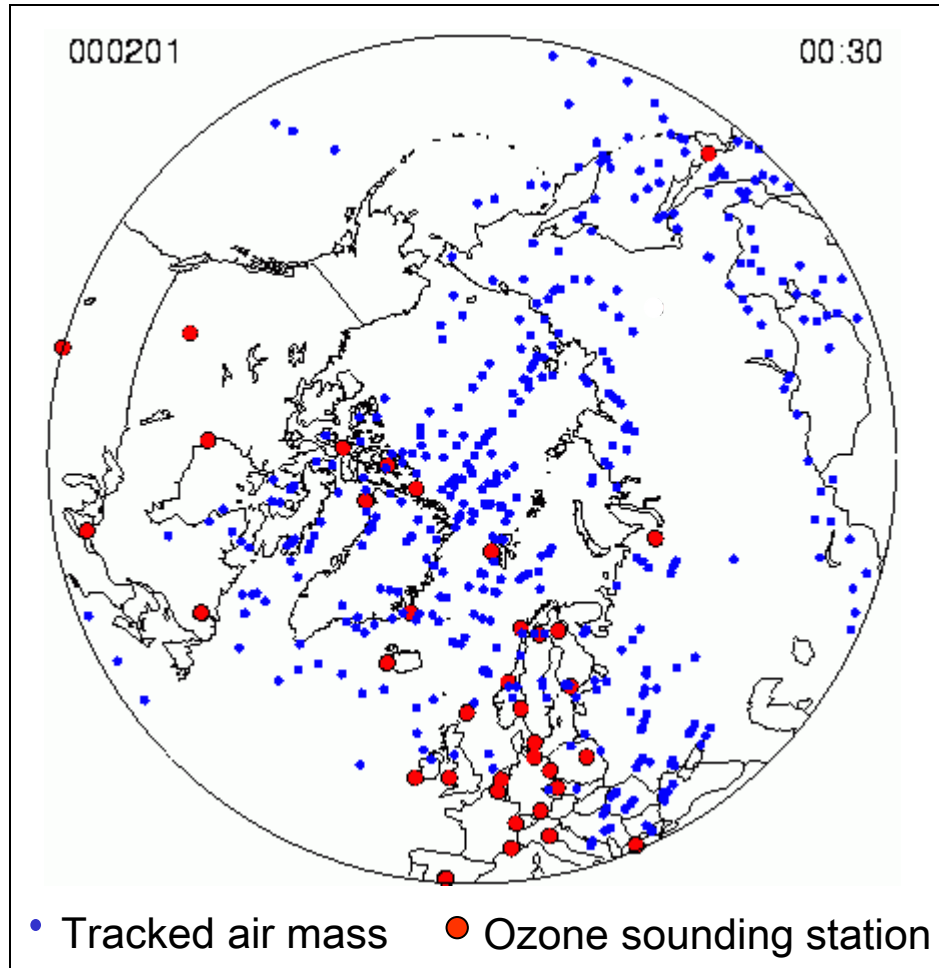


- Ozone sounding station ● Lidar station
- ➡ Air mass trajectory (day/night)

=> Chemical ozone loss rate



The Match project



- International network of ~30 Stations operated by 17 nations.
- ~500-1200 ozonesonde launches per winter.
- 15 Arctic and 2 Antarctic campaigns since the early 1990s.

e.g.:

Von der Gathen, Rex et al., *Nature*, 1995
Rex, von der Gathen et al, *Nature*, 1997
Manney, Santee, Rex et al., *Nature*, 2011

Multiple Sensor approach for a System of three sensors (e.g. (A) ILAS II, (B) POAM III, and (C) sondes)

Standard Matches:

i Matches of Type A-A:

$$\begin{aligned} \Delta O_3^{i1} &= -L t_s^{i1} + e^{i1} \\ \vdots \\ \Delta O_3^i &= -L t_s^i + e^i \end{aligned}$$

j Matches of Type B-B:

$$\begin{aligned} \Delta O_3^{i+j} &= -L t_s^{i+j} + e^{i+j} \\ \vdots \\ \Delta O_3^{i+j} &= -L t_s^{i+j} + e^{i+j} \end{aligned}$$

k Matches of Type C-C:

$$\begin{aligned} \Delta O_3^{i+j+k} &= -L t_s^{i+j+k} + e^{i+j+k} \\ \vdots \\ \Delta O_3^{i+j+k} &= -L t_s^{i+j+k} + e^{i+j+k} \end{aligned}$$

Mixed Matches:

l of Matches Type A-B:

$$\begin{aligned} \Delta O_3^{i+j+k+l} &= \text{Bias}_B \\ \vdots \\ \Delta O_3^{i+j+k+l} &= \text{Bias}_B \end{aligned}$$

m Matches of Type A-C:

$$\begin{aligned} \Delta O_3^{i+j+k+l+m} &= \text{Bias}_C \\ \vdots \\ \Delta O_3^{i+j+k+l+m} &= \text{Bias}_C \end{aligned}$$

p Matches of Type B-C:

$$\begin{aligned} \Delta O_3^{i+\dots+o+1} &= \text{Bias}_C - \text{Bias}_B \\ \vdots \\ \Delta O_3^{i+\dots+p} &= \text{Bias}_C - \text{Bias}_B - L t_s^{i+\dots+p} + e^{i+\dots+p} \end{aligned}$$

Caution: The errors in these ensembles of Matches are correlated. Covariance matrix is important for uncertainty estimates.
c.f. Lehman et al., 2005

$$\begin{aligned} &+k+l+1 + e^{i+j+k+l+1} \\ &j+k+l+m + e^{i+j+k+l+m} \\ &L t_s^{i+j+k+l+m+n+1} + e^{i+j+k+l+m+n+1} \\ &L t_s^{i+j+k+l+m+n+o} + e^{i+j+k+l+m+n+o} \\ &- L t_s^{i+\dots+p+1} + e^{i+\dots+p+1} \\ &\vdots \\ &\Delta O_3^{i+\dots+p+q} = \text{Bias}_C - \text{Bias}_B - L t_s^{i+\dots+p+q} + e^{i+\dots+p+q} \end{aligned}$$

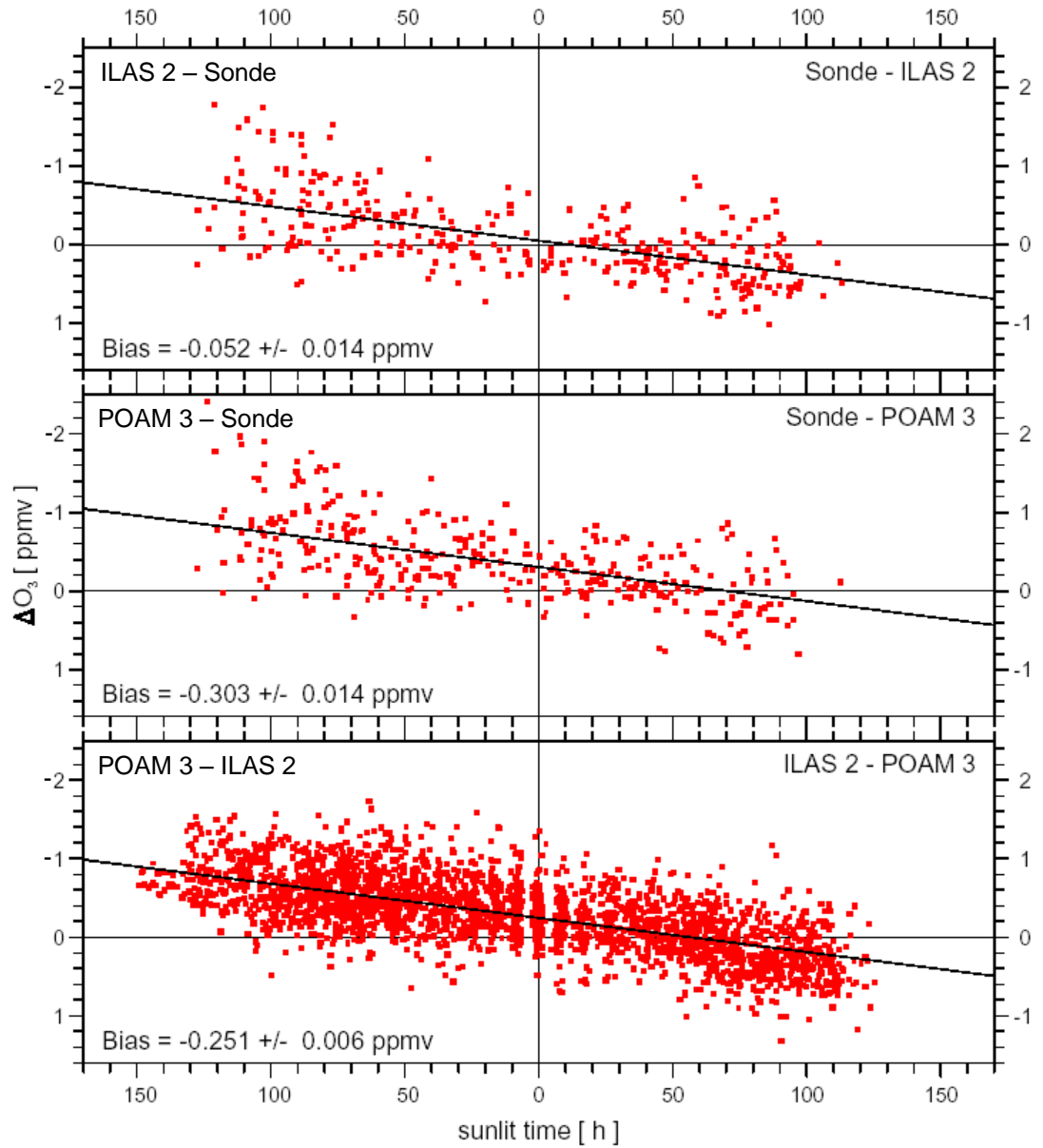
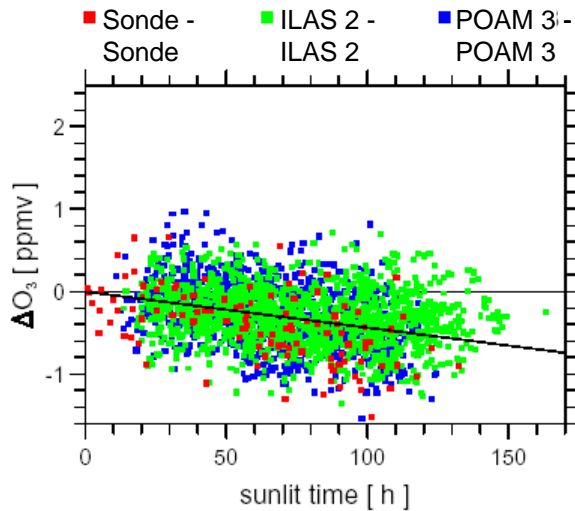
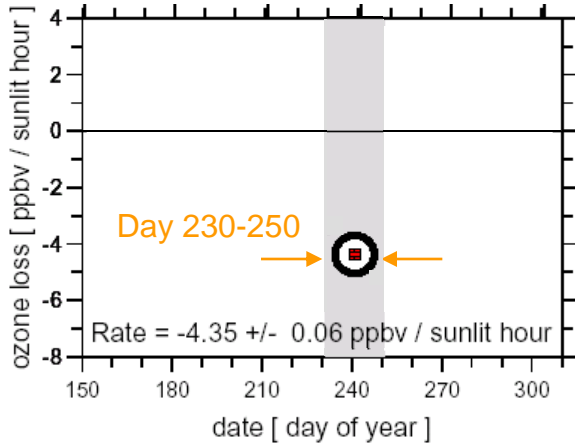


$i+j+k+l+m+n+o+p+q$ equations for the parameters L , Bias_B , and Bias_C , which can be determined simultaneously by a multivariate regression analysis

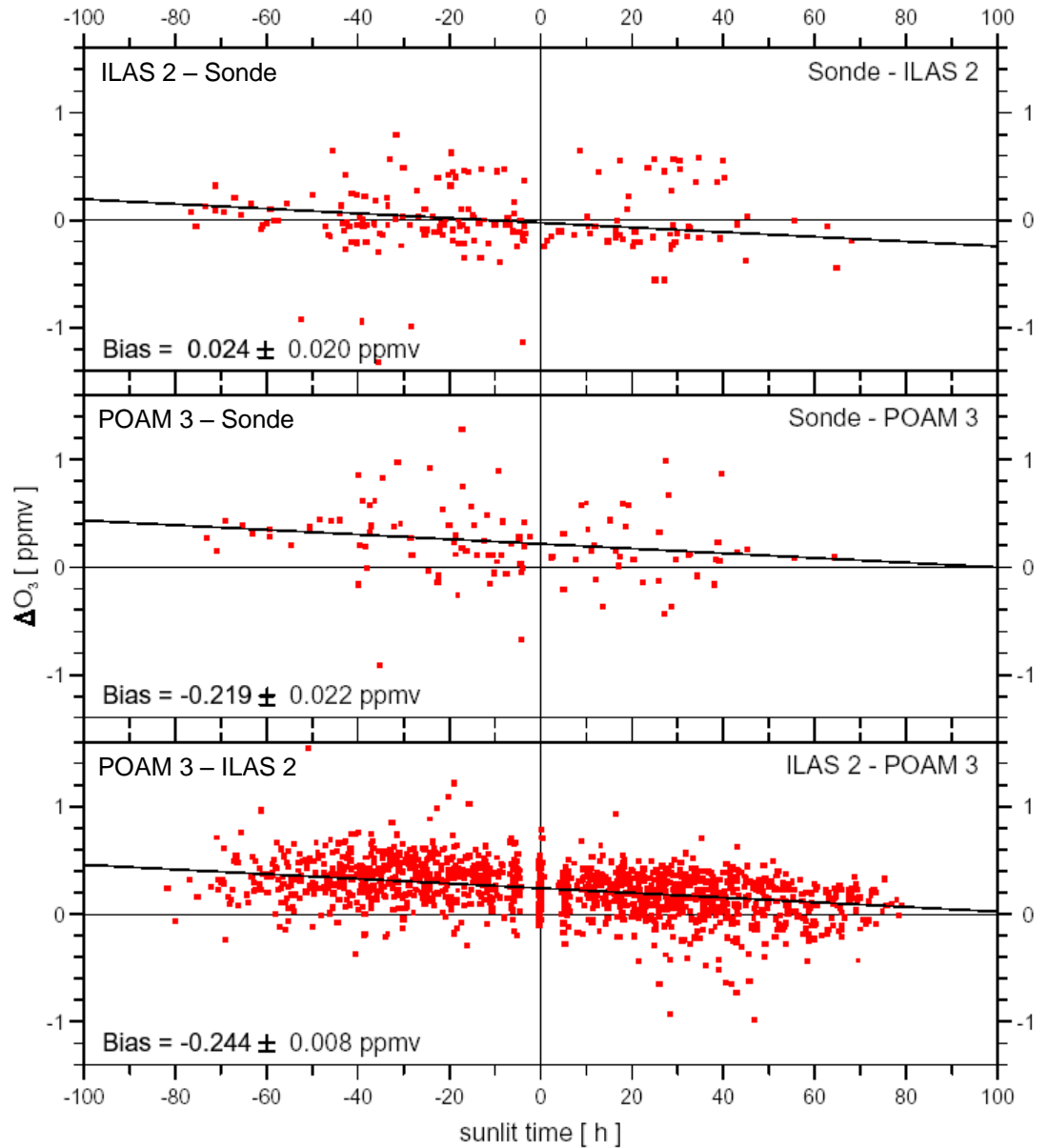
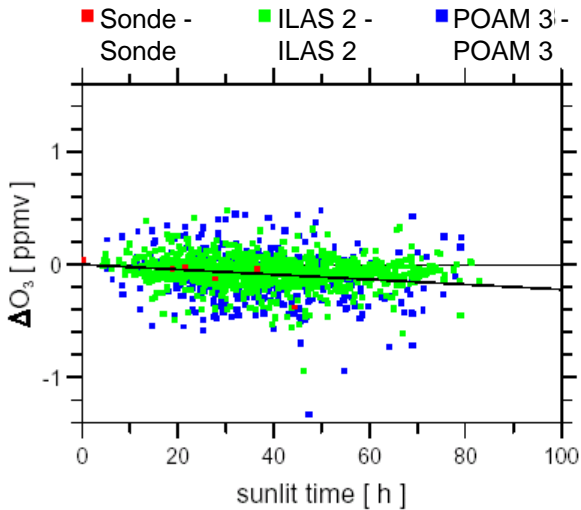
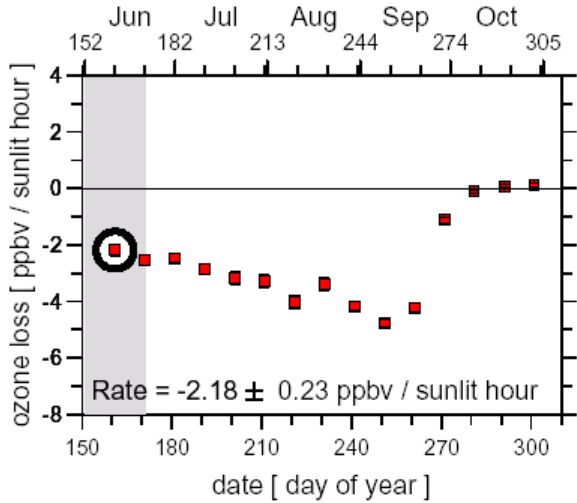


Compared to standard Match: Ninefold increase in the number of equations for a threefold increase in the number of parameters

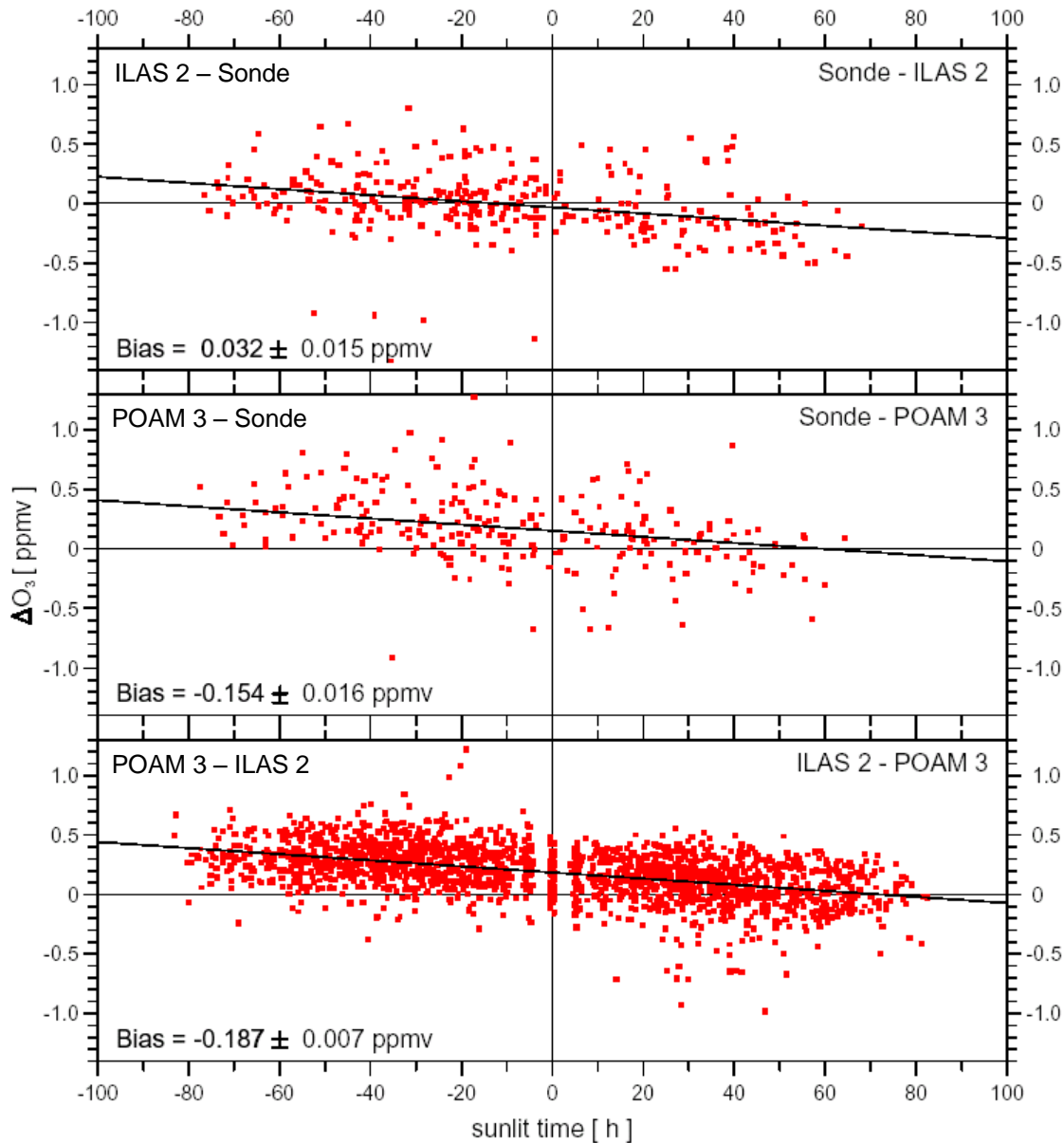
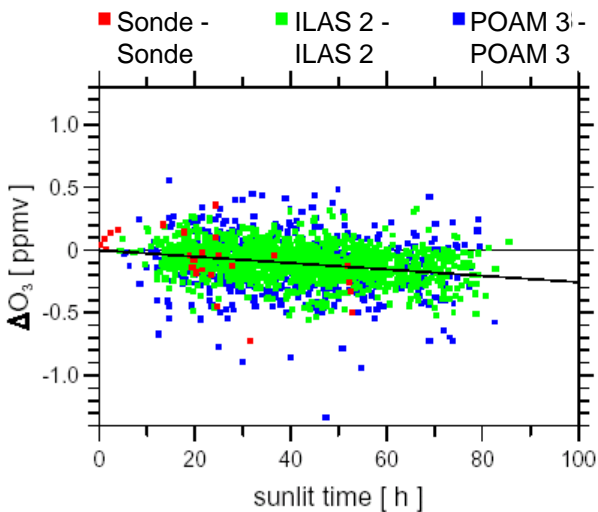
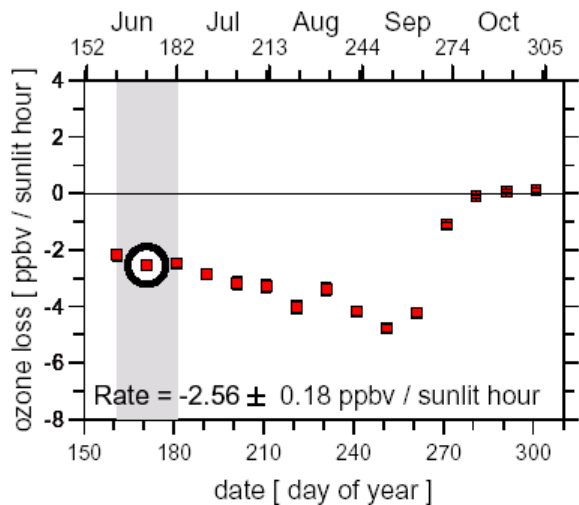
465 K - 485 K



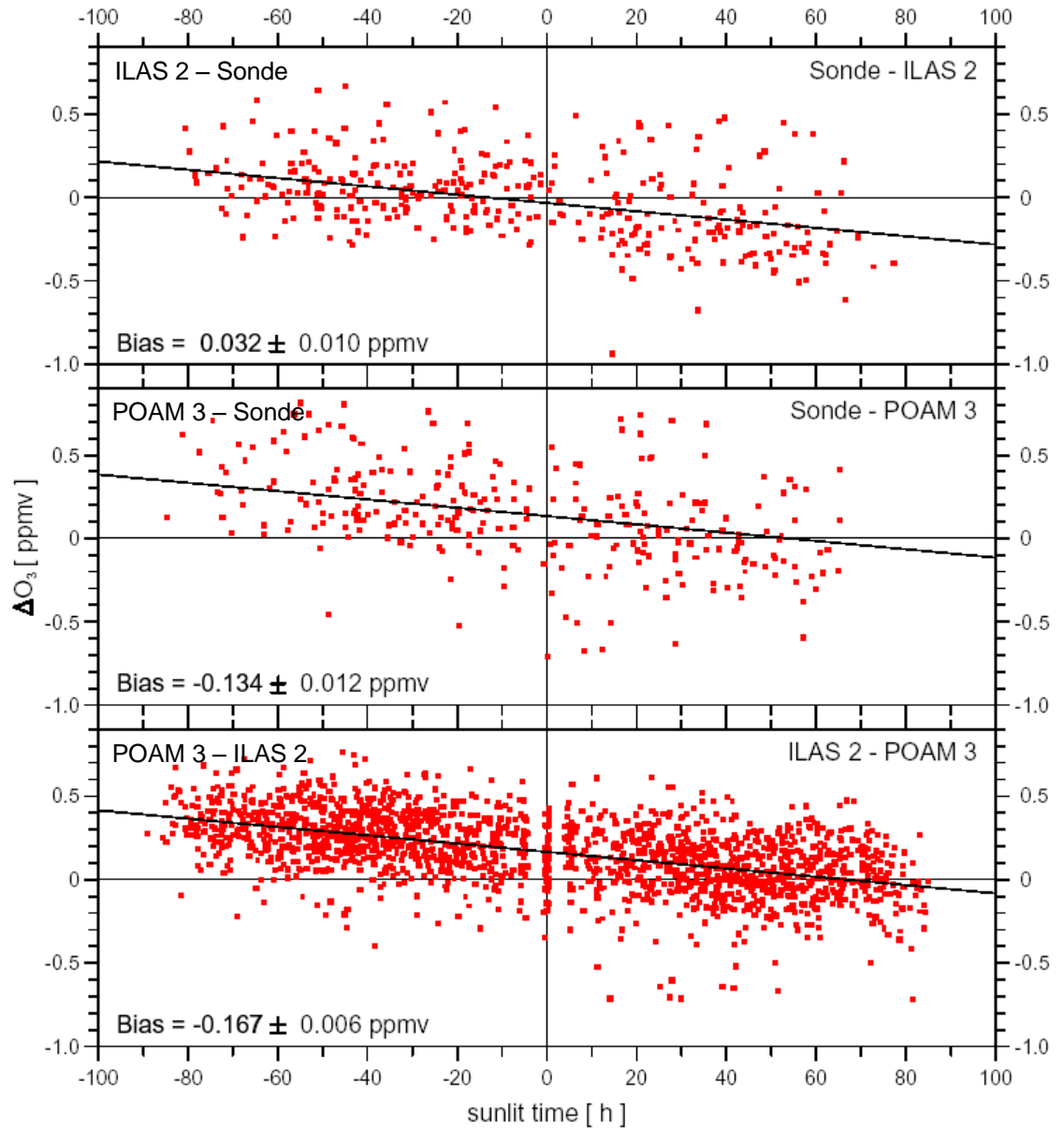
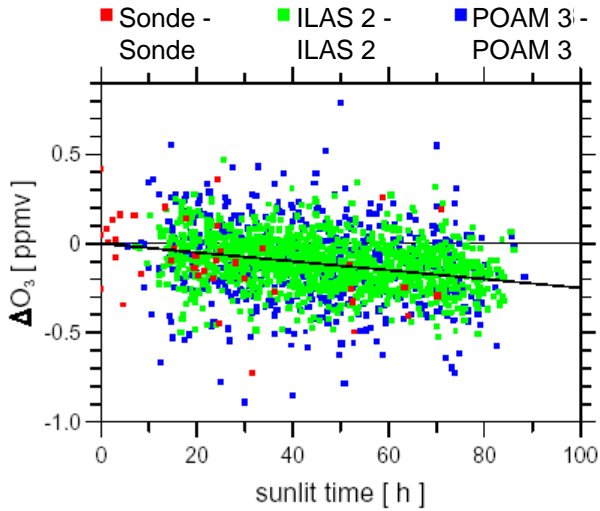
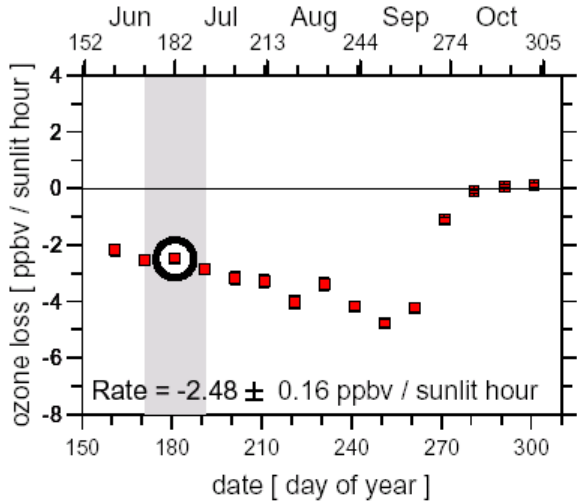
465 K - 485 K



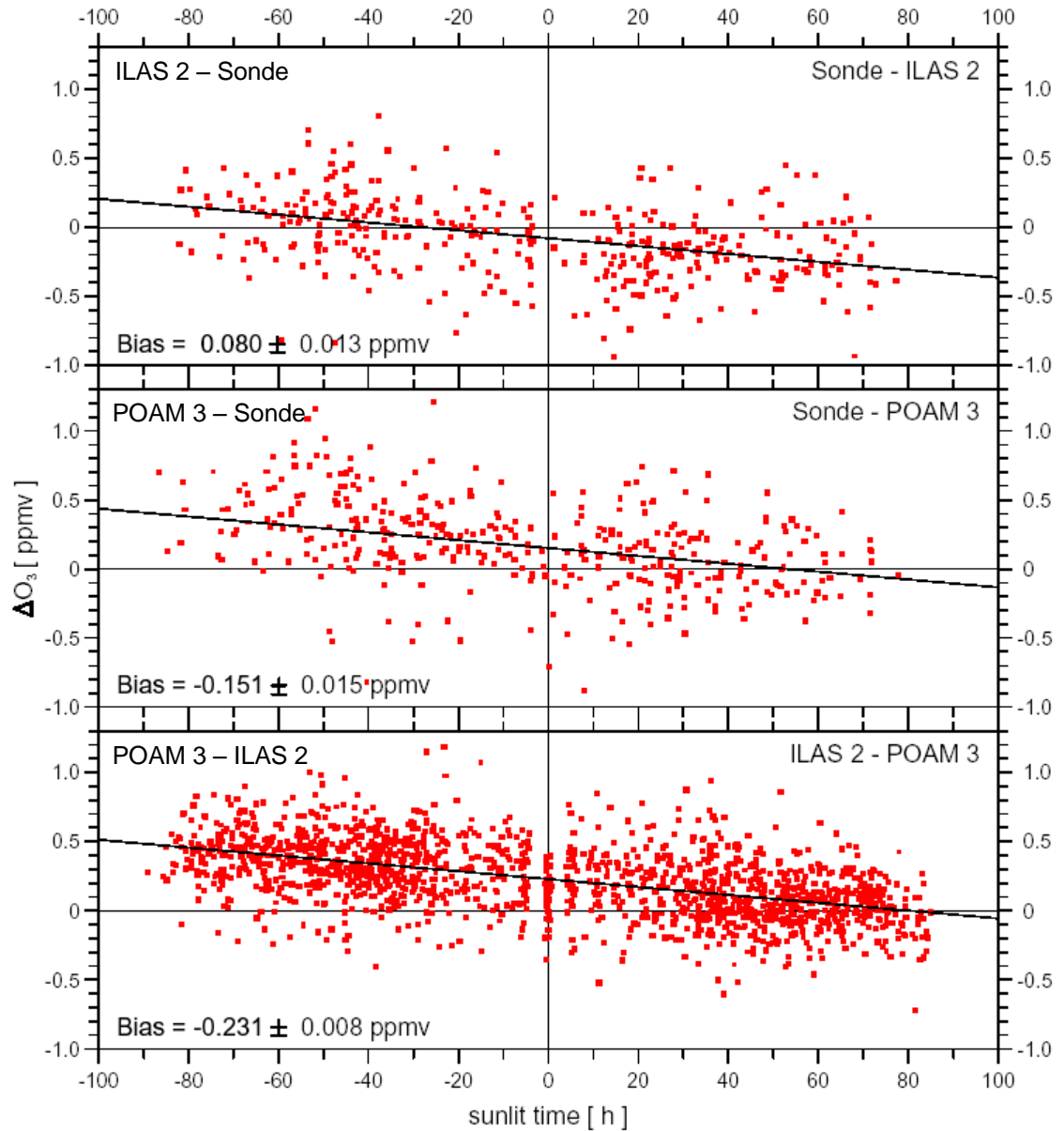
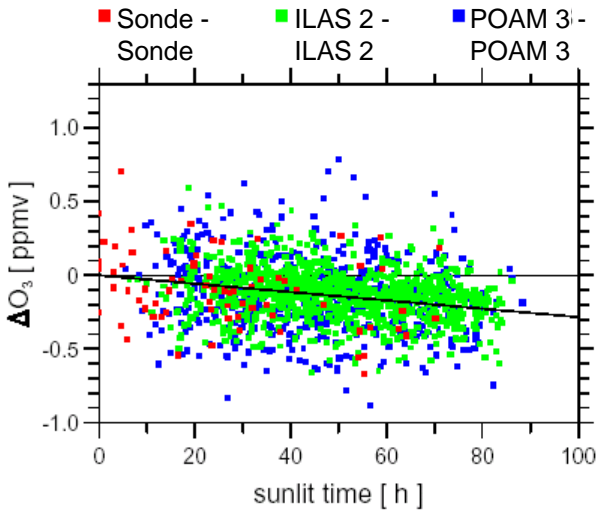
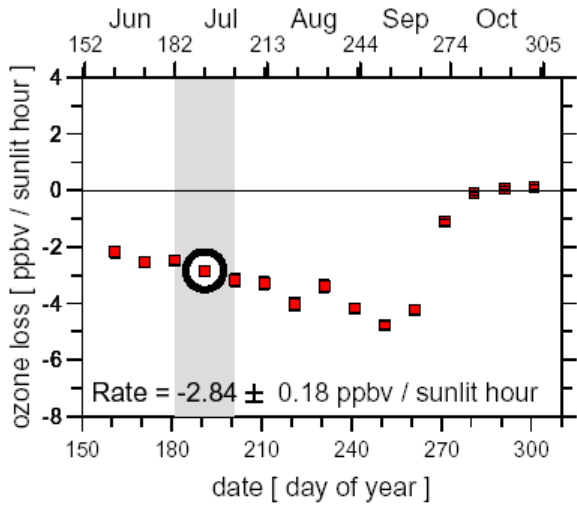
465 K - 485 K



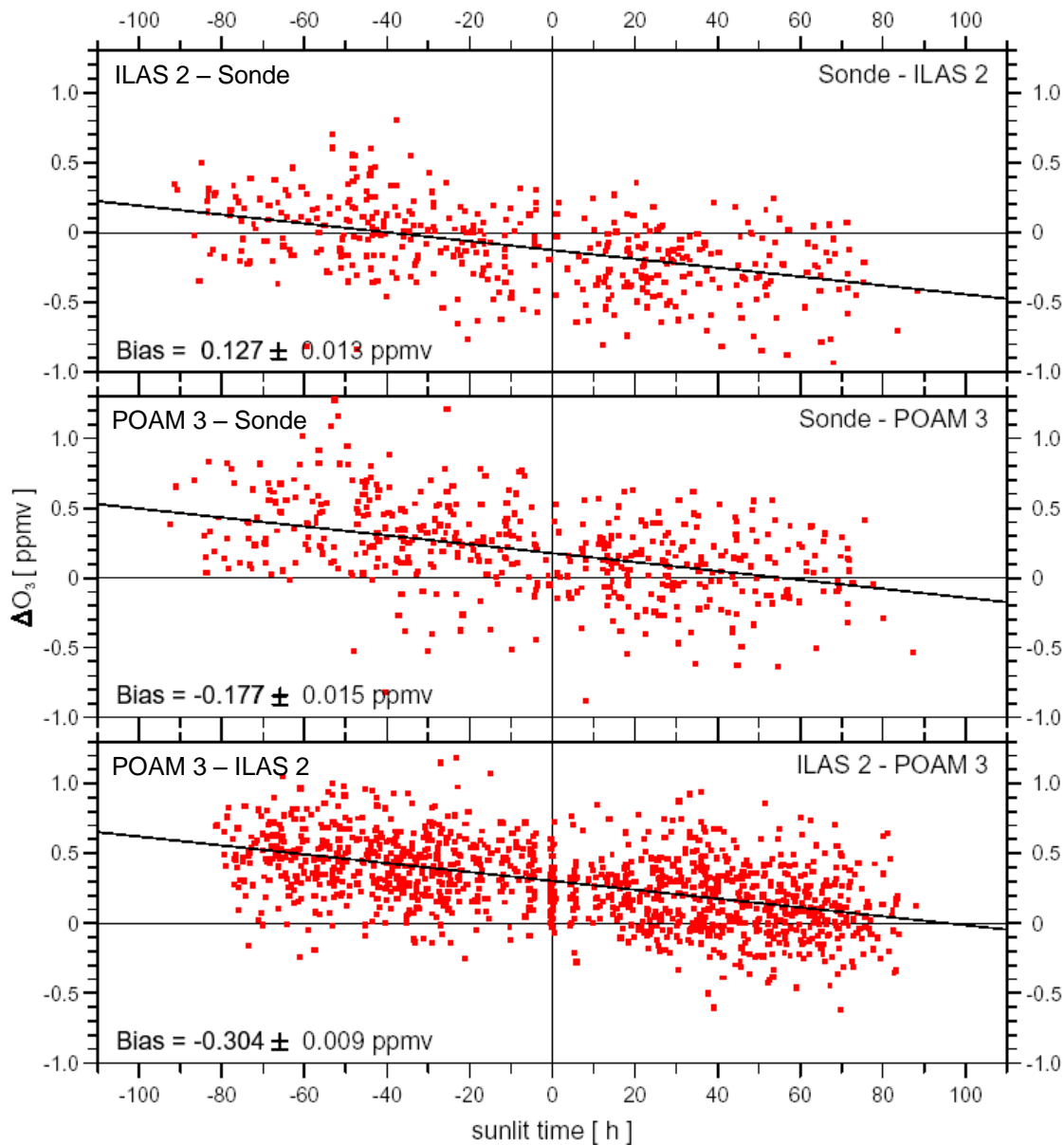
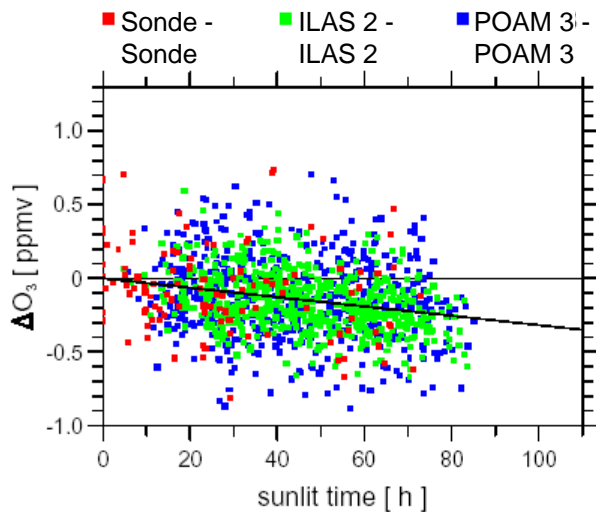
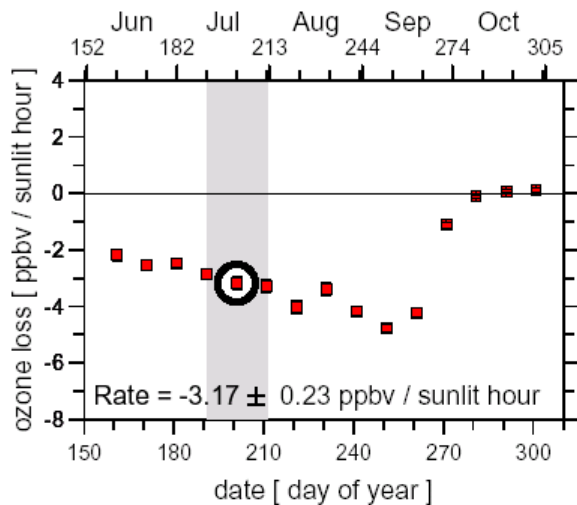
465 K - 485 K



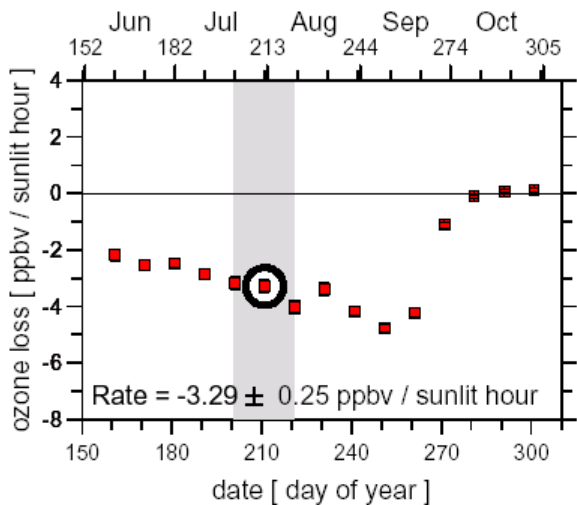
465 K - 485 K



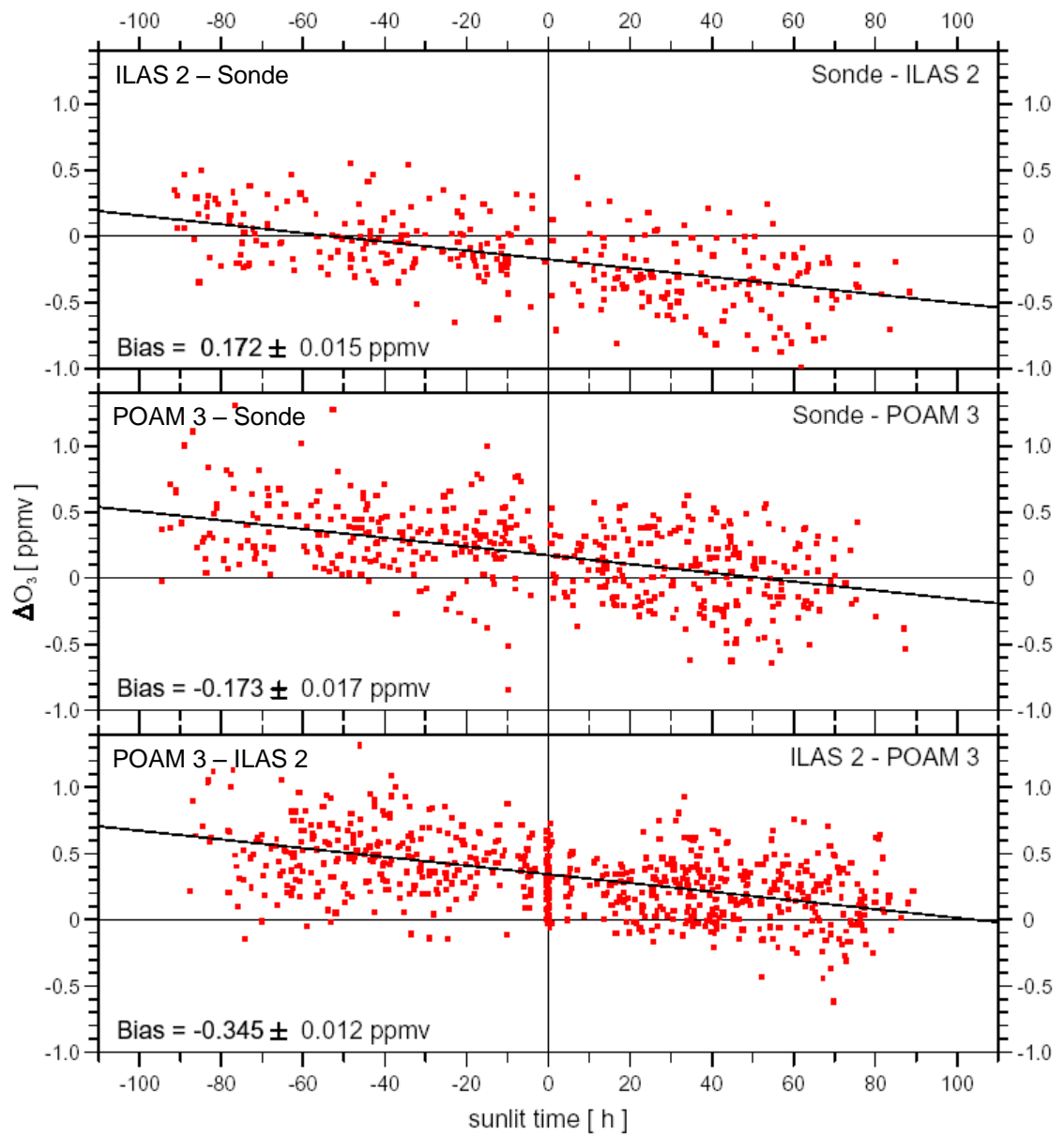
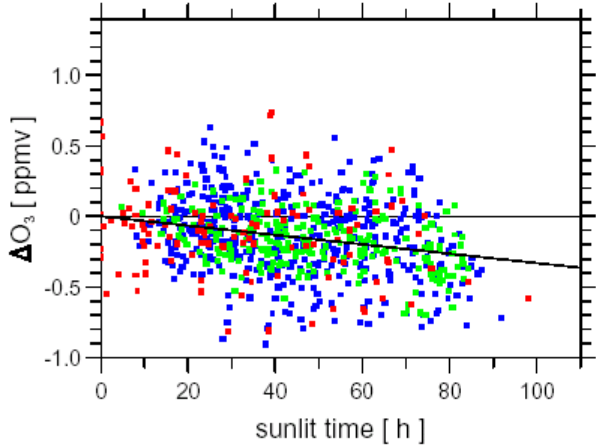
465 K - 485 K



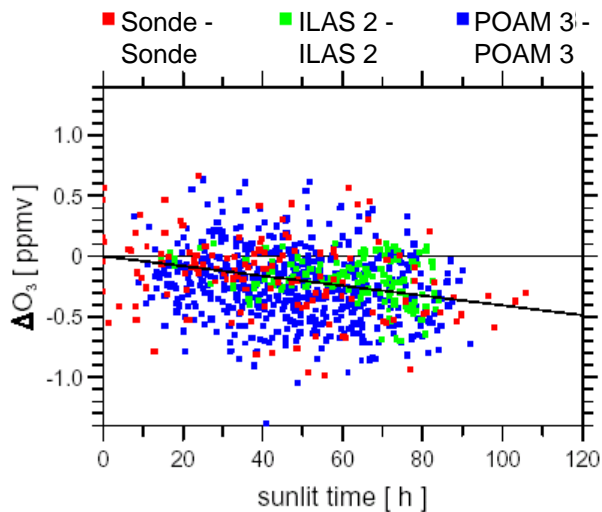
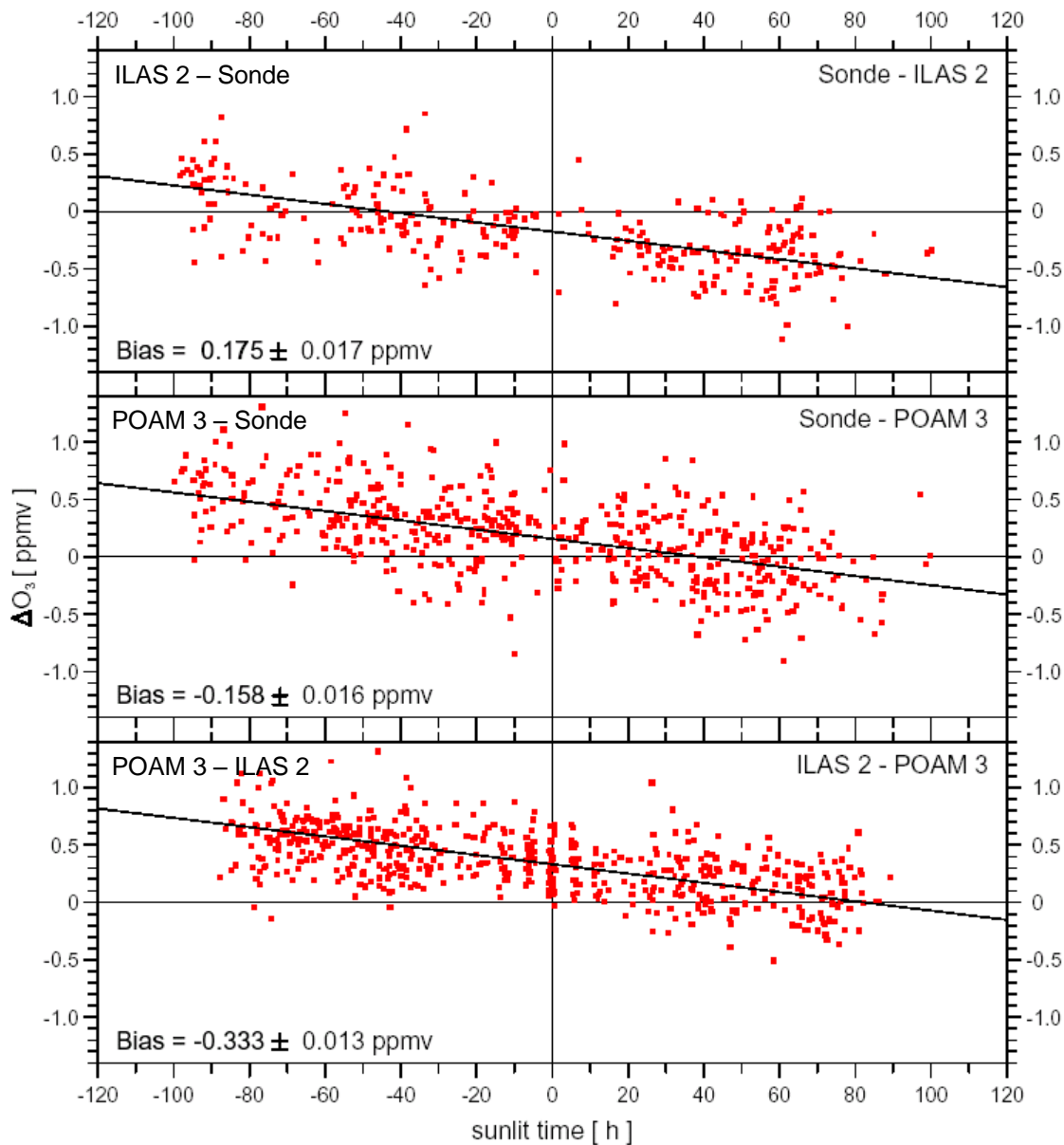
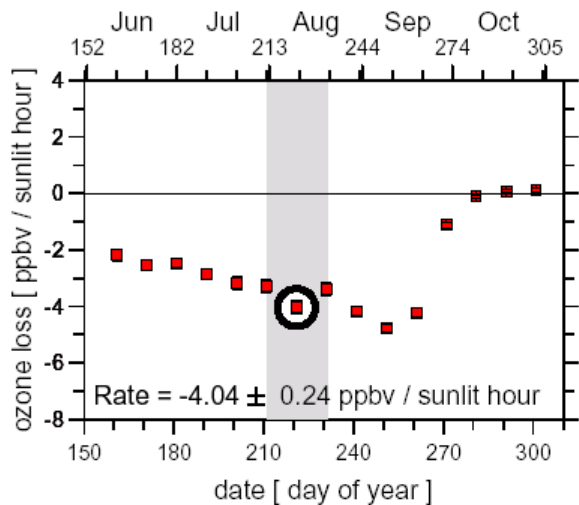
465 K - 485 K



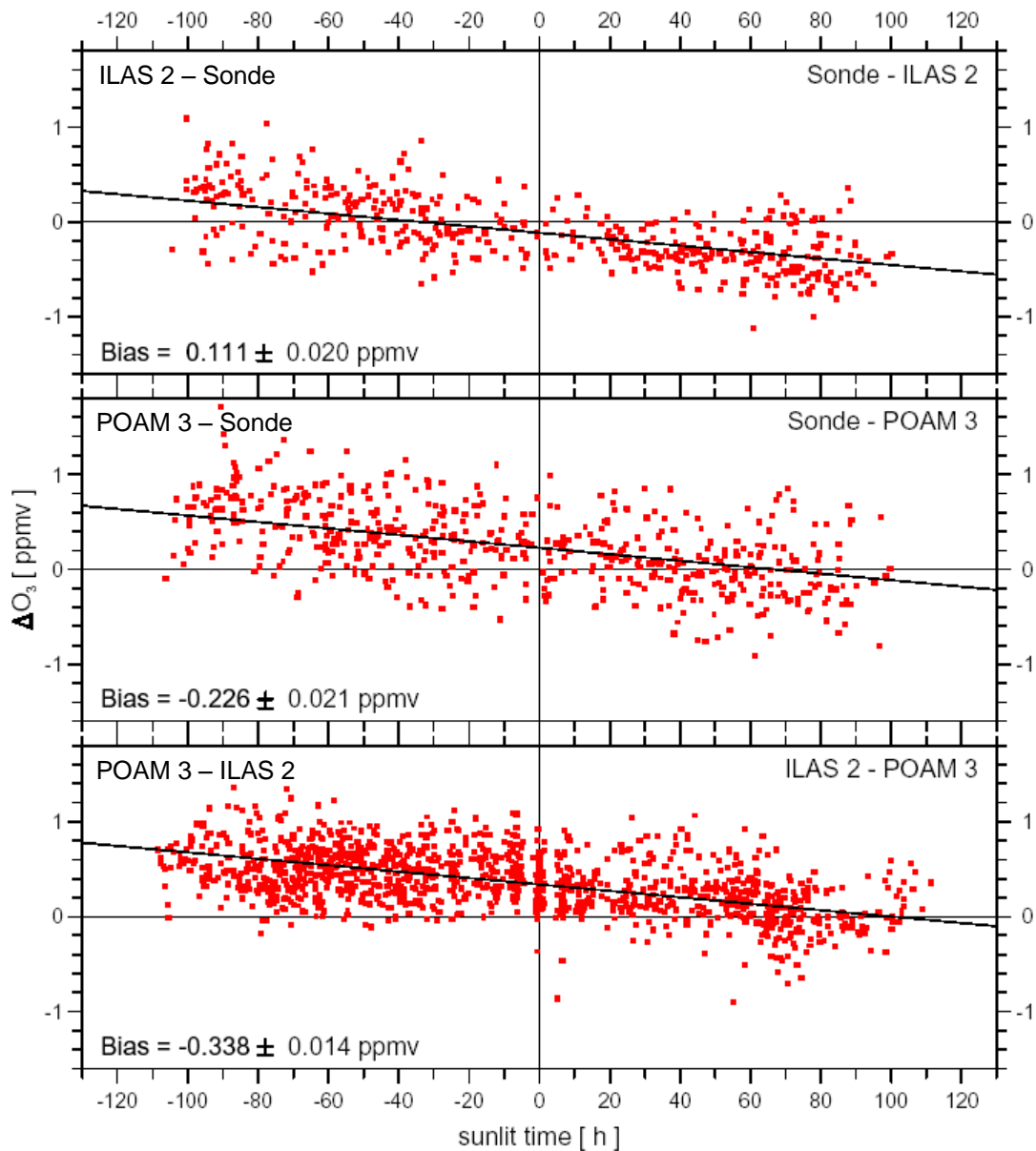
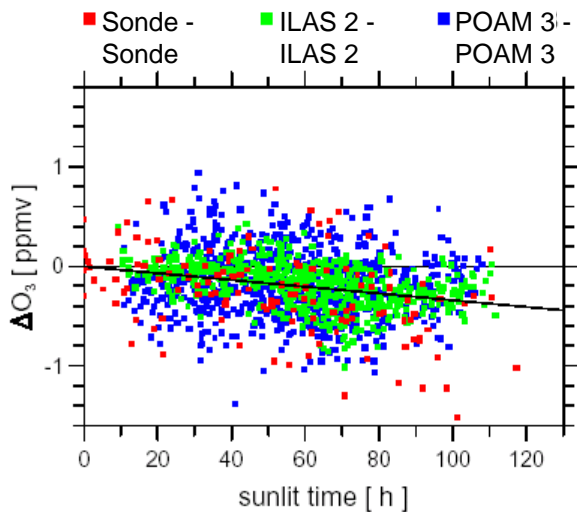
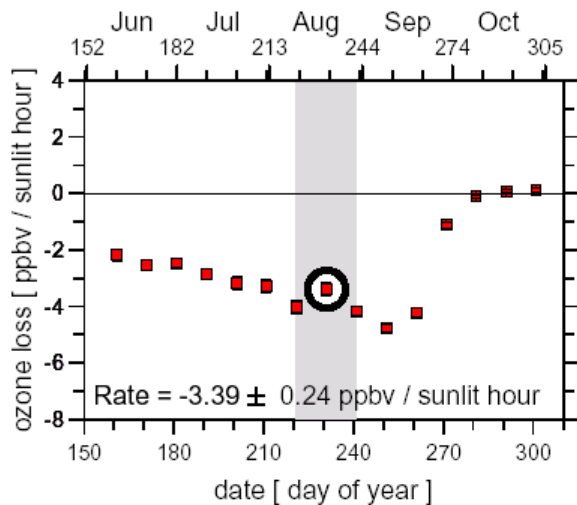
■ Sonde - Sonde
 ■ ILAS 2 - ILAS 2
 ■ POAM 3 - POAM 3



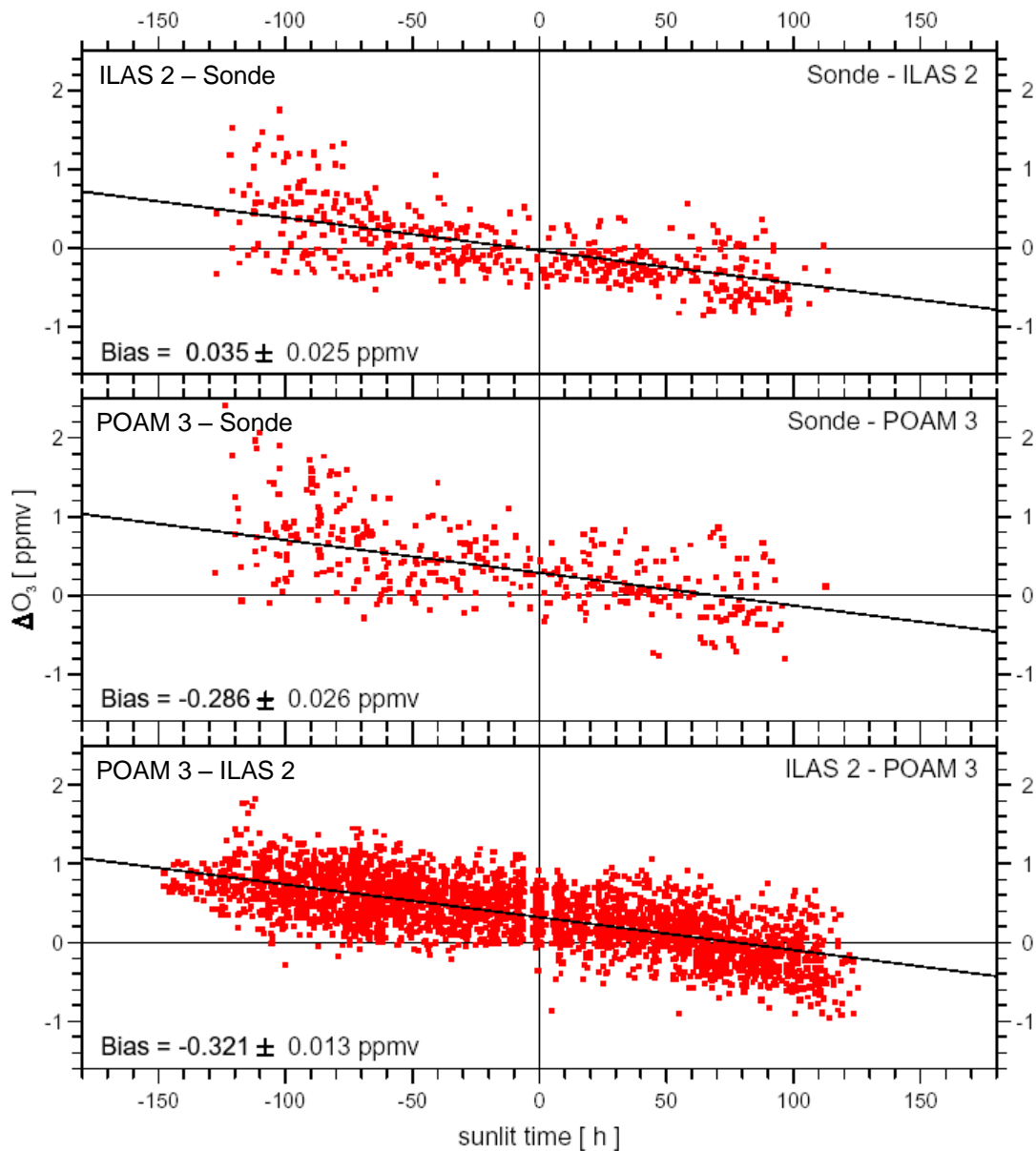
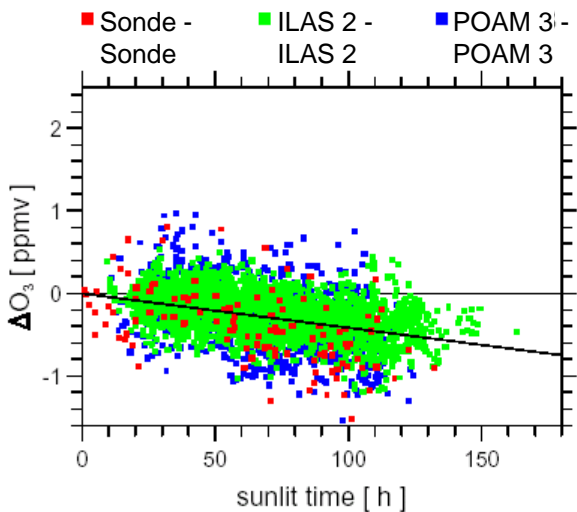
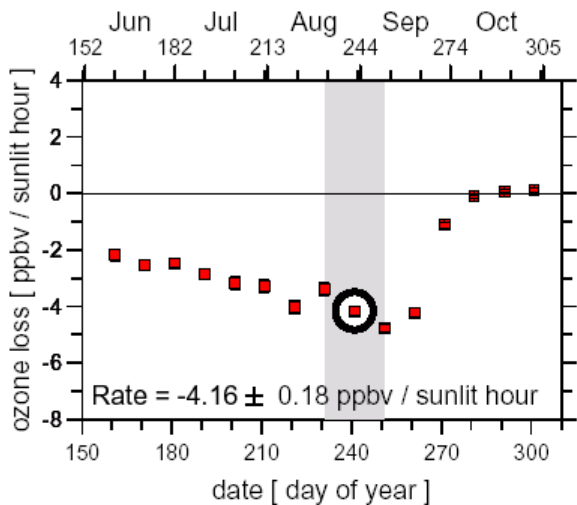
465 K - 485 K



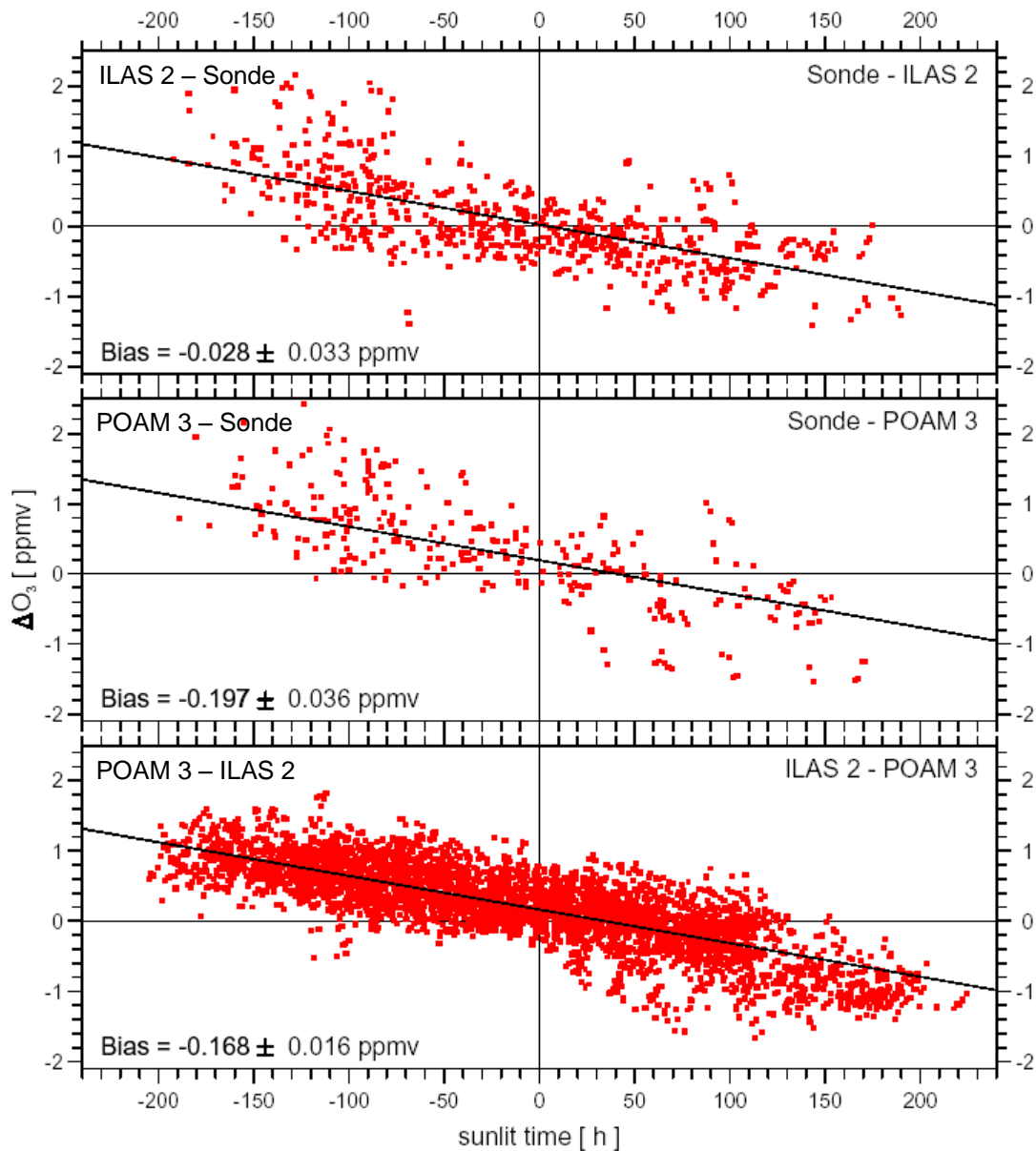
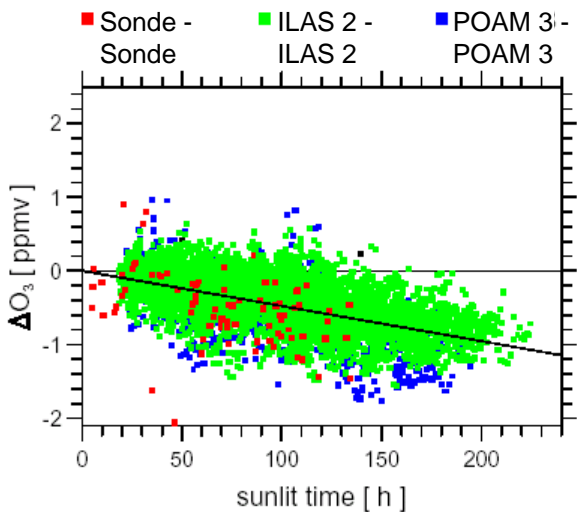
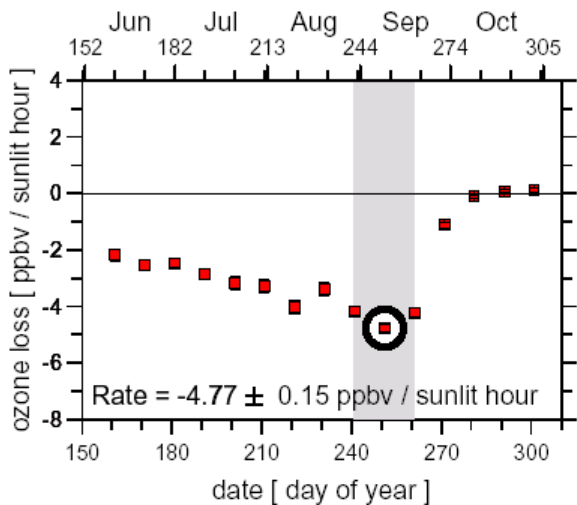
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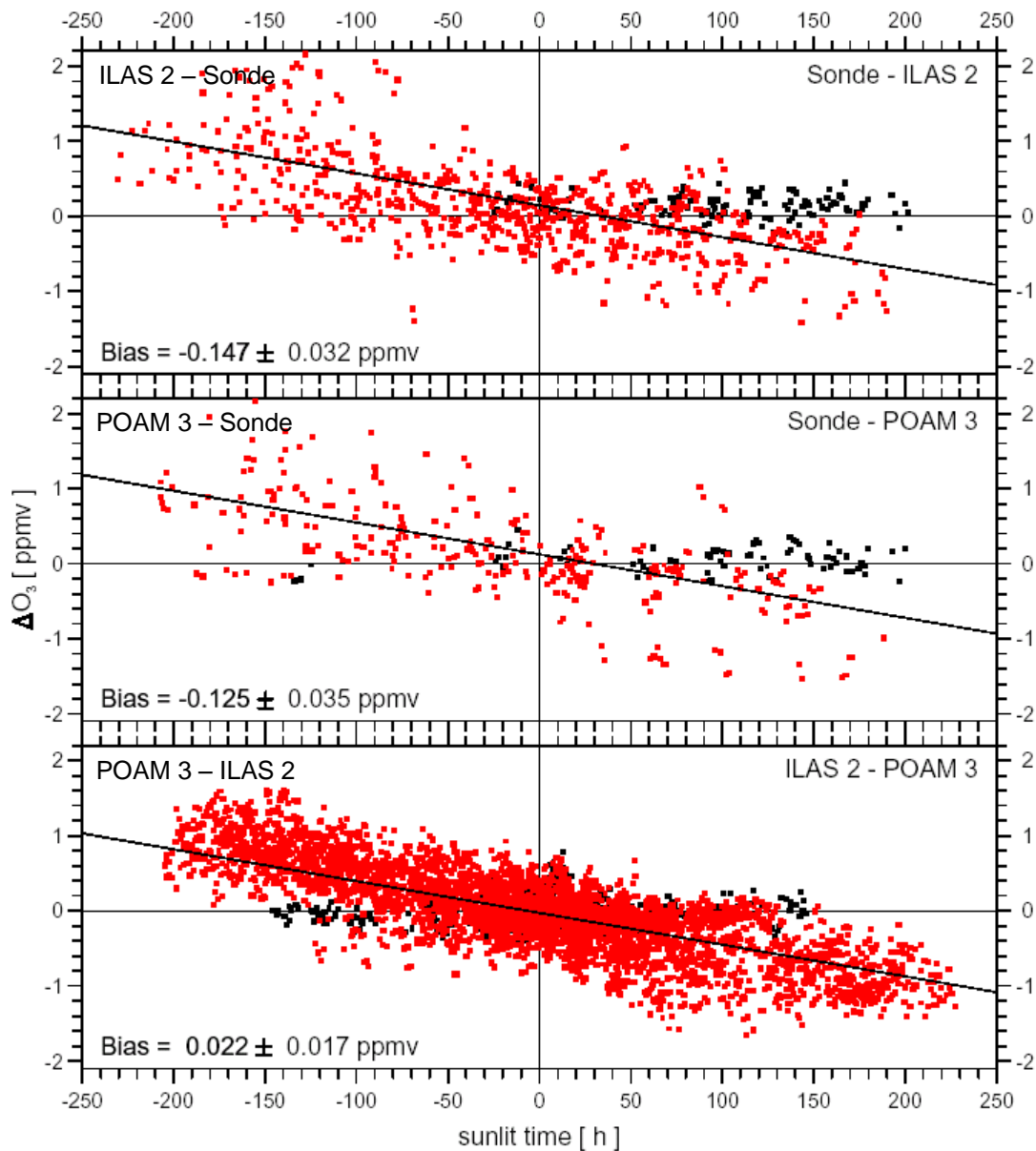
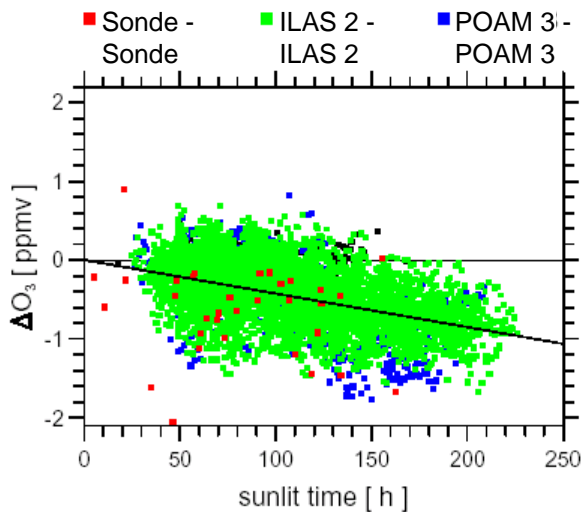
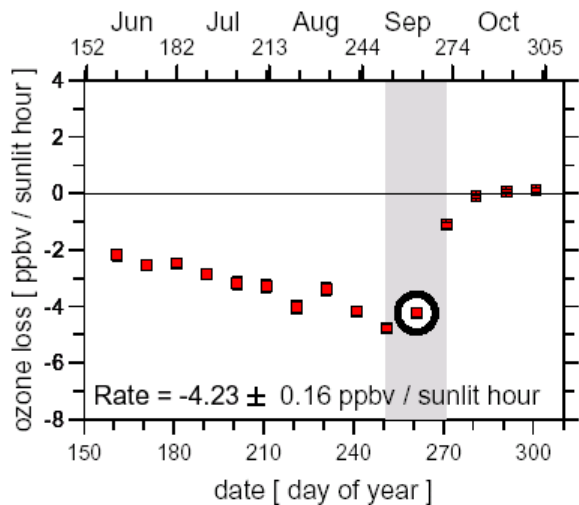
465 K - 485 K



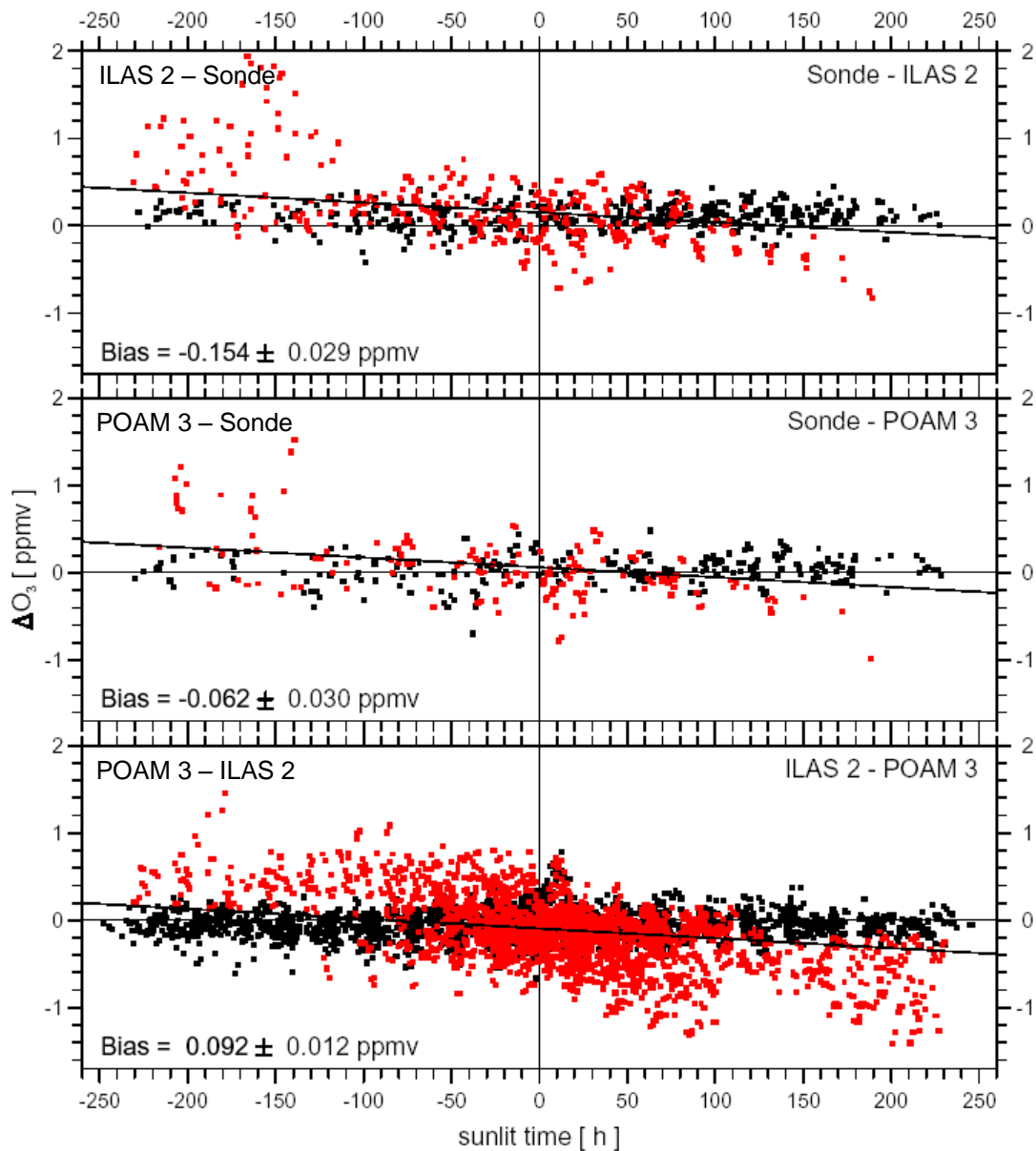
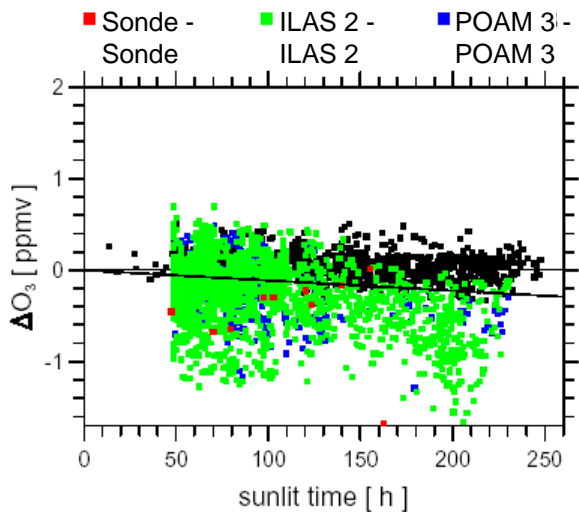
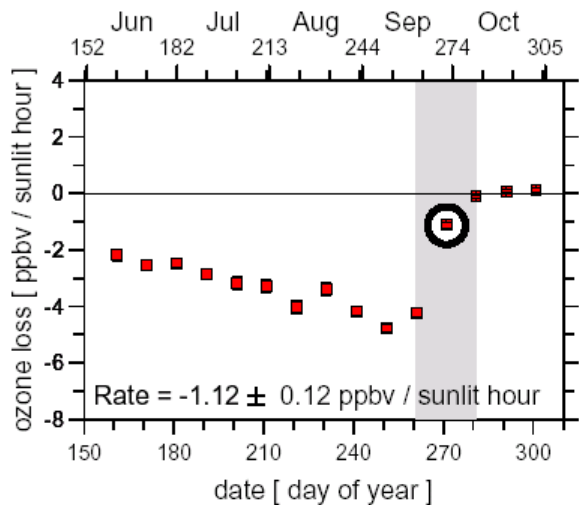
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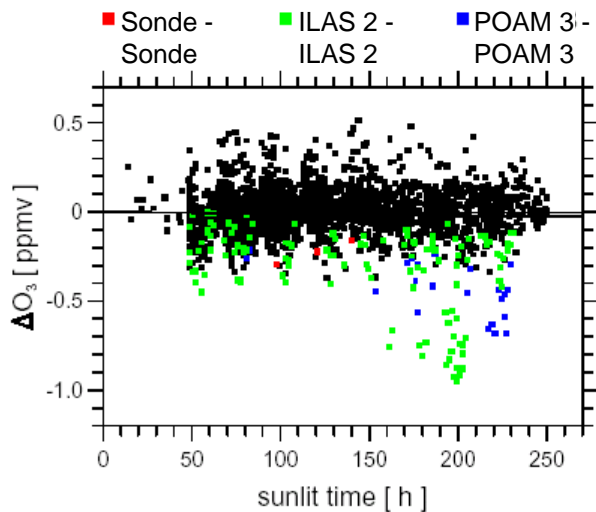
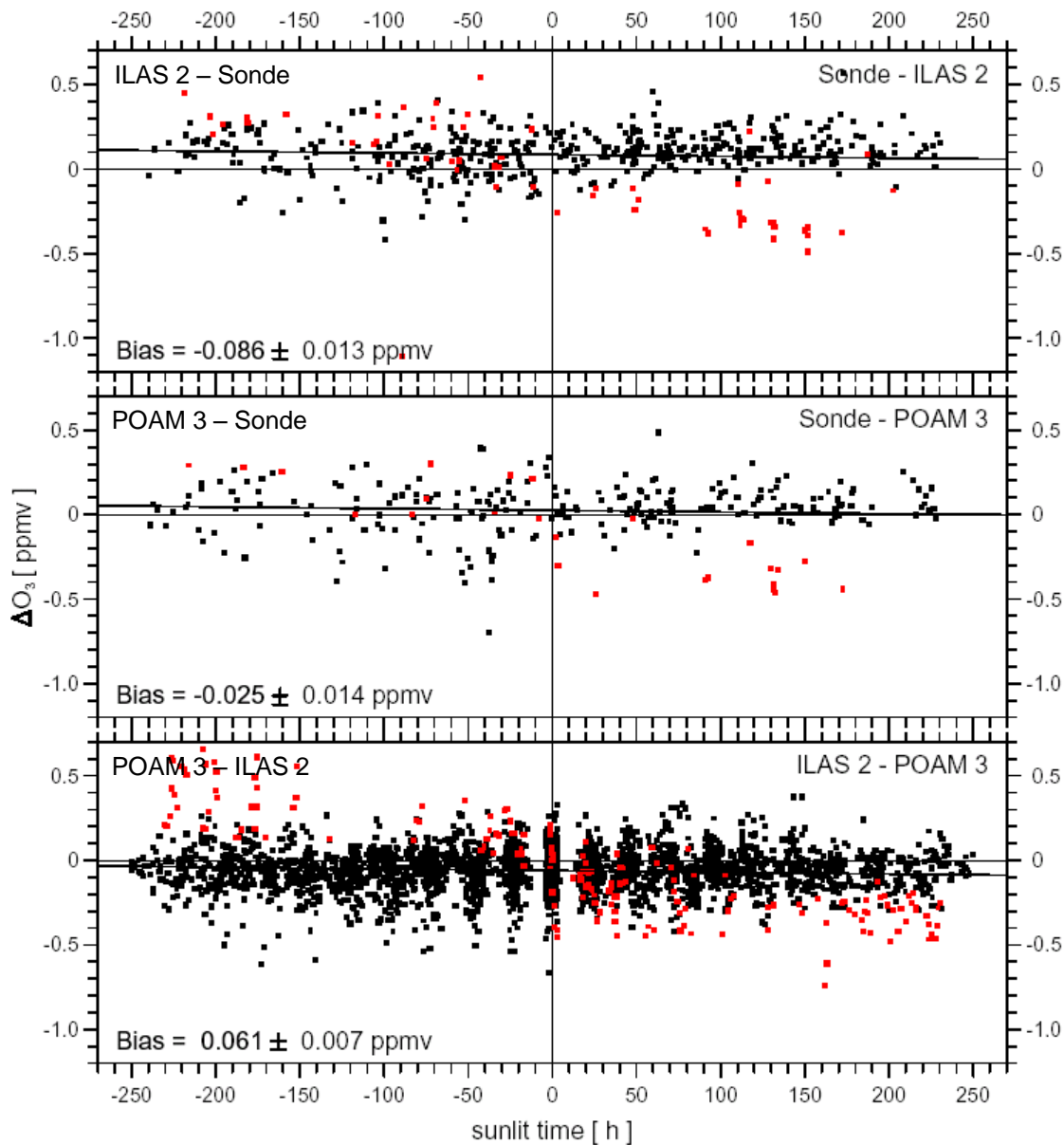
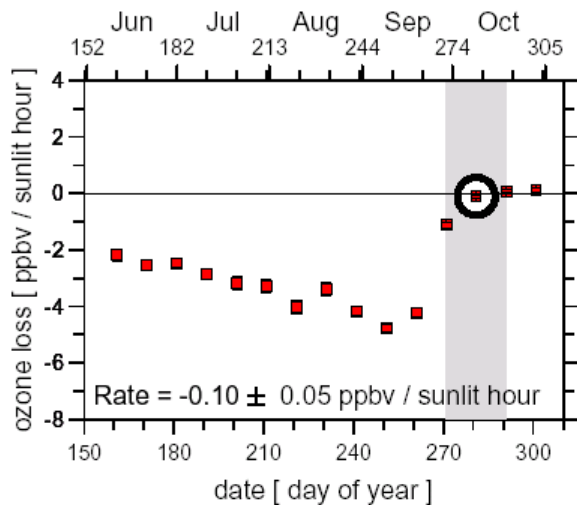
465 K - 485 K



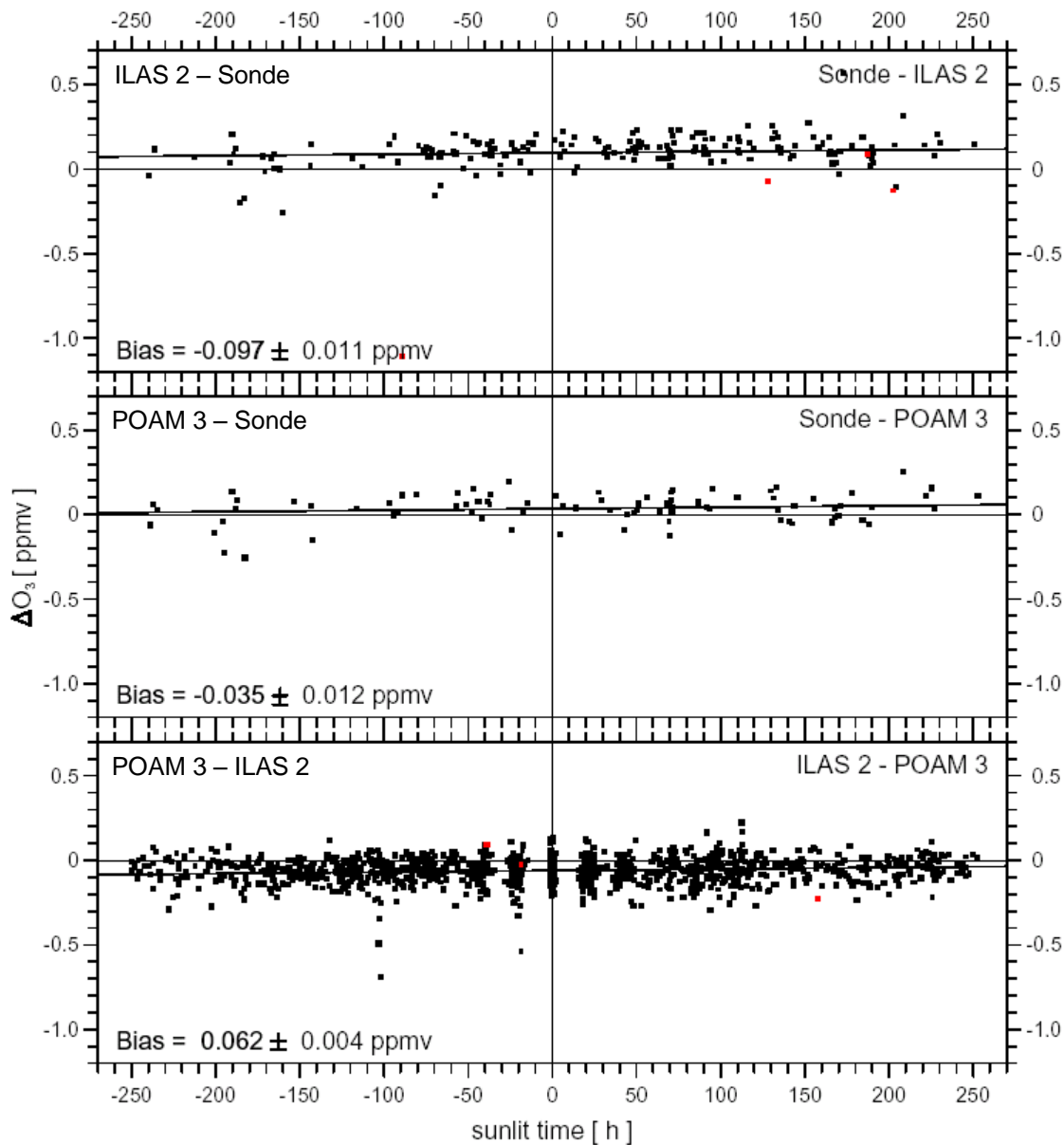
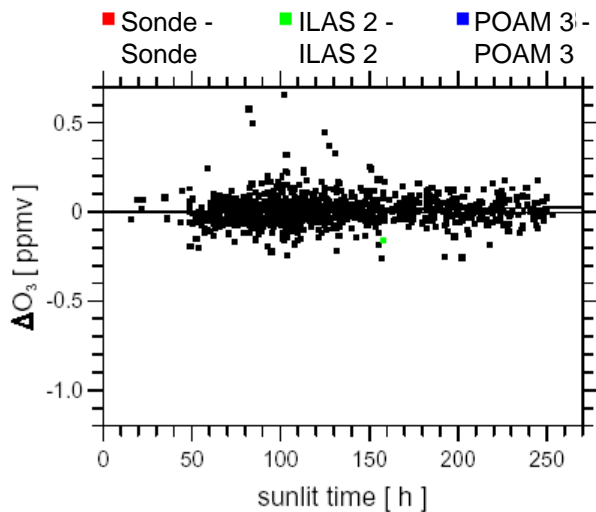
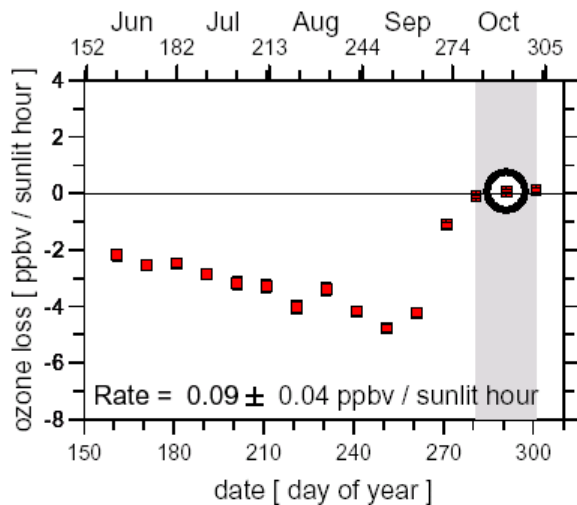
465 K - 485 K



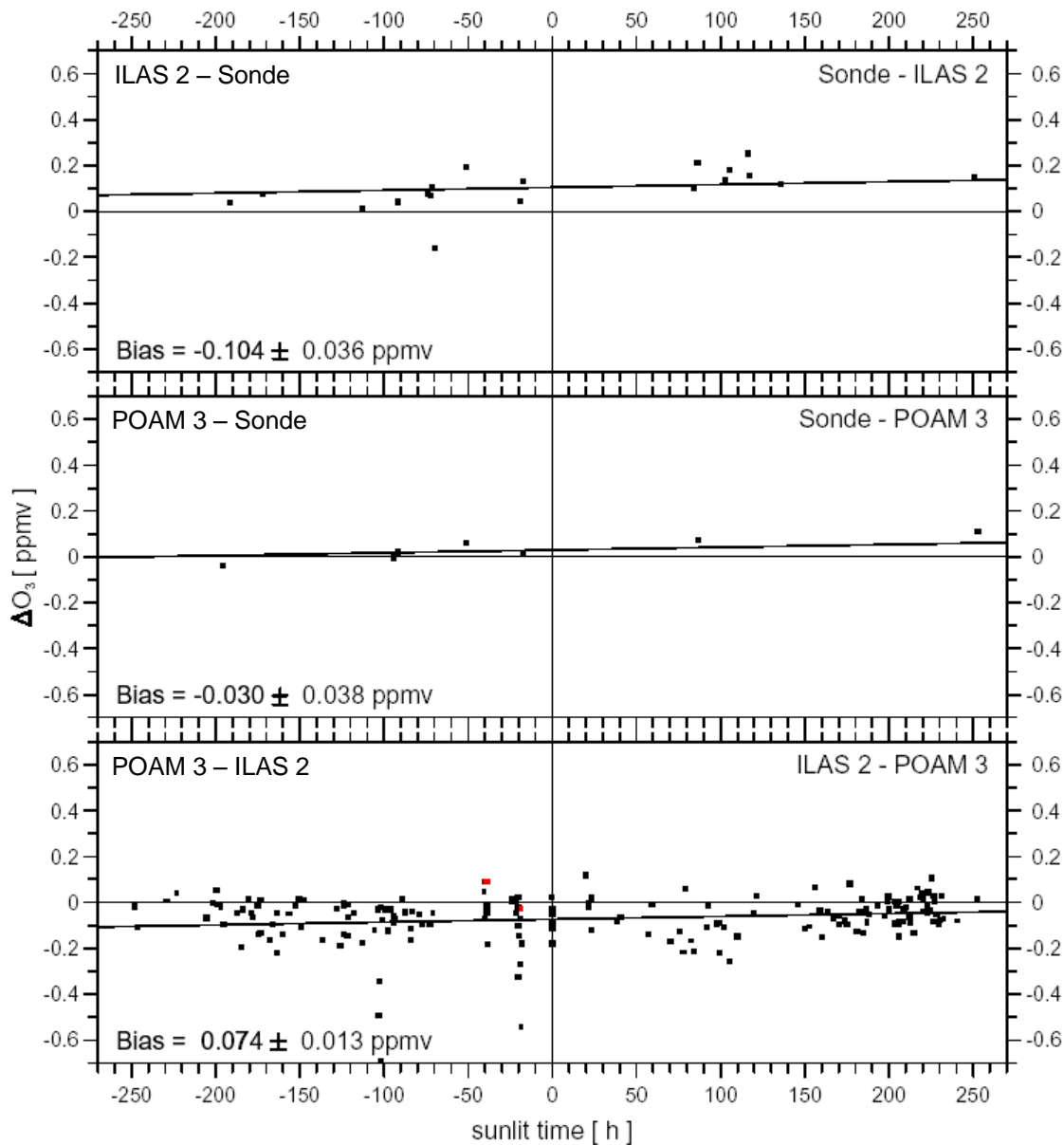
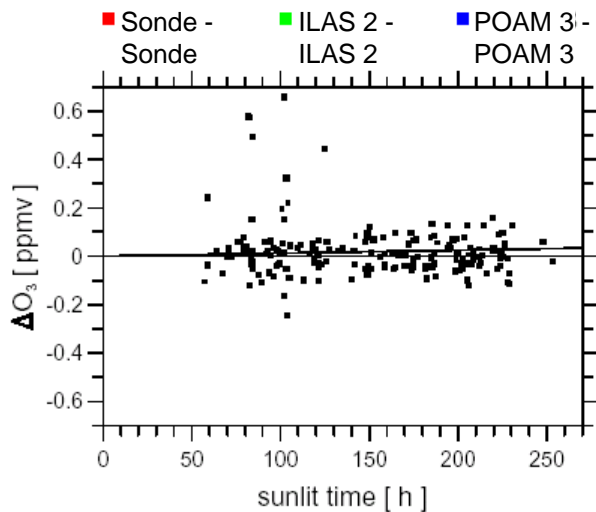
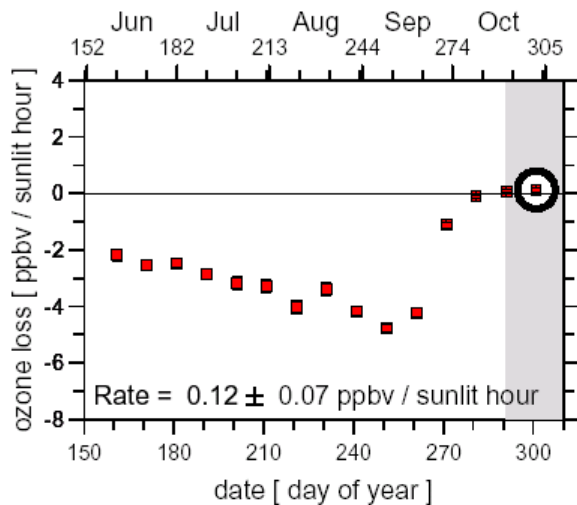
465 K - 485 K



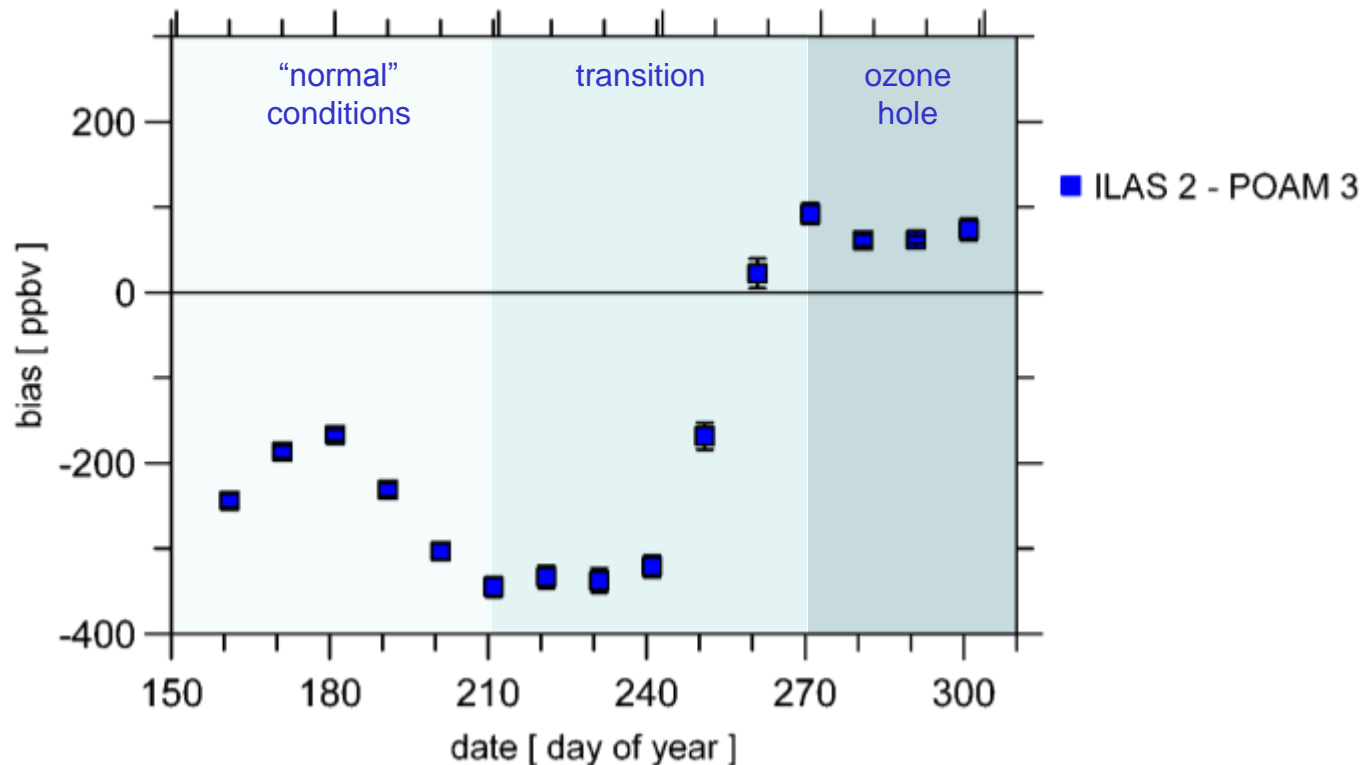
465 K - 485 K



465 K - 485 K

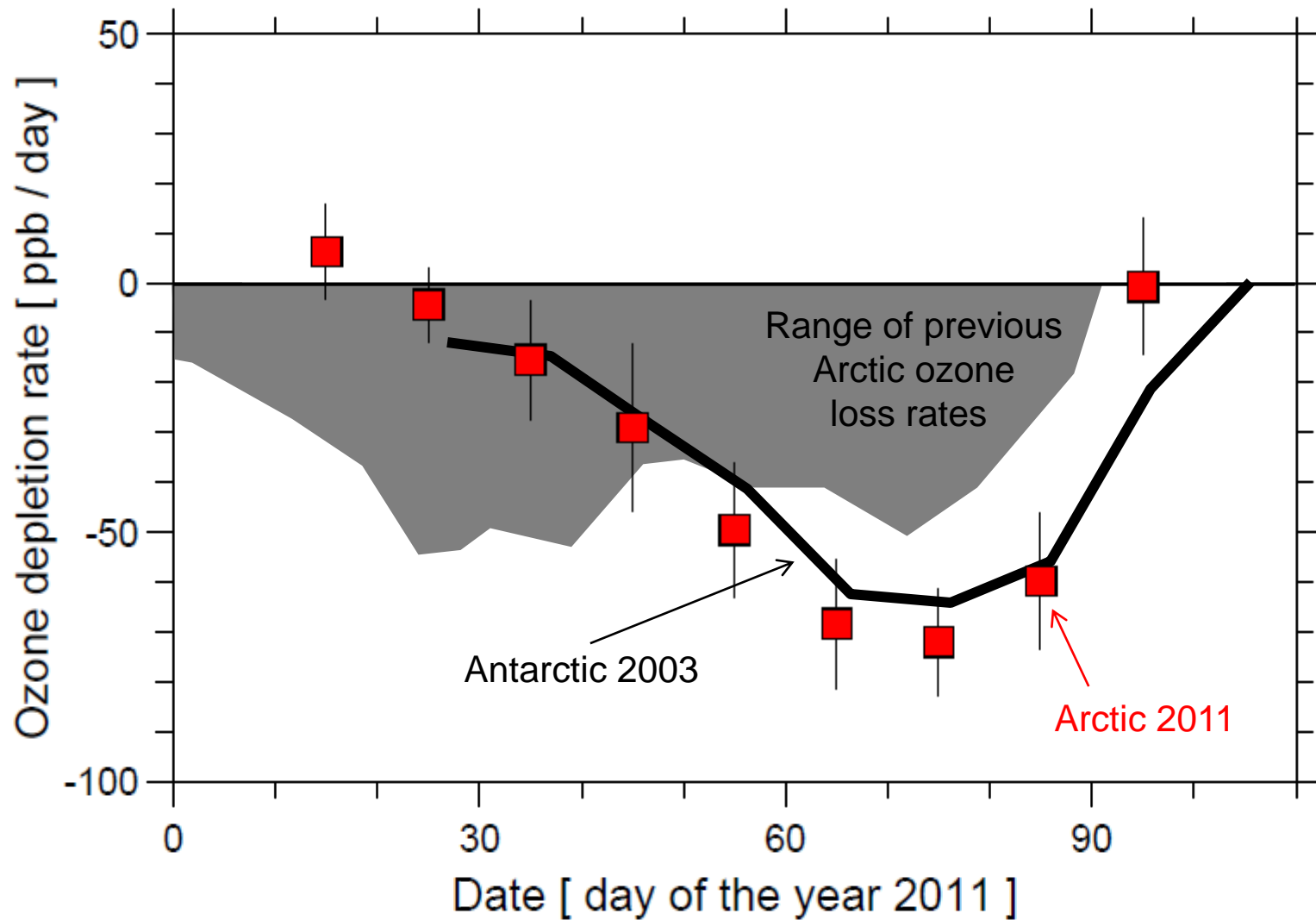


Biases of ILAS2 versus POAM 3

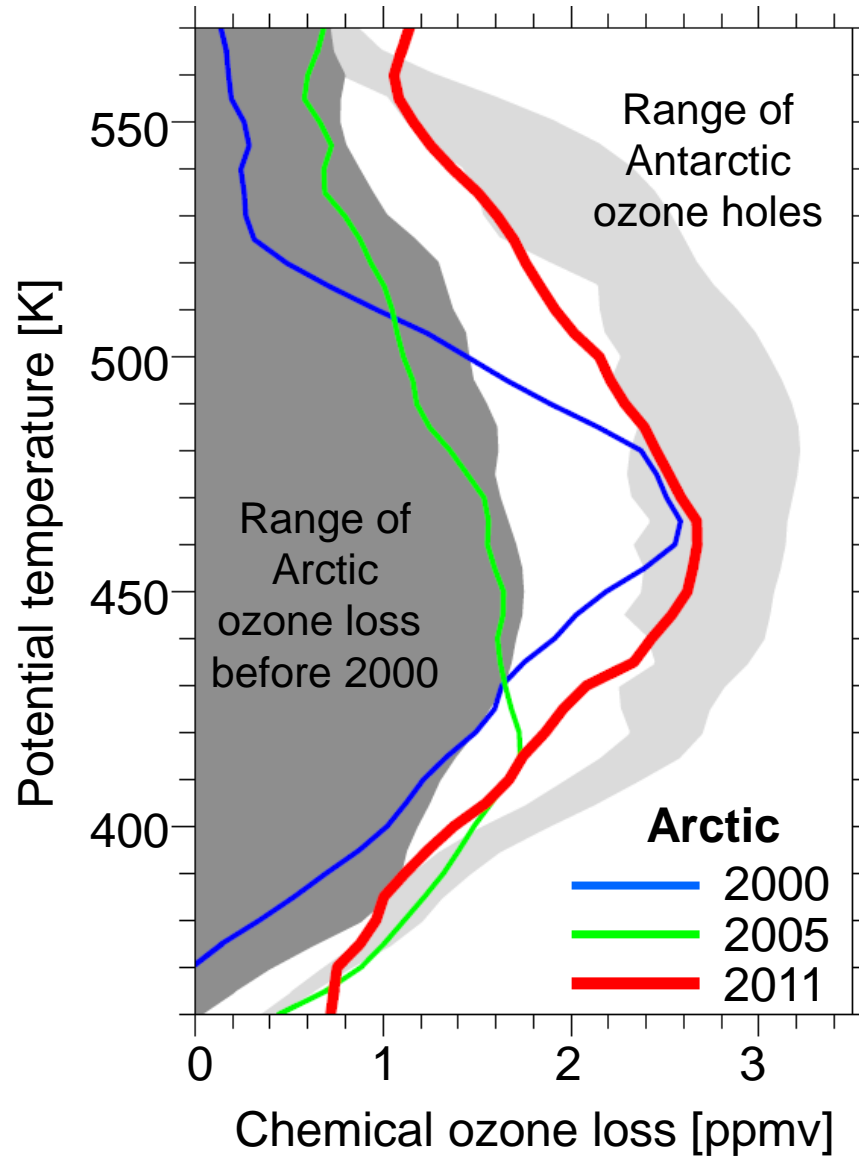


For background conditions ILAS II v1.4 is about 200-350 ppbv lower than POAM III. For very low ozone under ozone hole conditions it is about 50 ppbv higher.

Match: Chemical ozone loss rates @ e⁺~465K (unmixed vortex air)

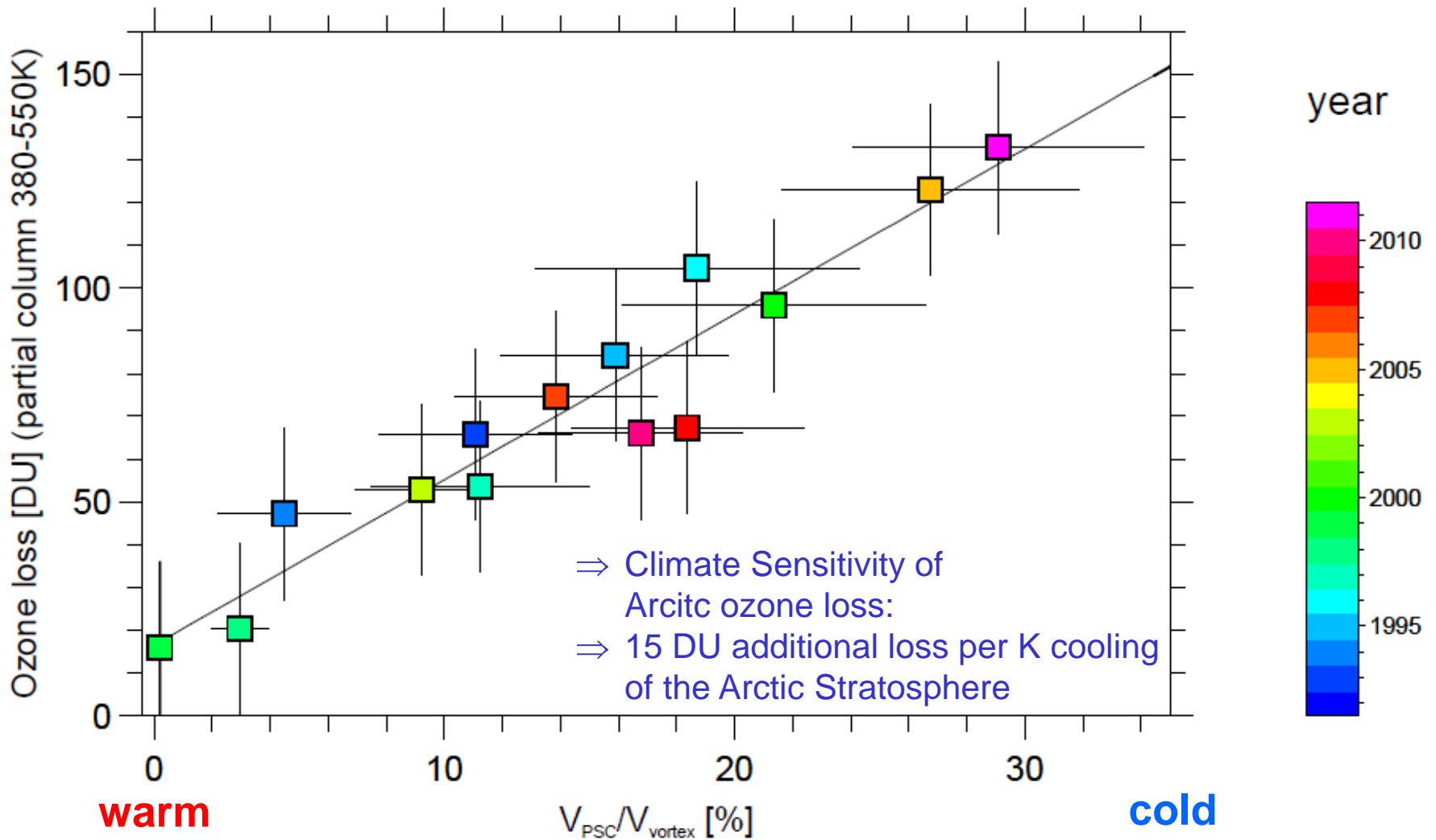


Ozone loss profiles

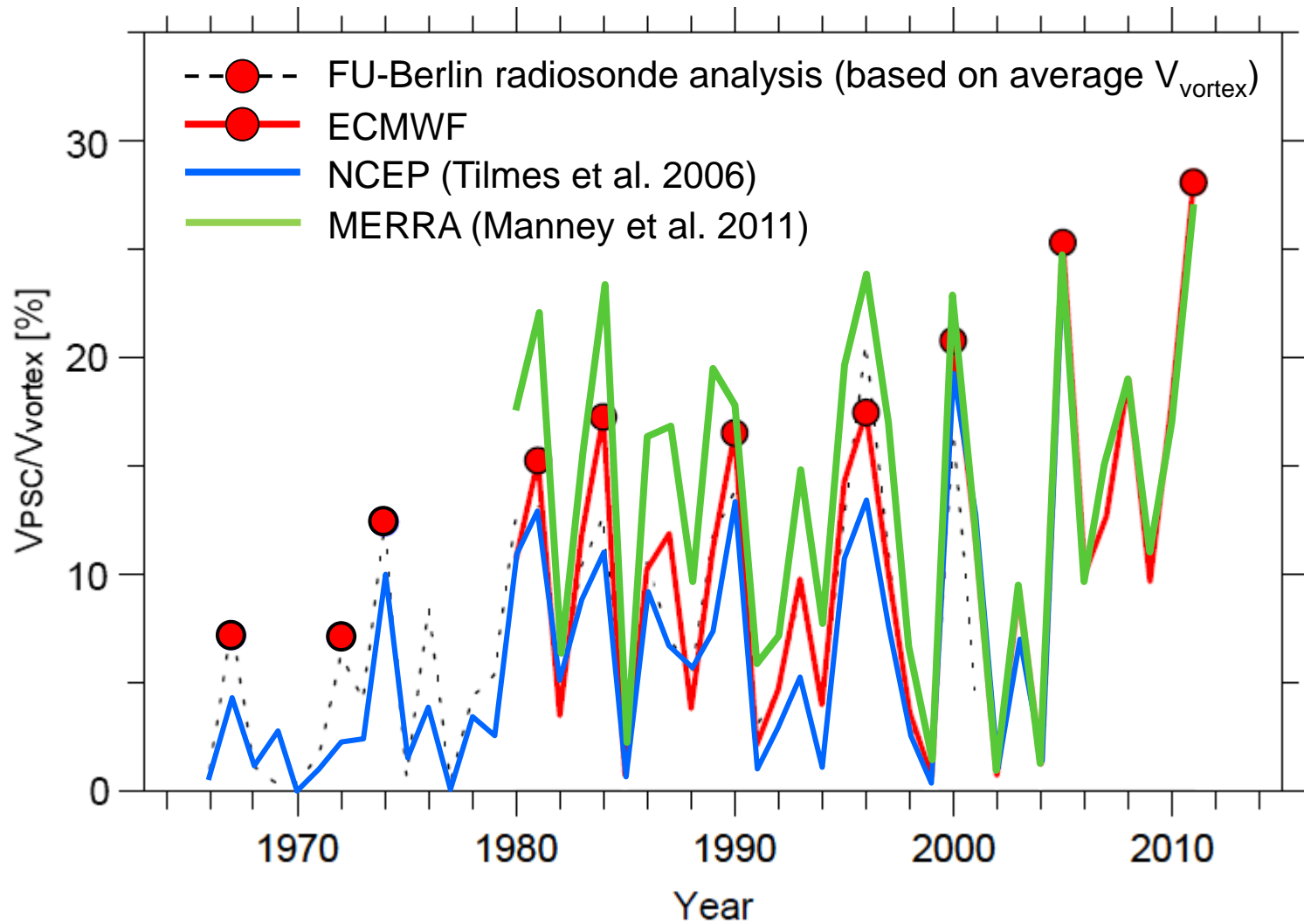


Manney, Santee,
Rex et al., Nature, 2011

Ozone loss versus PSC formation potential (V_{PSC})



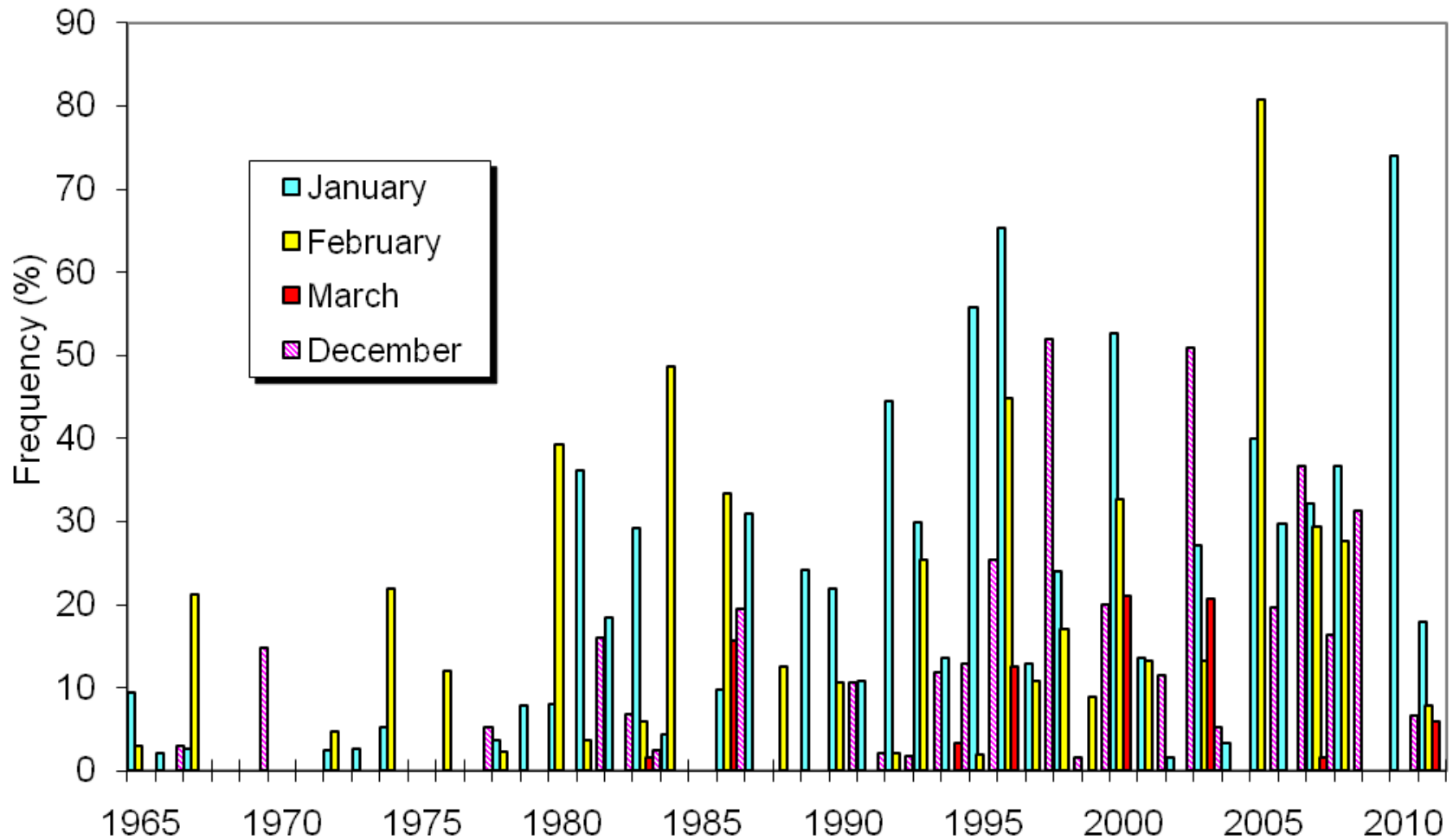
Update of Rex et al., 2004; Rex et al., 2006; WMO 2007; WMO 2011



Update of Rex et al., 2004; 2006; WMO 2007; 2011

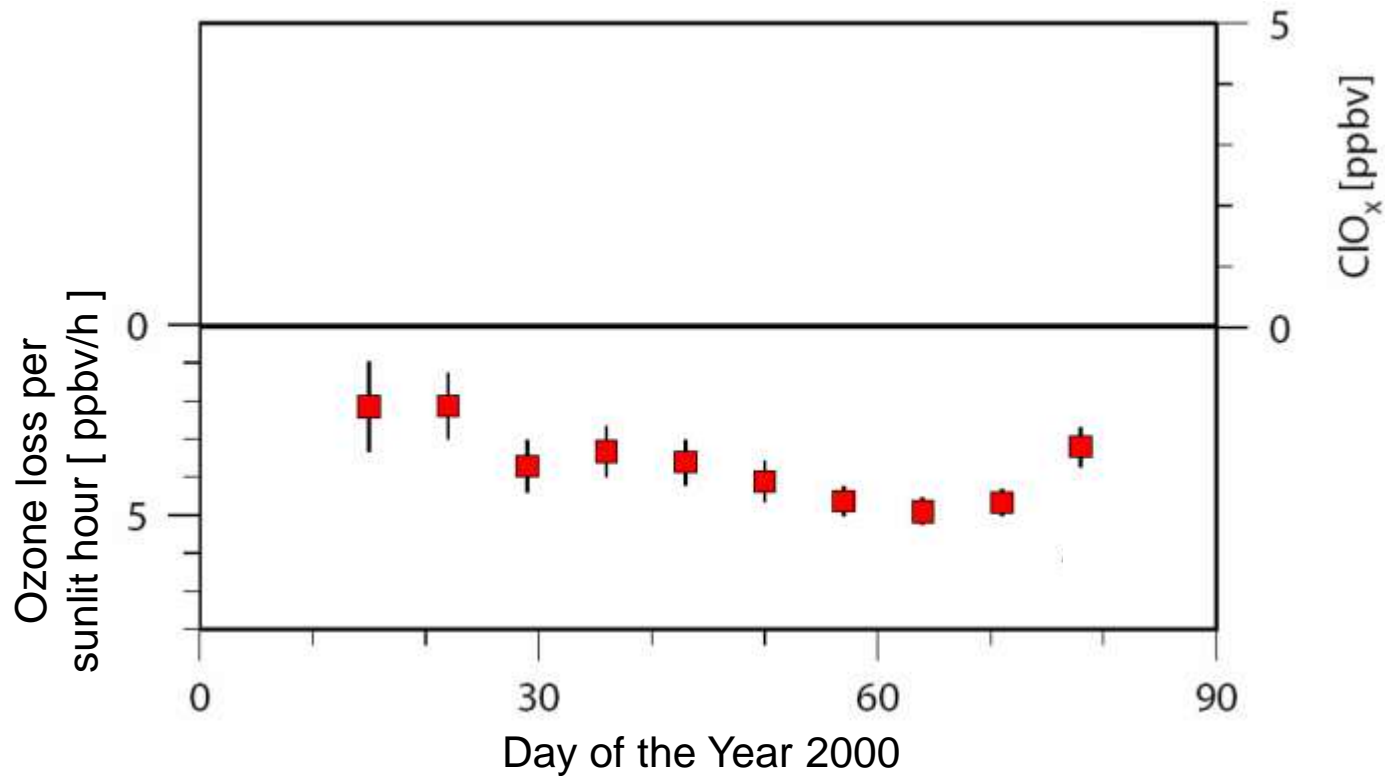
$T < T_{NAT}$ occurrence frequency over Sodankylä

50hPa, based on radiosonde data



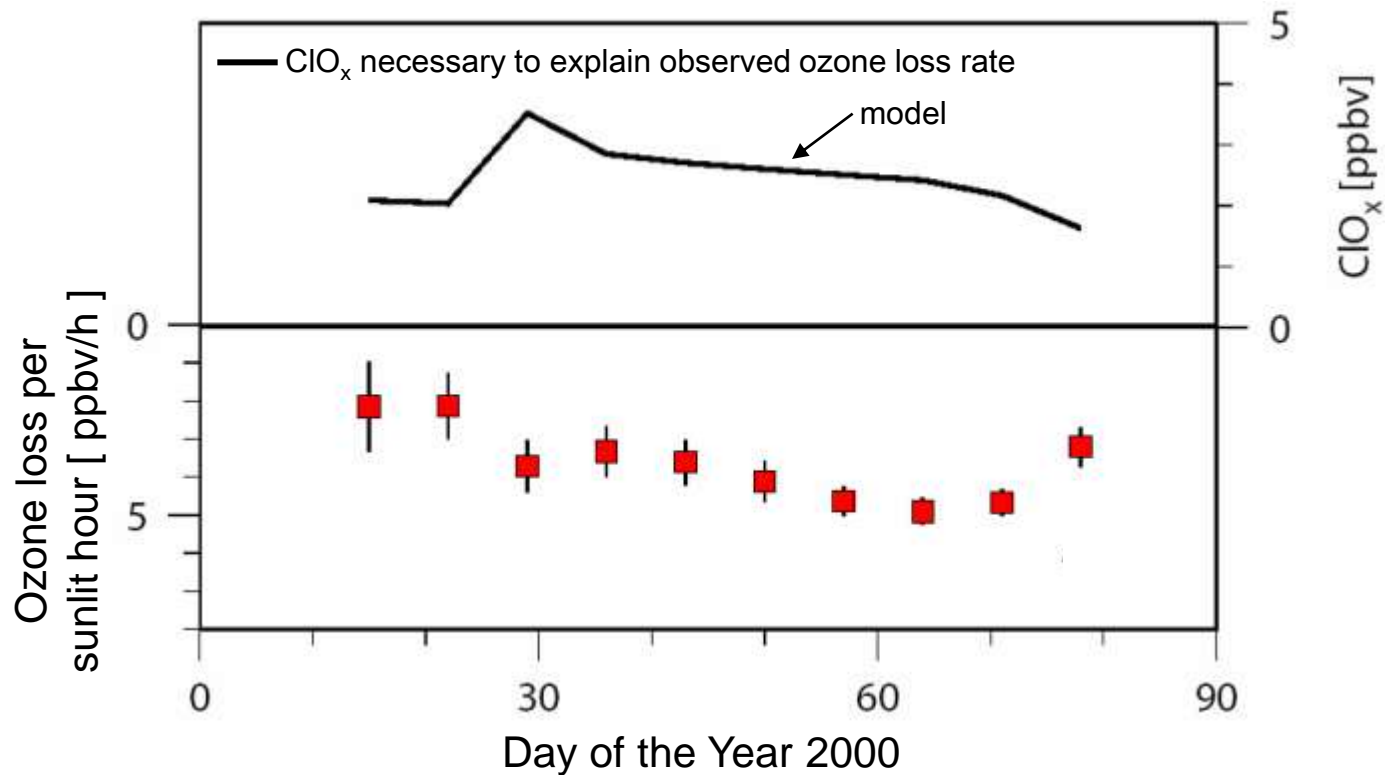
Rigel Kivi, FMI

Theoretical understanding of polar ozone loss process



Frieler et al., GRL 2006; WMO 2007

Theoretical understanding of polar ozone loss process



Frieler et al., GRL 2006; WMO 2006

Geophysika – 2003, 2010

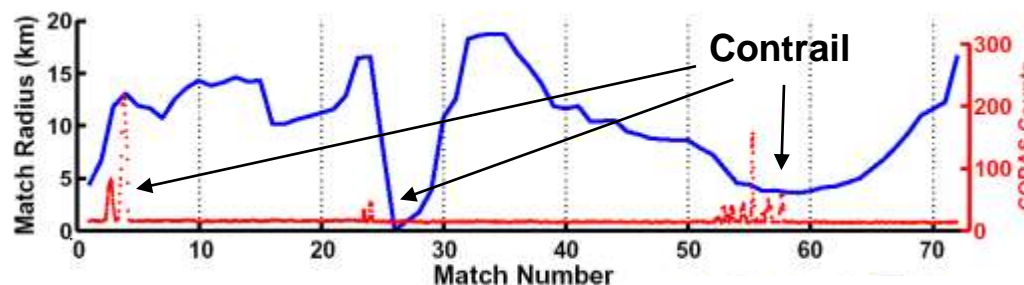
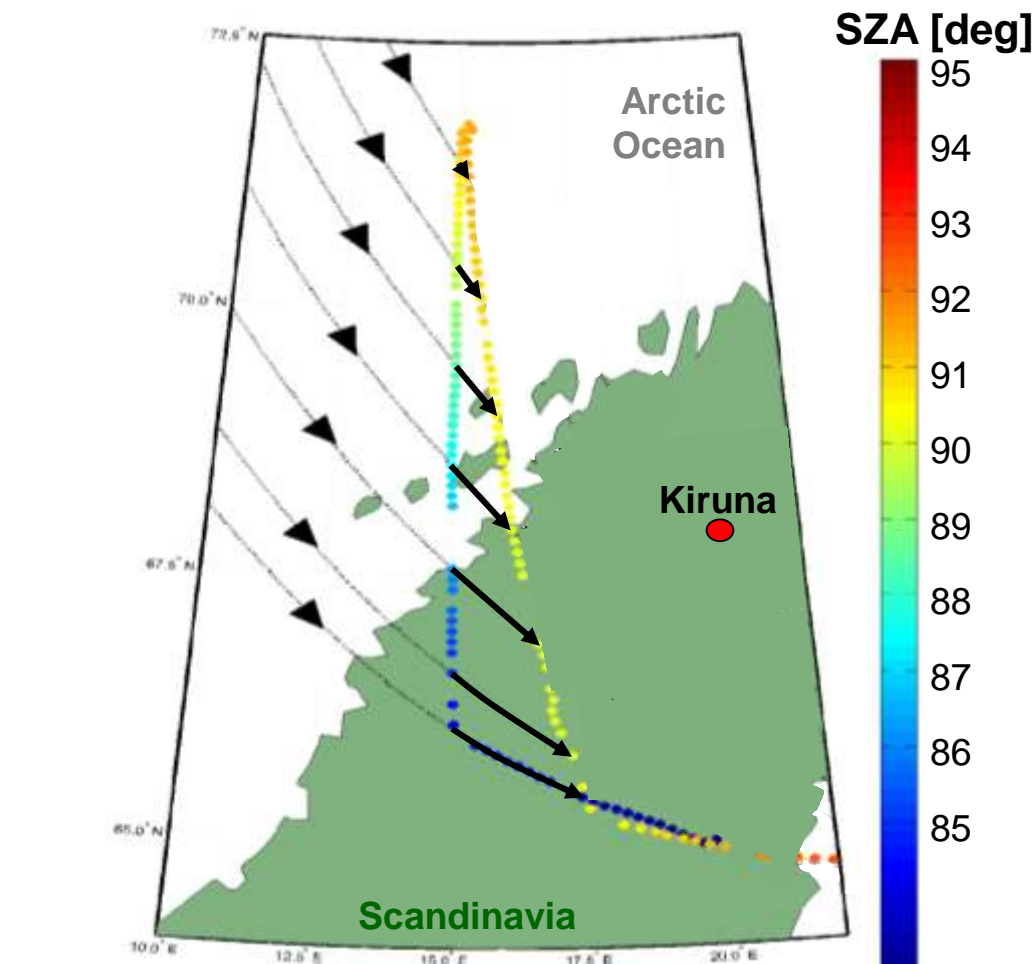


Kiruna, March 2010

Lagrangian aircraft measurements:

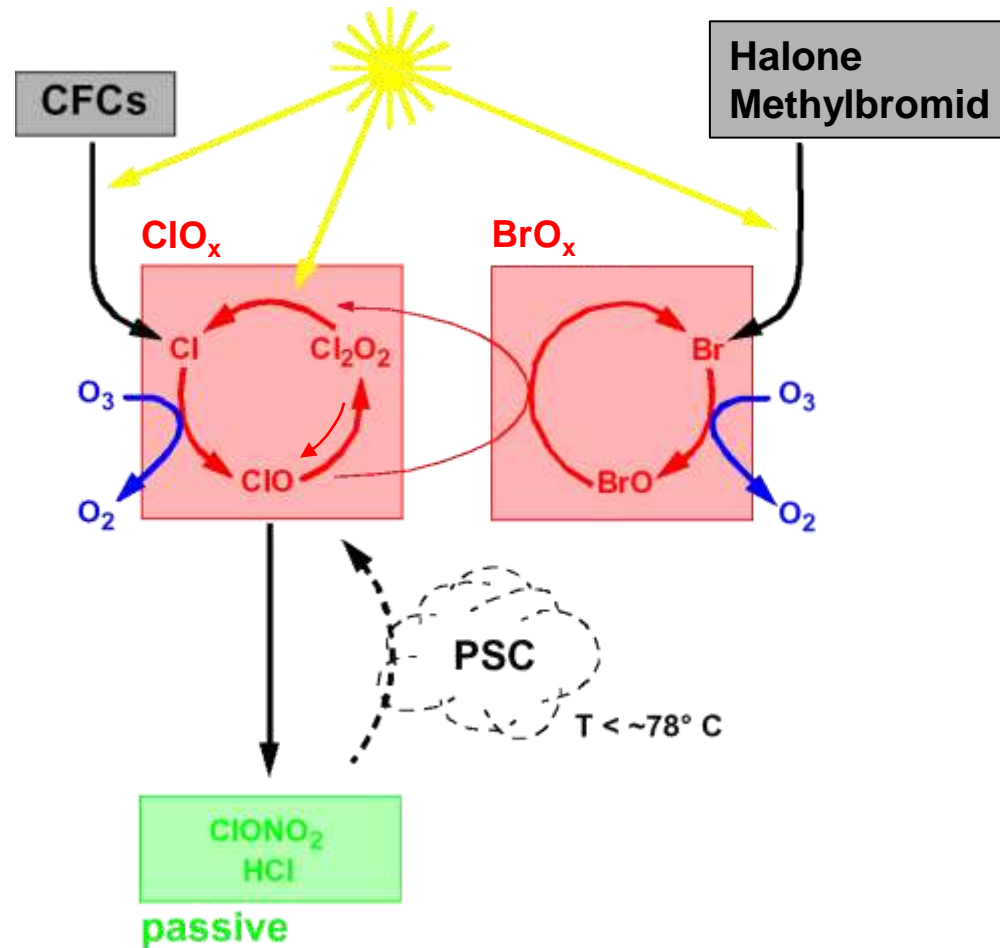
Self-Match-flights

- Flight pattern designed to probe air masses before and after sunset/sunrise
- Success of flight planning confirmed by contrail encounter
- **EUPLEX:**
 - 1 Self-Match flight in 2003
- **RECONCILE: Jan-March 2010**
 - 3 Self-Match flights
 - 1 Match/Self-Match flight

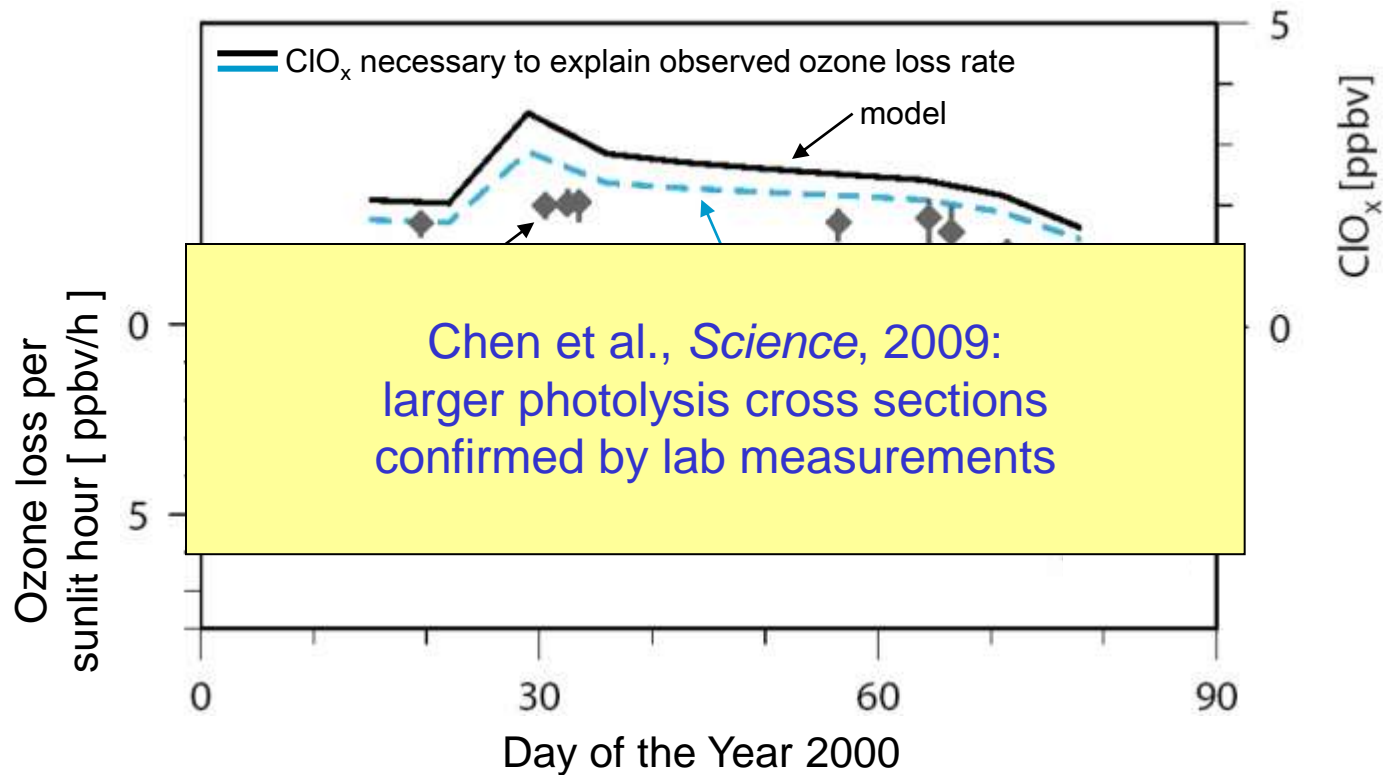


Schofield et al., GRL, 2008

Polar ozone loss process

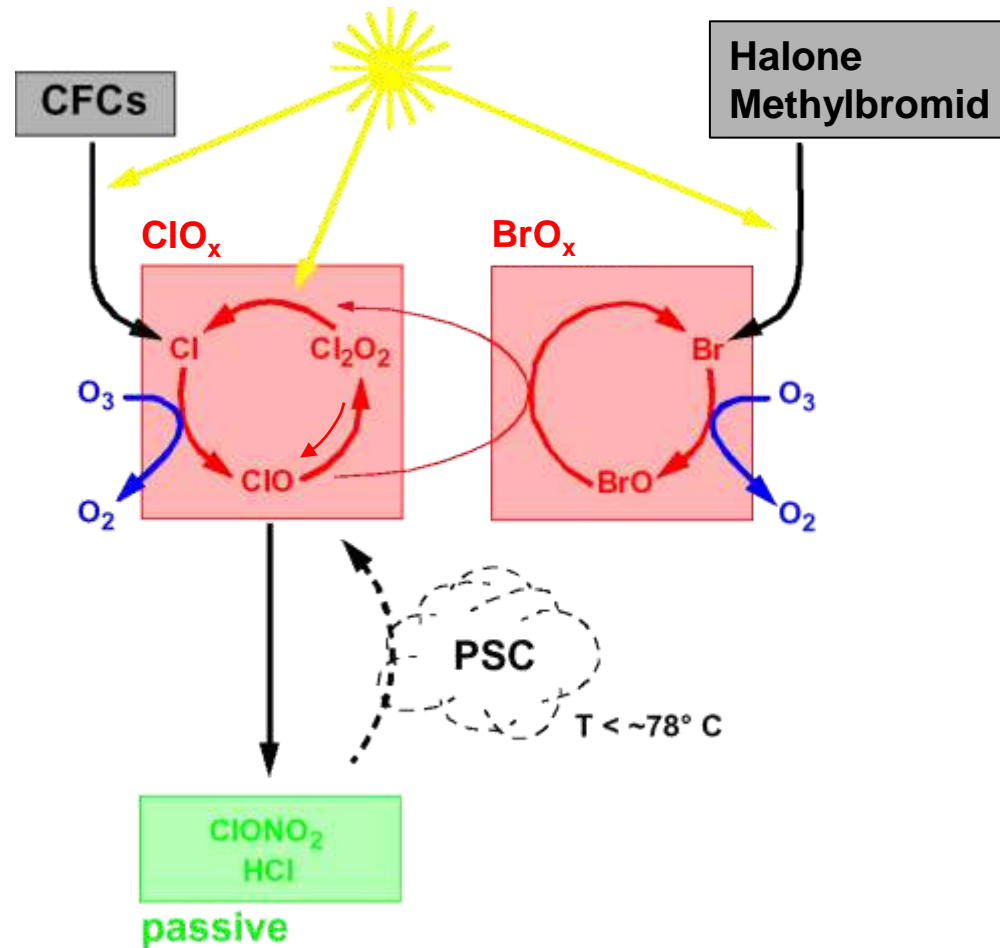


Theoretical understanding of polar ozone loss process

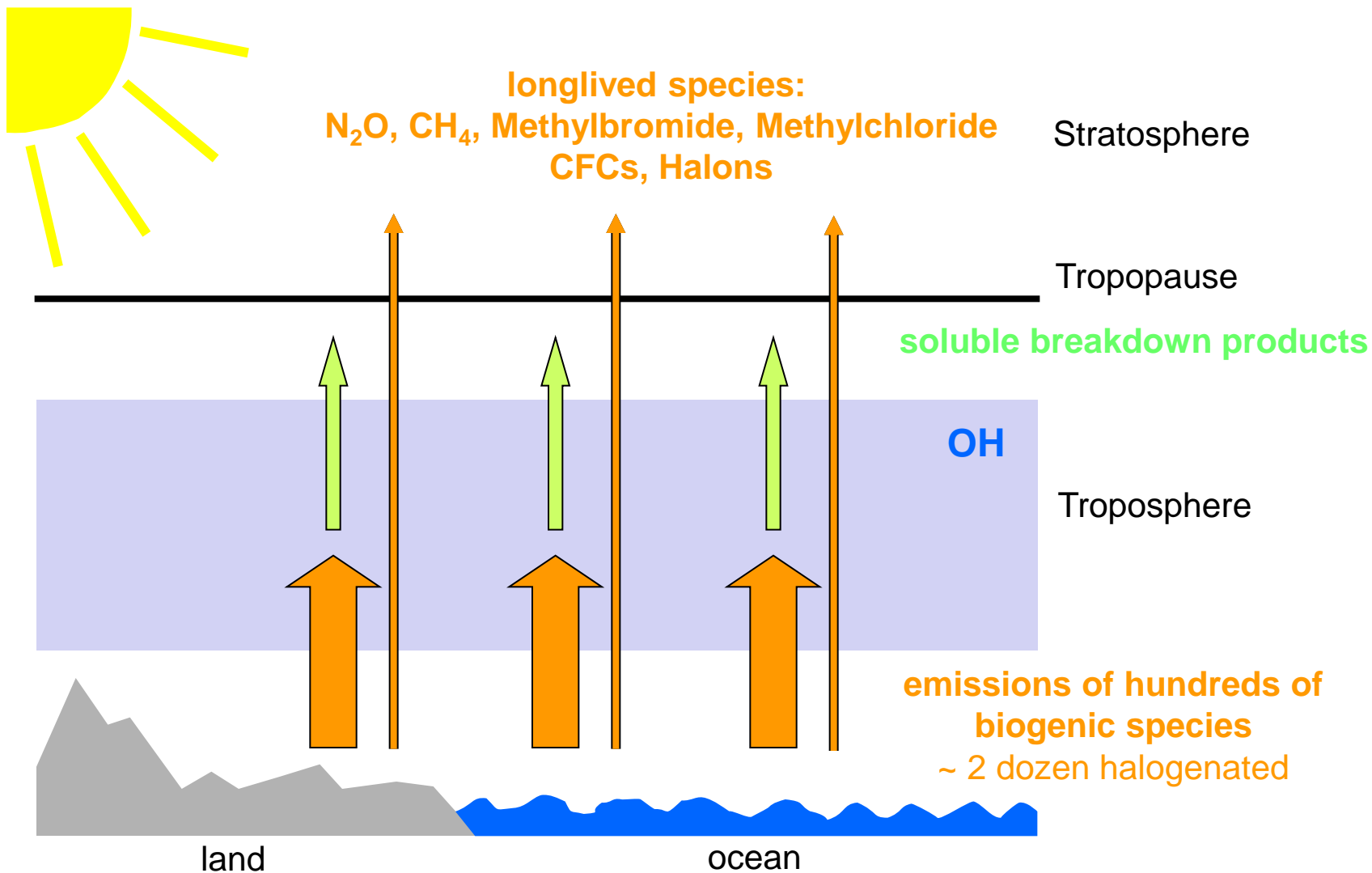


Update of Frieler et al., GRL 2006; WMO 2007

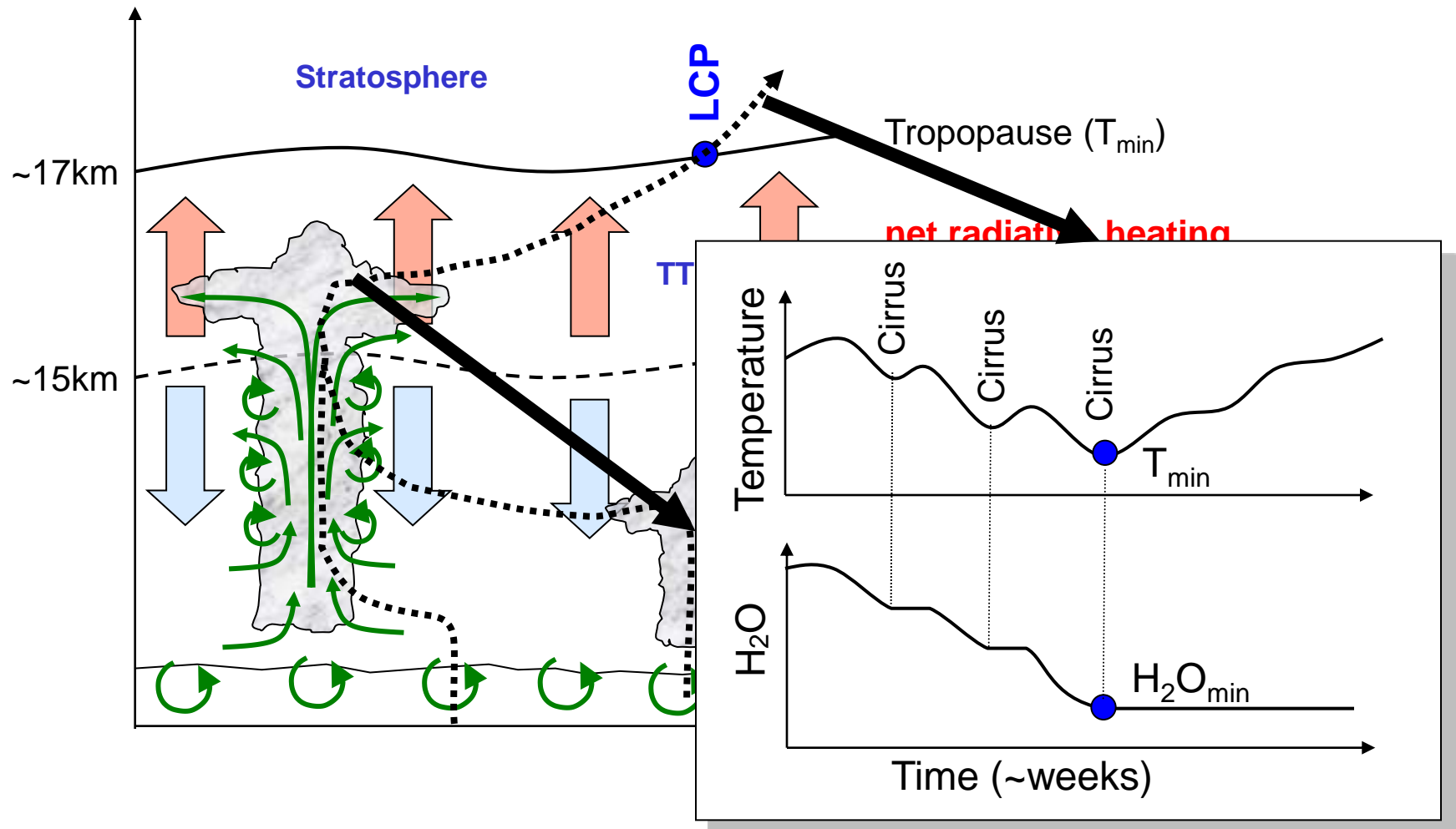
Polar ozone loss process



The „OH shield“



Transport into the Stratosphere



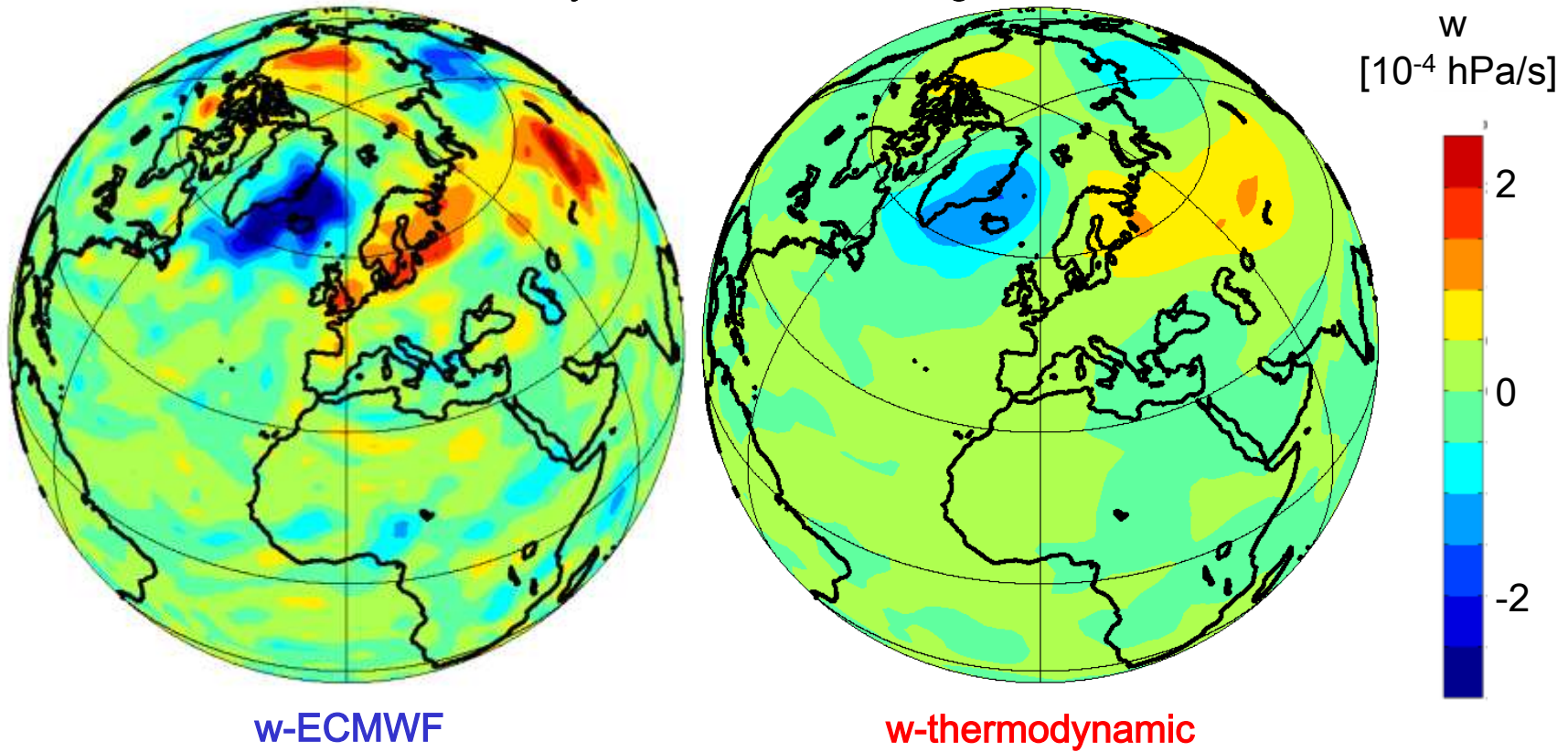
Thermodynamic vertical wind

w from the continuity equation for mass:

From the continuity equation for thermal energy follows:

$$w = (Q - \partial_t \theta - \frac{u}{a \cos \varphi} \partial_\lambda \theta - \frac{v}{a} \partial_\varphi \theta) / \partial_z \theta$$

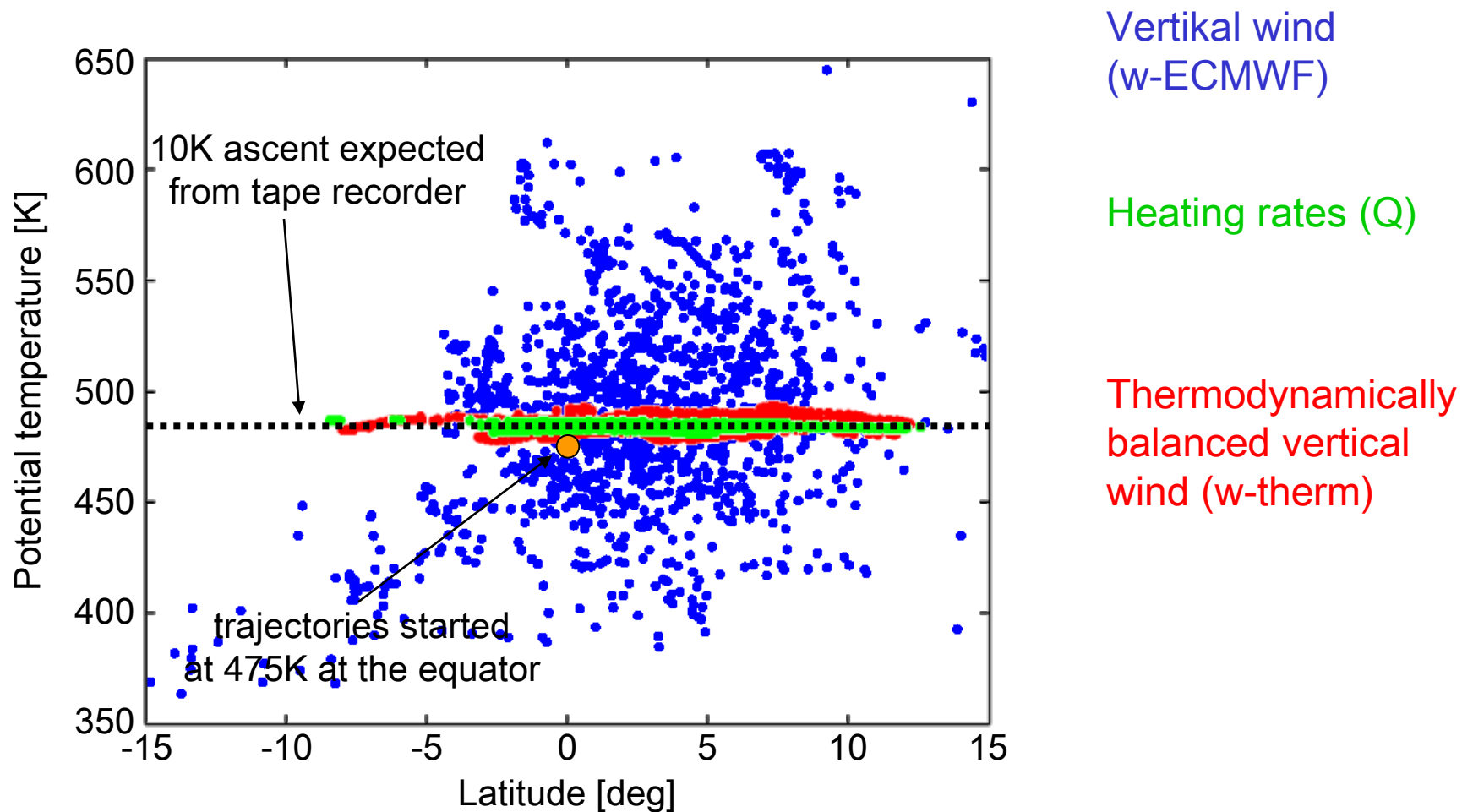
1 January 2000, 24 hour average



Wohltmann and Rex, ACP, 2008

20 days lagrangian transport in ATLAS

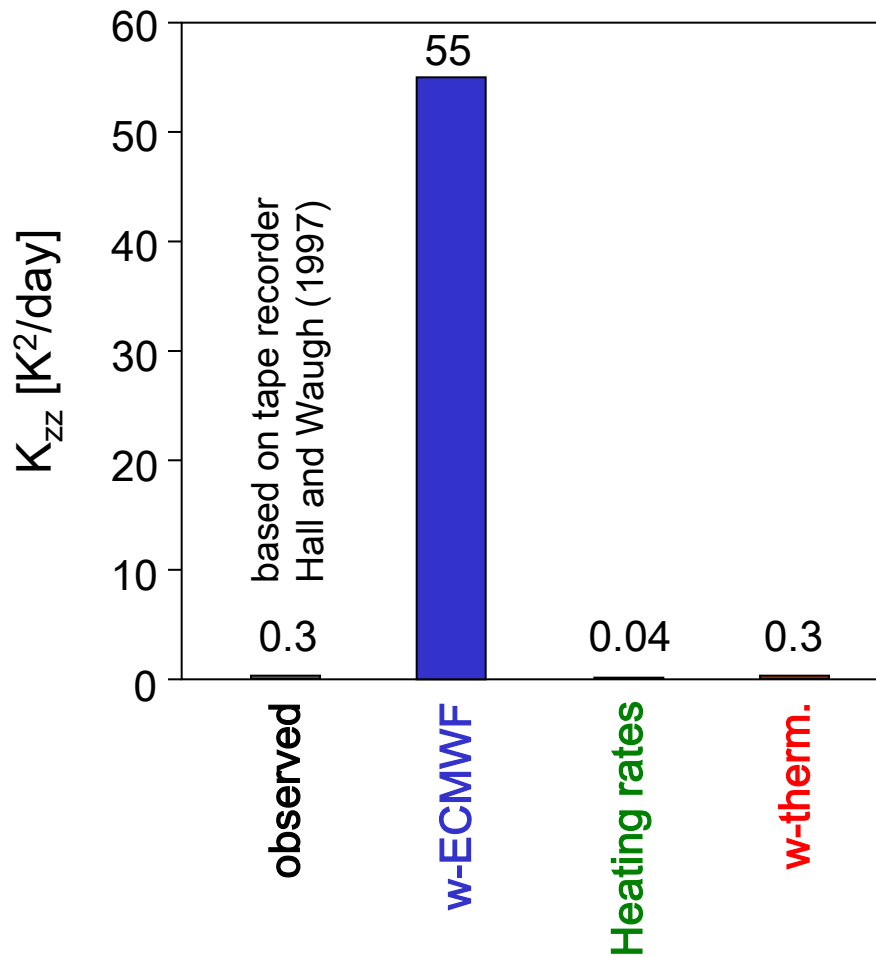
starting 1 January 2000



Wohltmann and Rex, ACPD, 2007

Vertical diffusion rate

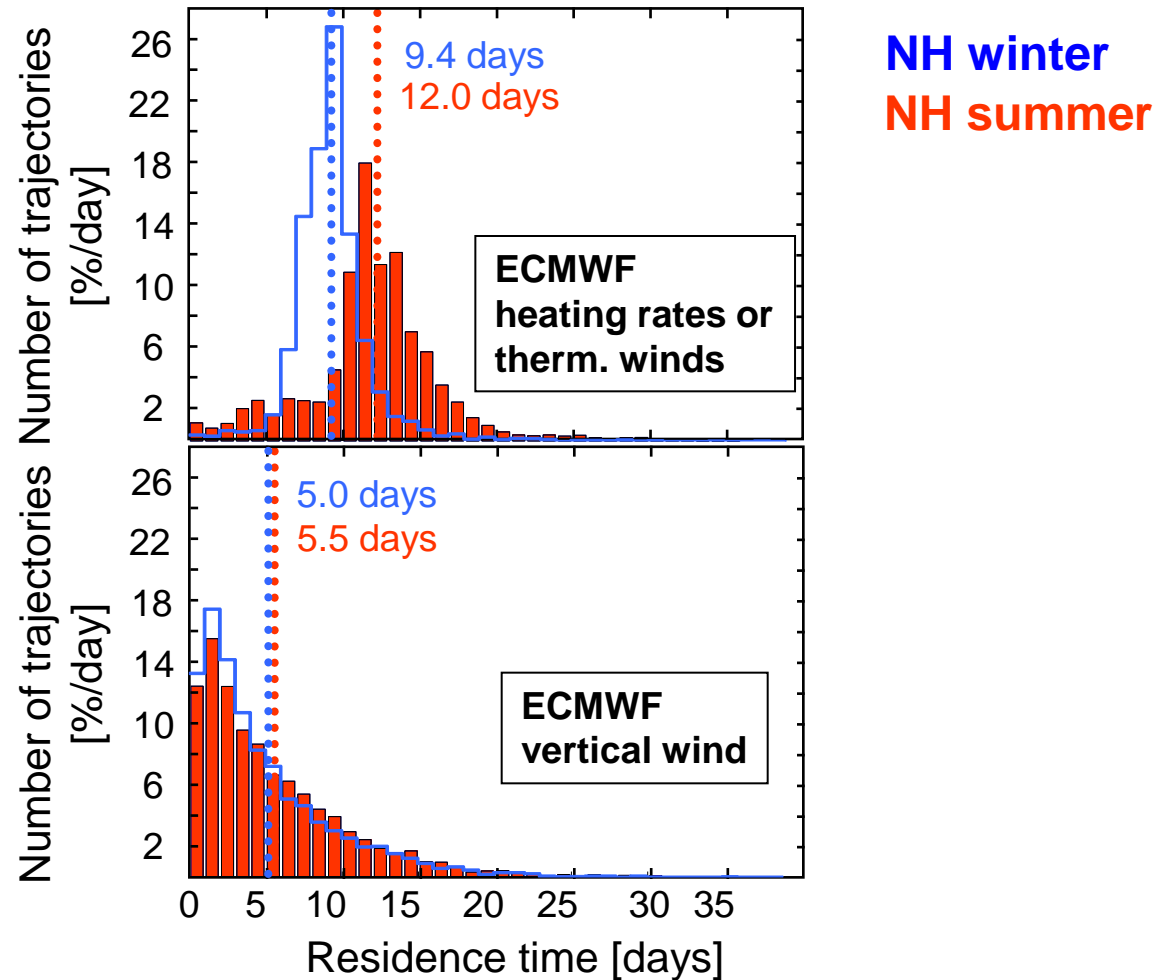
20 km altitude



Wohltmann and Rex, ACPD, 2007

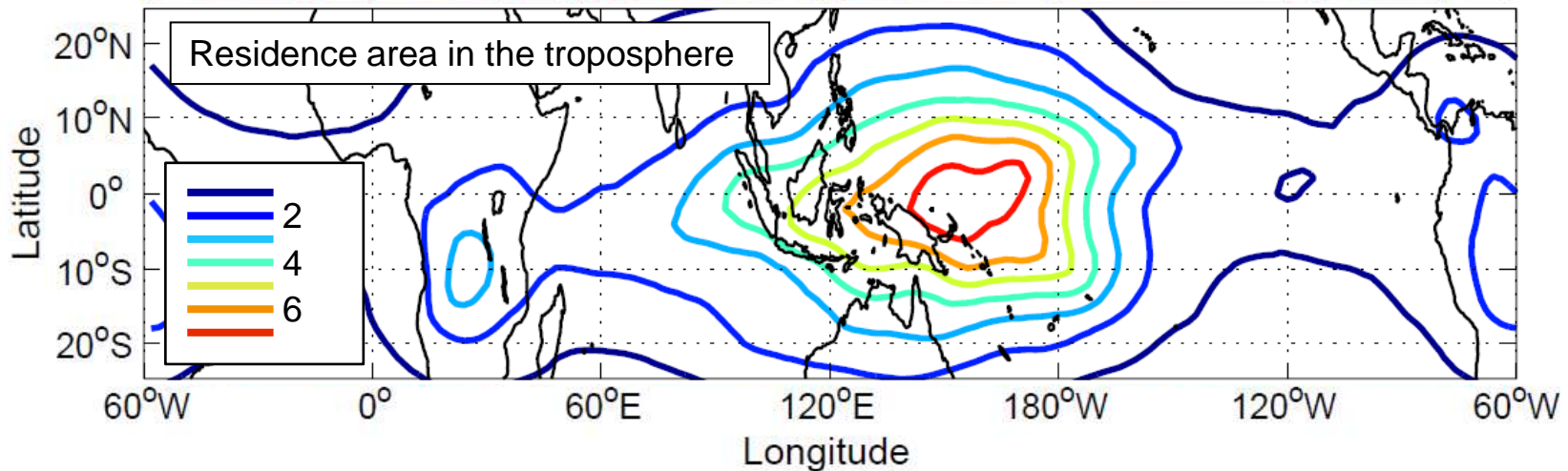
Residence time in upper TTL

(between 385-395 K)



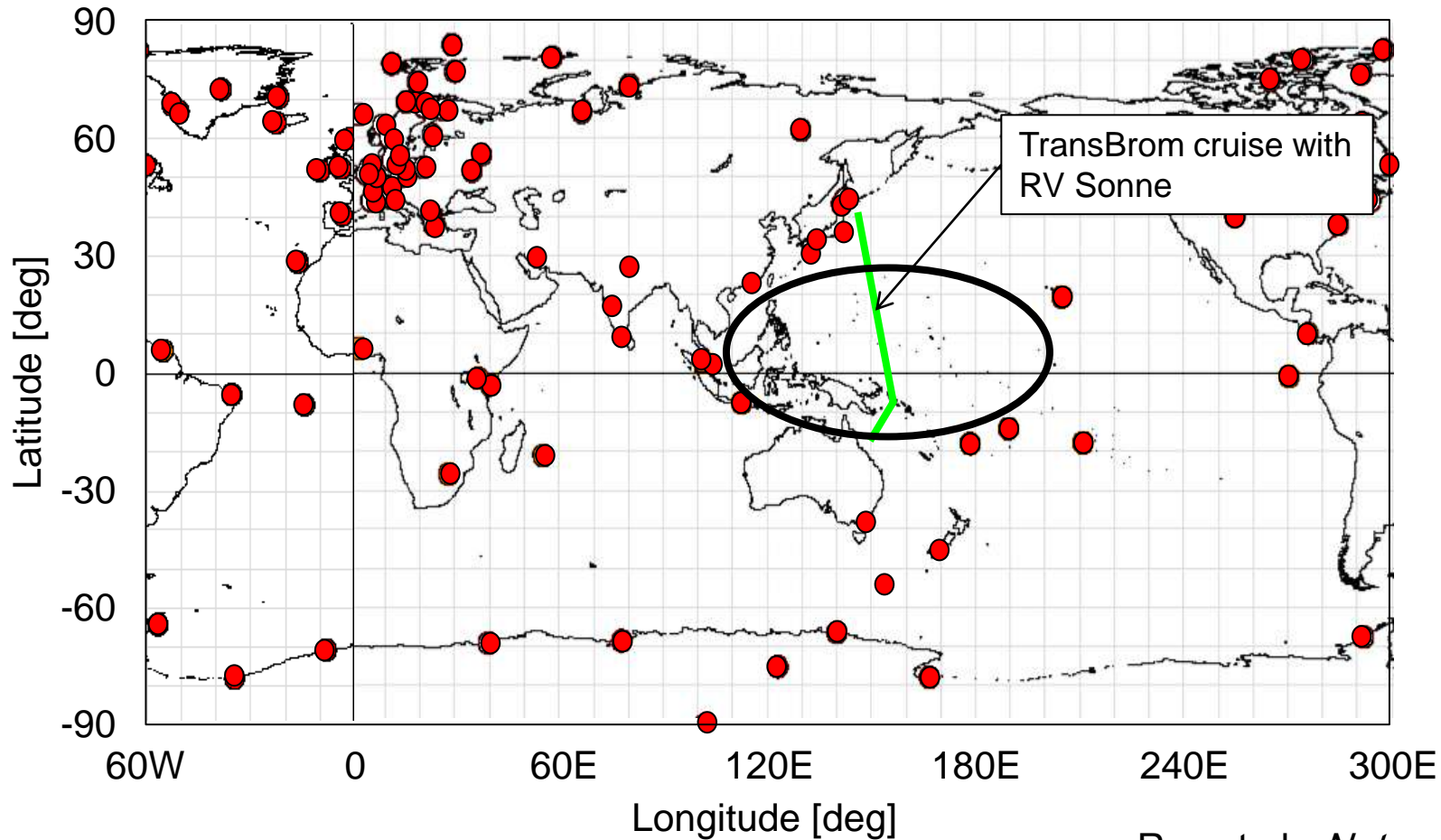
Geographic distribution of time spent in the troposphere for stratospheric air

During transport from the boundary layer to the LCP
Given relative to the tropical average
Based on ATLAS



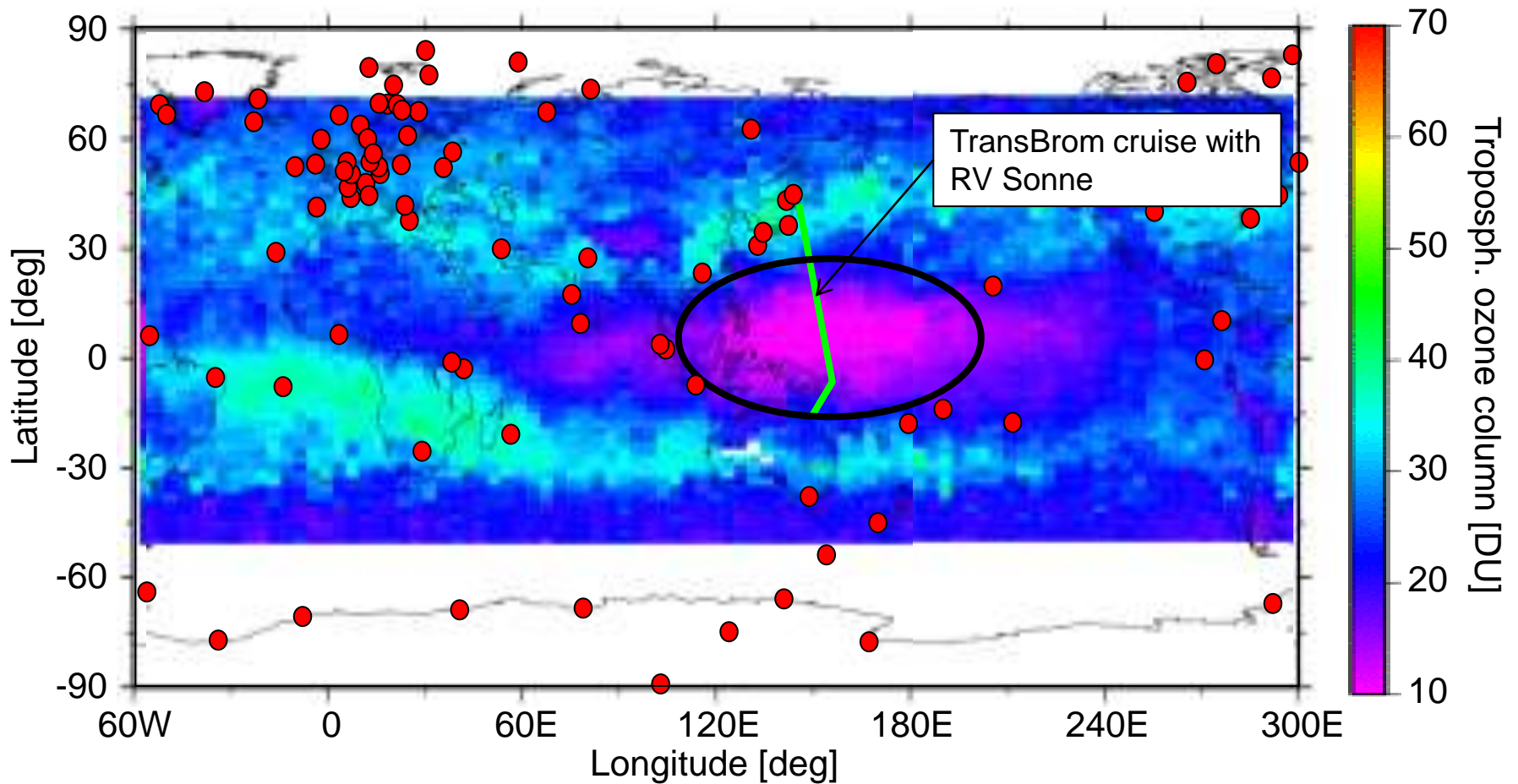
Rex et al., *Nature*, under review

Global ozonesonde station network



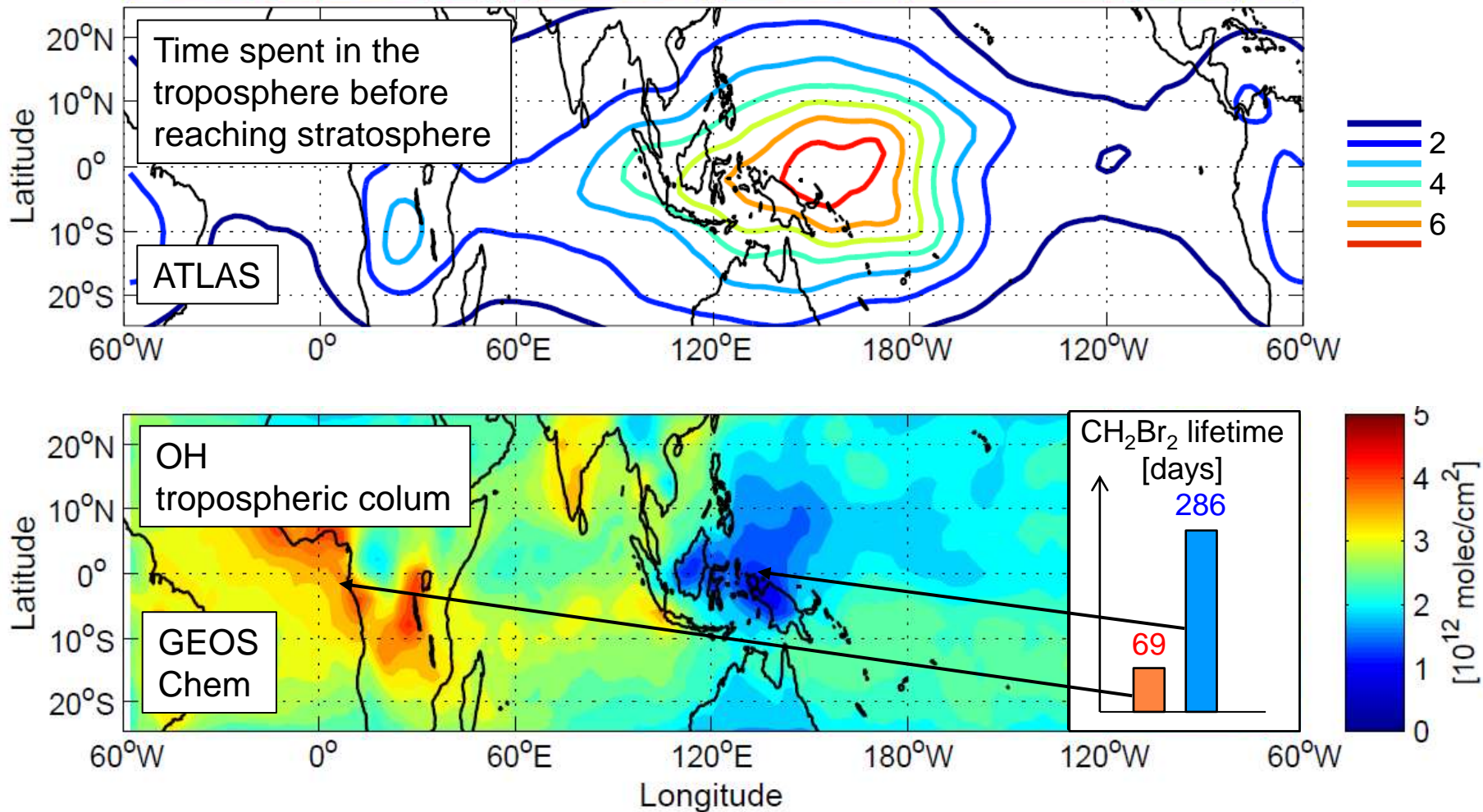
Rex et al., *Nature*, under review

TES tropospheric ozone column October 2009

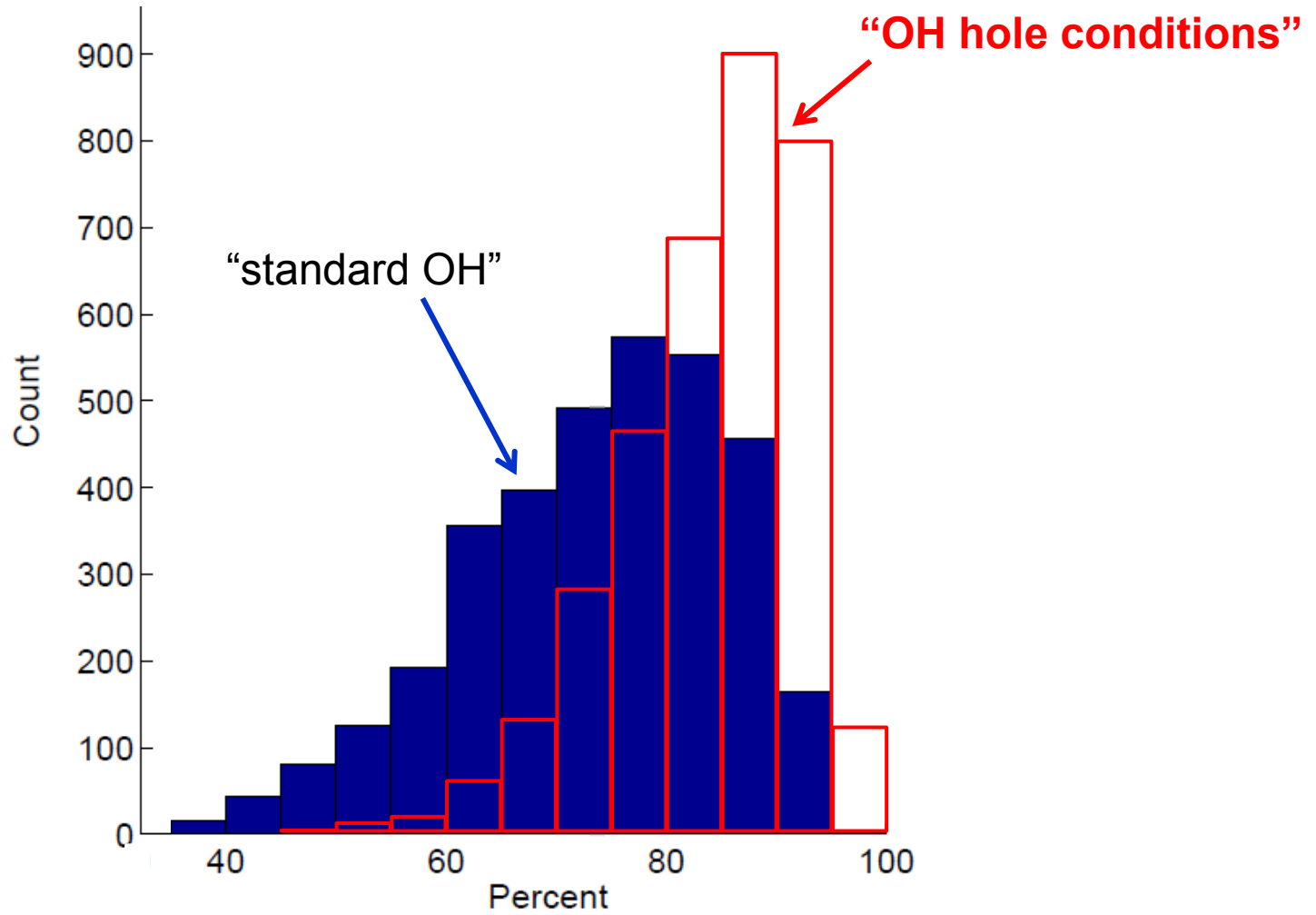


Rex et al., *Nature*, under review

Source region for stratospheric air OH distribution / SO₂ lifetimes

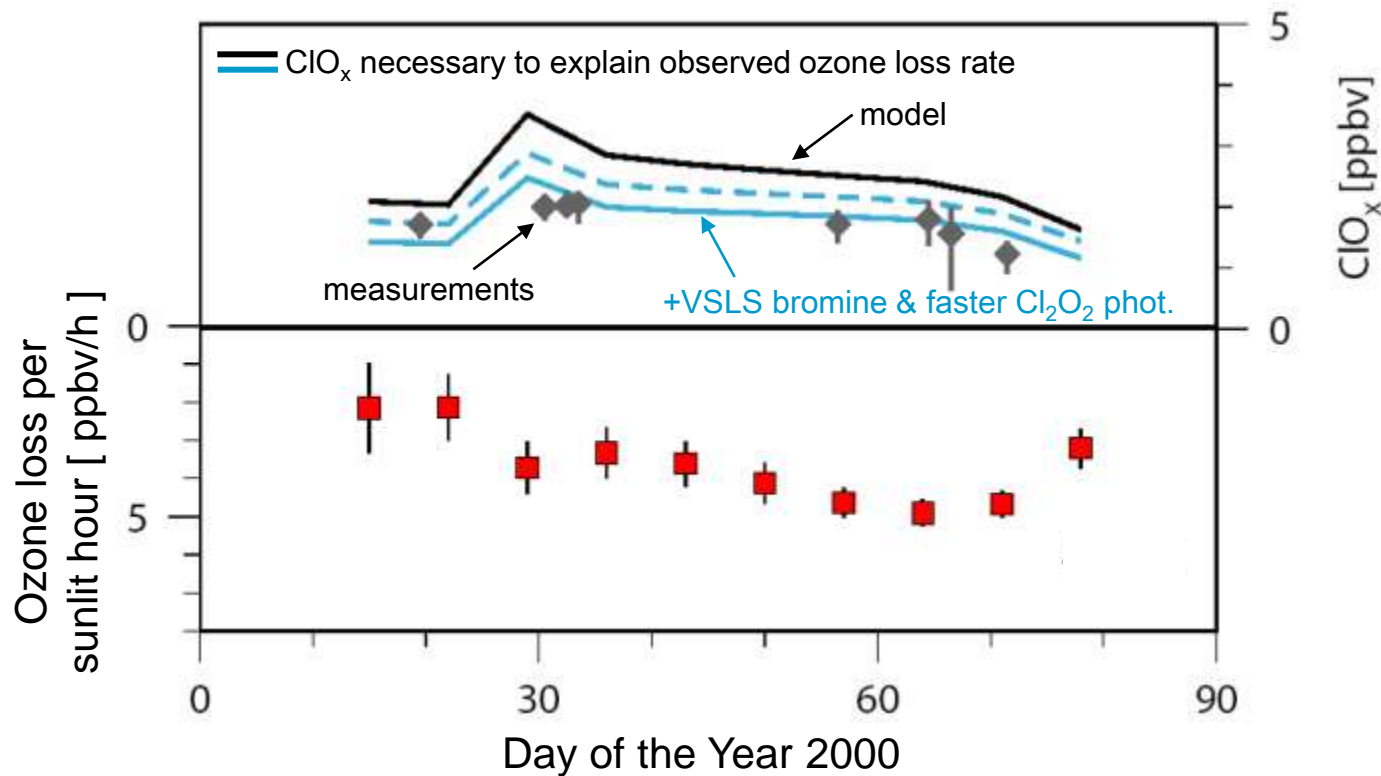


Fraction of CH_2Br_2 reaching the stratosphere



Rex et al., *Nature*, under review

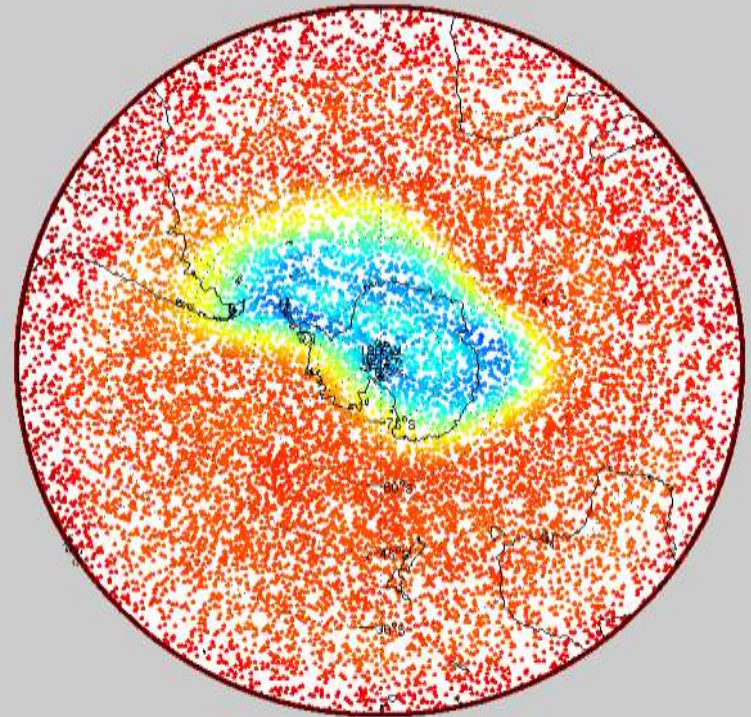
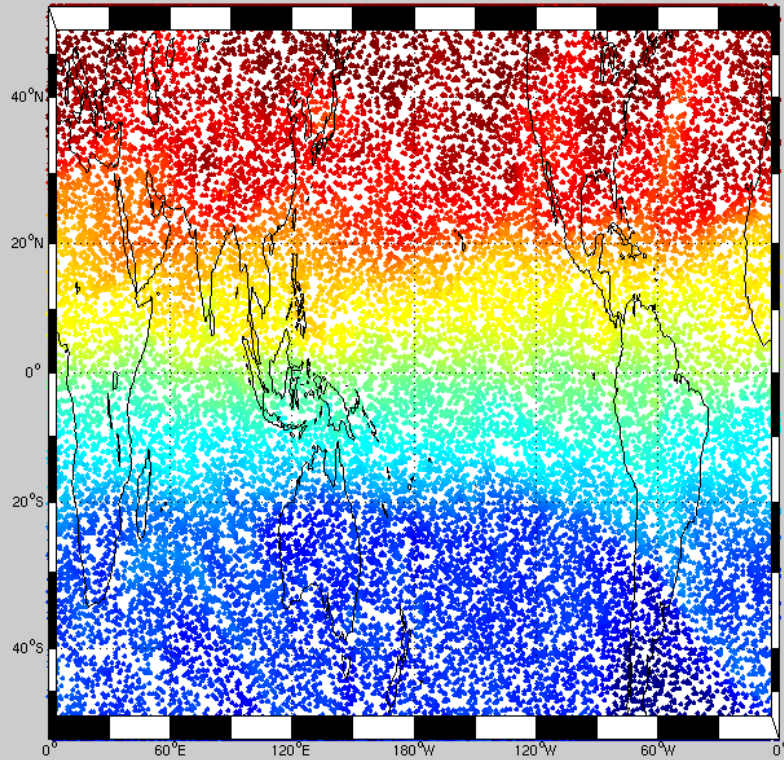
Theoretical understanding of polar ozone loss process



Update of Frieler et al., GRL 2006; WMO 2007

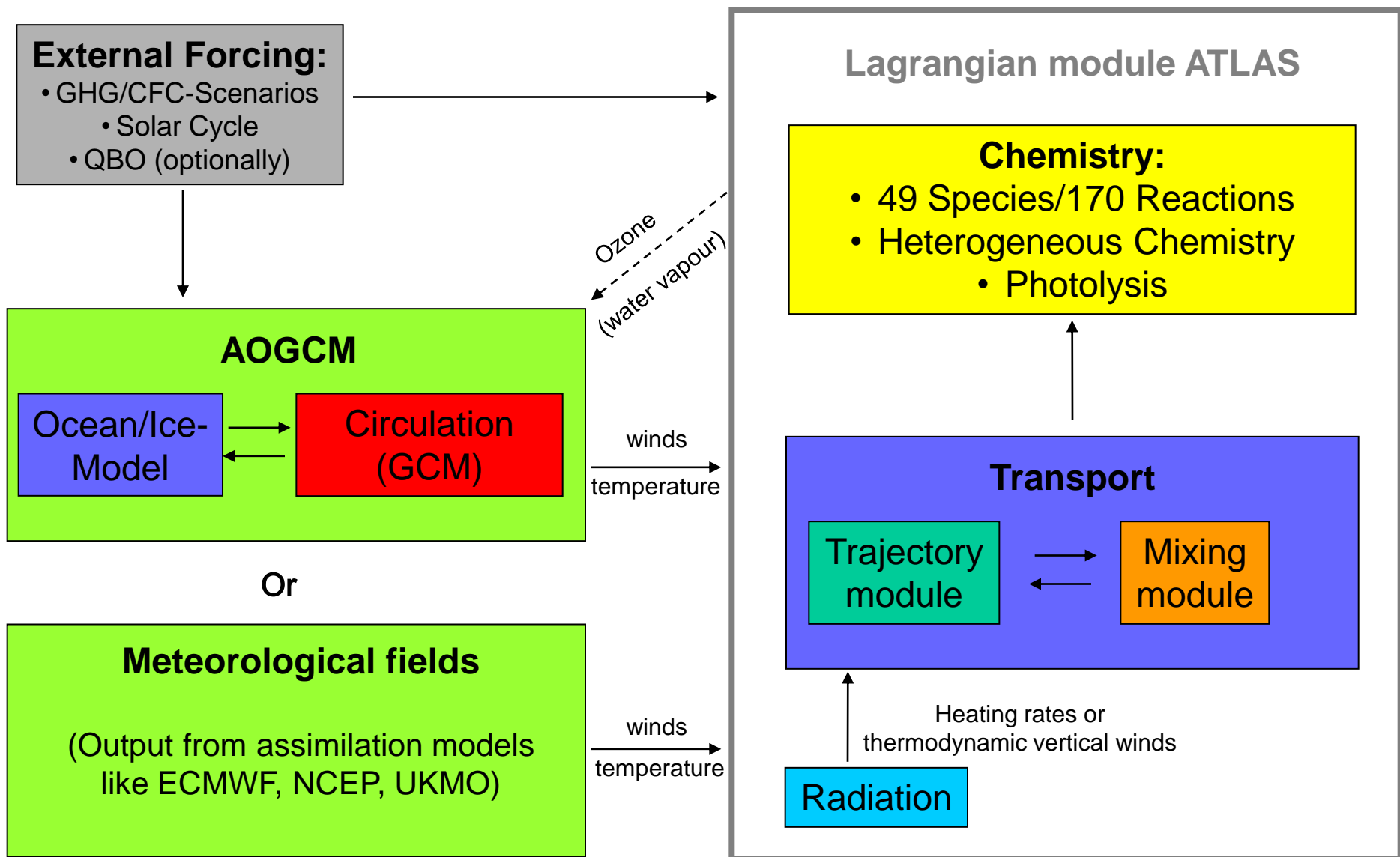
Lagrangesche Modellierung – ATLAS

~20km altitude, 20 model days, dynamical tracer (PV), ~50km resolution run



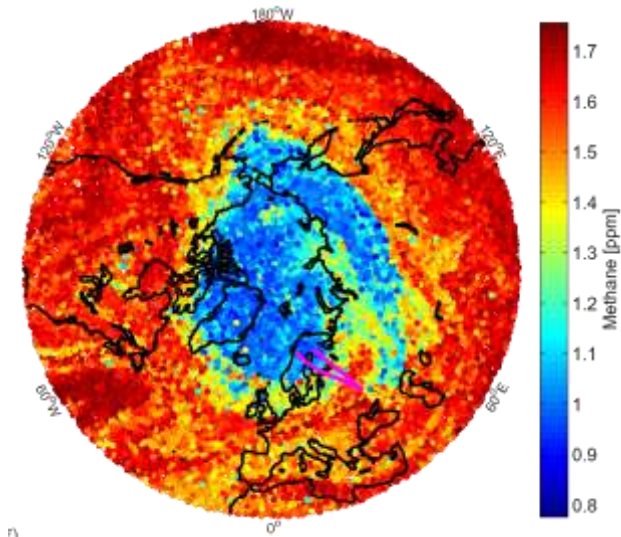
- detailed homogeneous and heterogeneous chemistry
- Lagrangian particle sedimentation scheme
- no numerical diffusion, sophisticated 3d mixing scheme
- full parallel architecture – long integrations for climate runs feasible

ATLAS – fully lagrangian chemistry-/transport modell

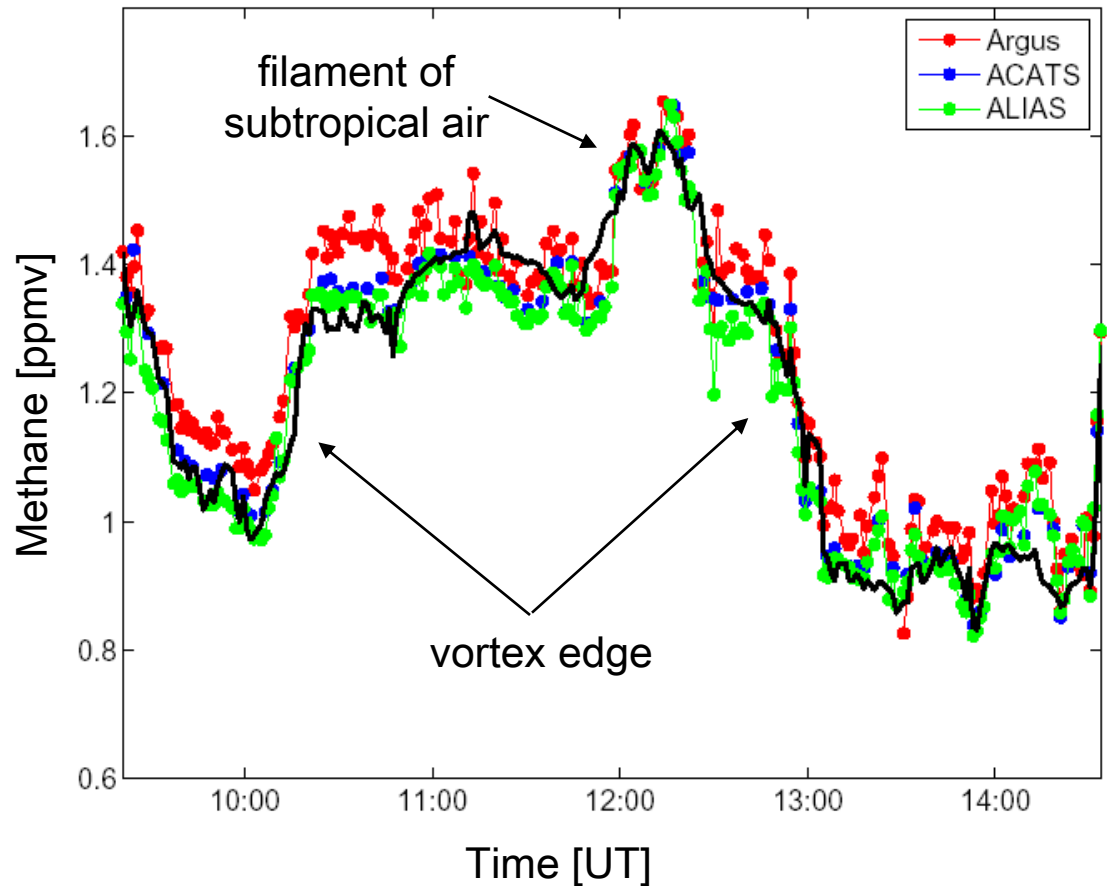


Wohltmann & Rex, 2009; Wohltmann, Lehmann, and Rex, 2010

ATLAS vs. SOLVE ER-2 27 January 2000



Methane



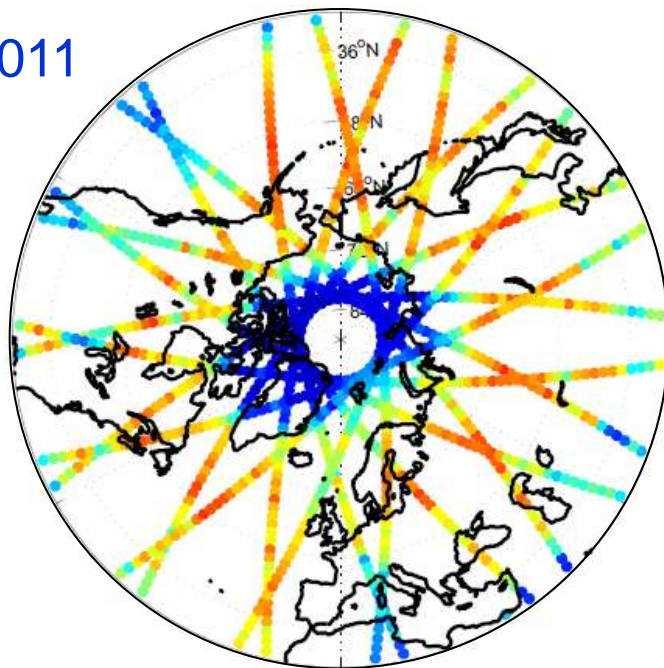
Wohltmann, Lehmann, and Rex, 2010

Ozone

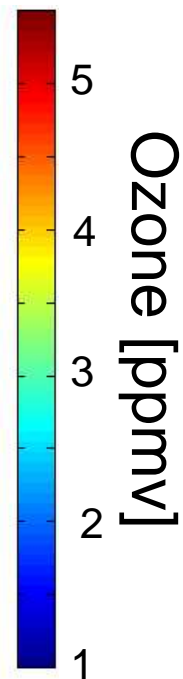
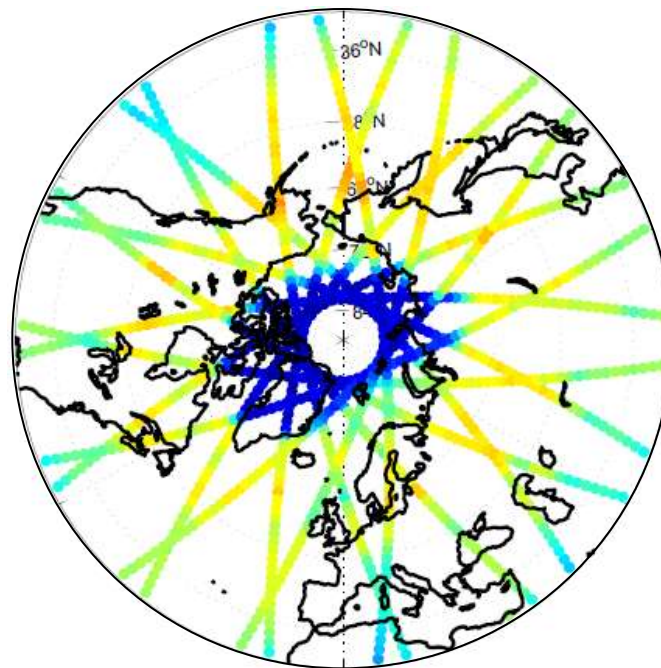
16 March 2011

46 hPa

ATLAS

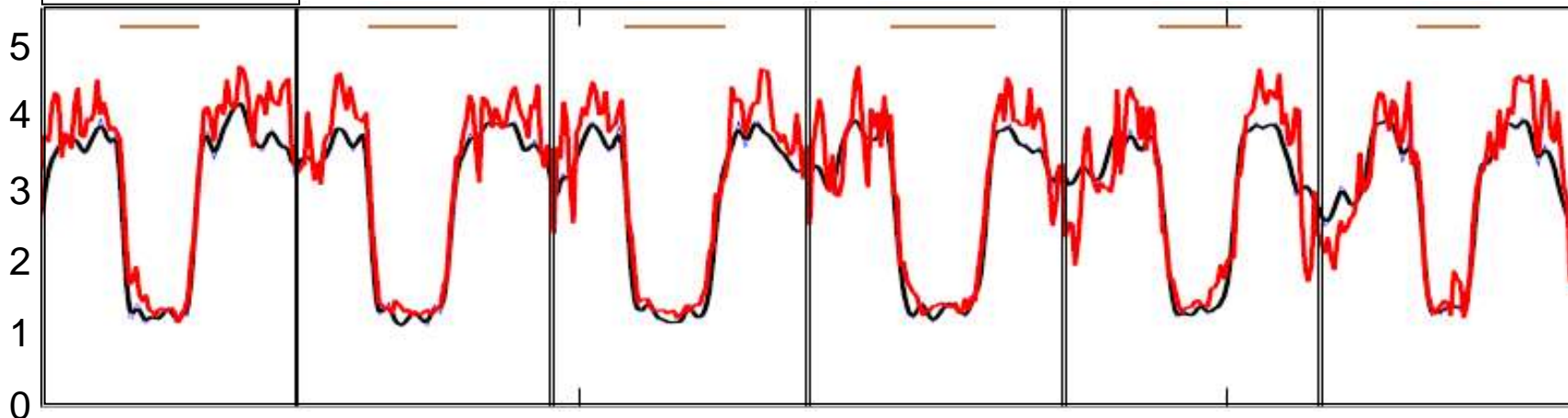


MLS

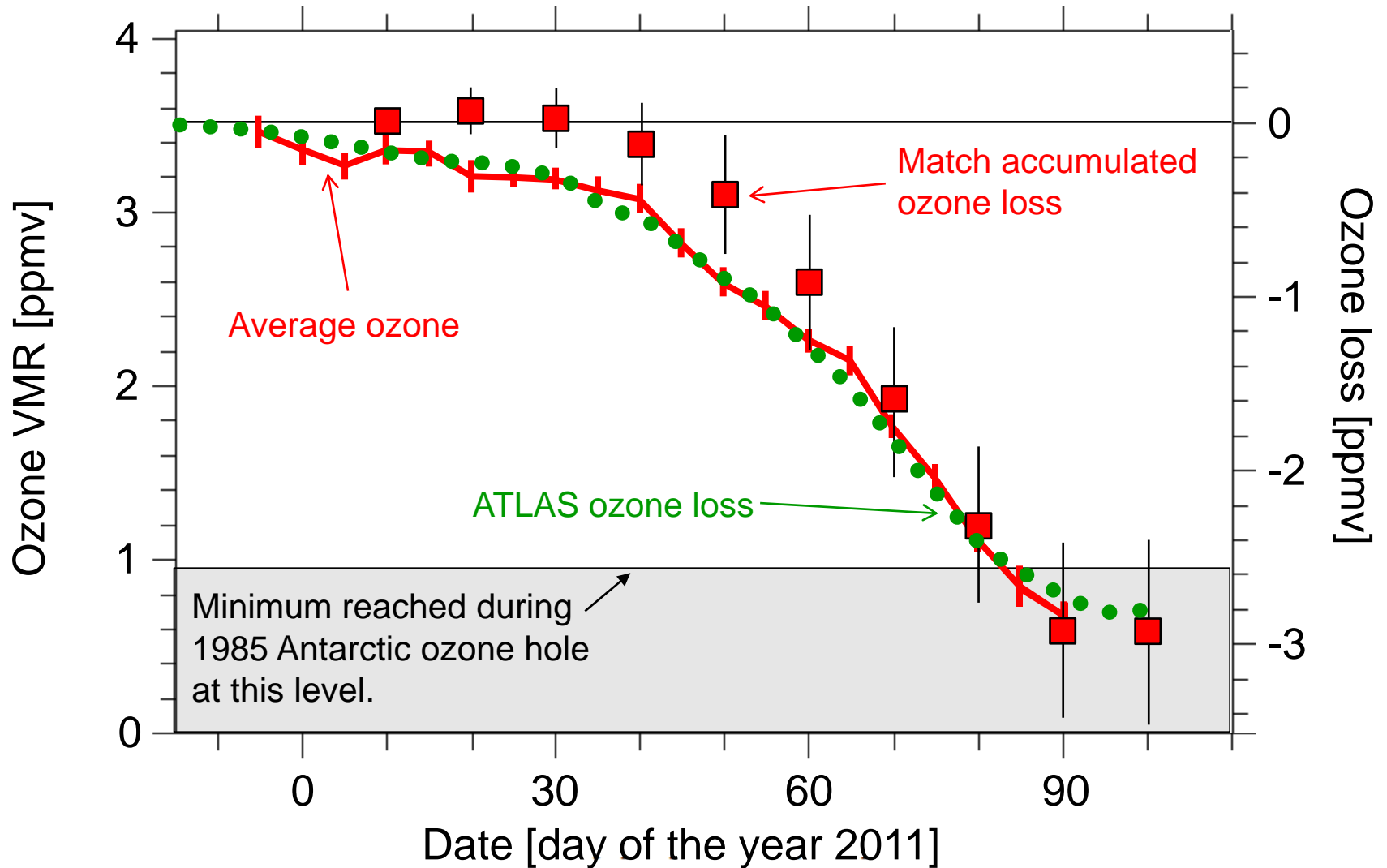


MLS ATLAS

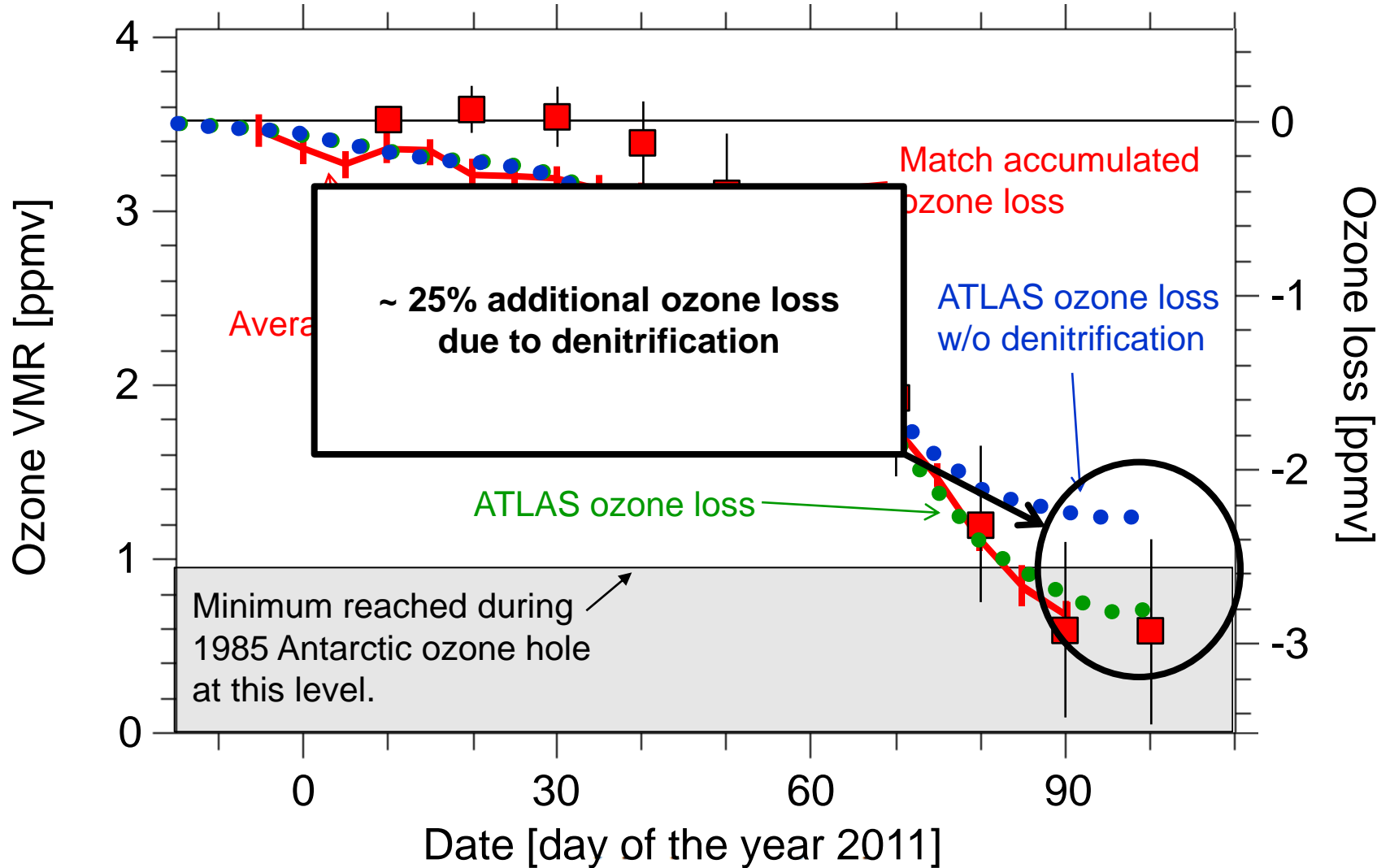
Ozone [ppmv]



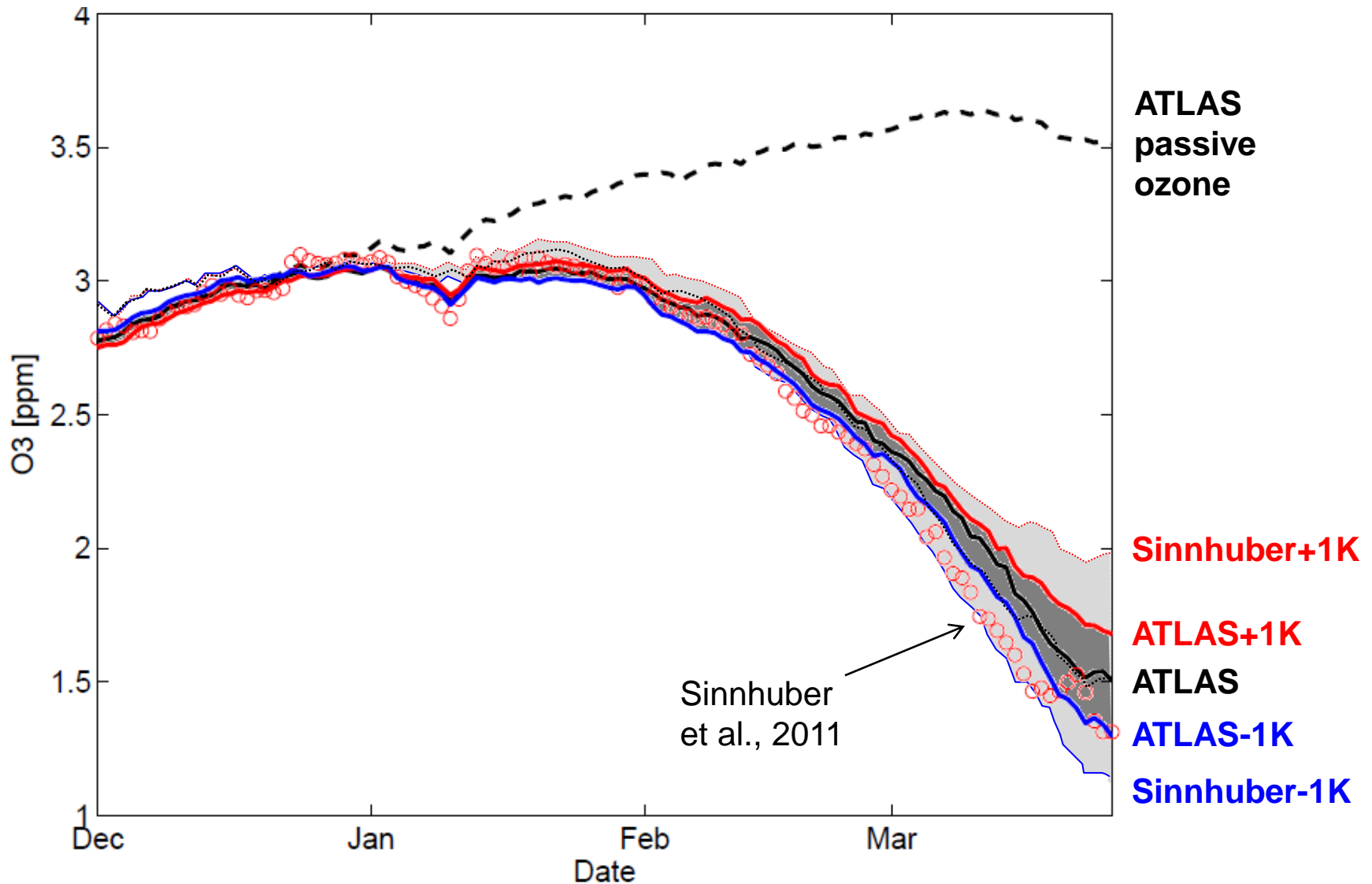
Ozone and ozone loss inside vortex @ $e_{\ominus}=465\text{K}$ (unmixed vortex air)



Ozone and ozone loss inside vortex @ e₀=465K (unmixed vortex air)



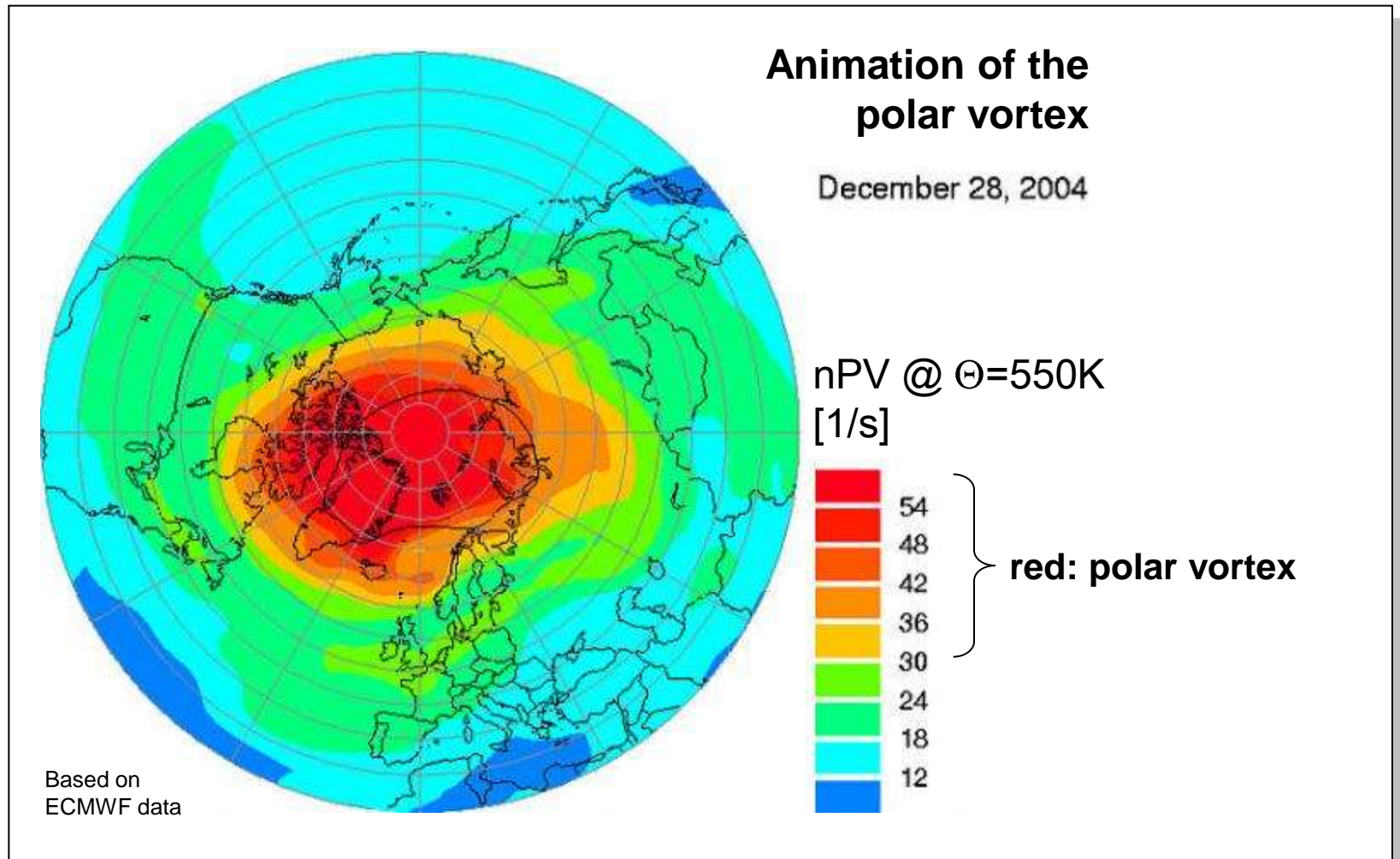
Ozone and passive ozone inside vortex @ fixed $\Theta=475K$ (full vortex average, including mixed areas)



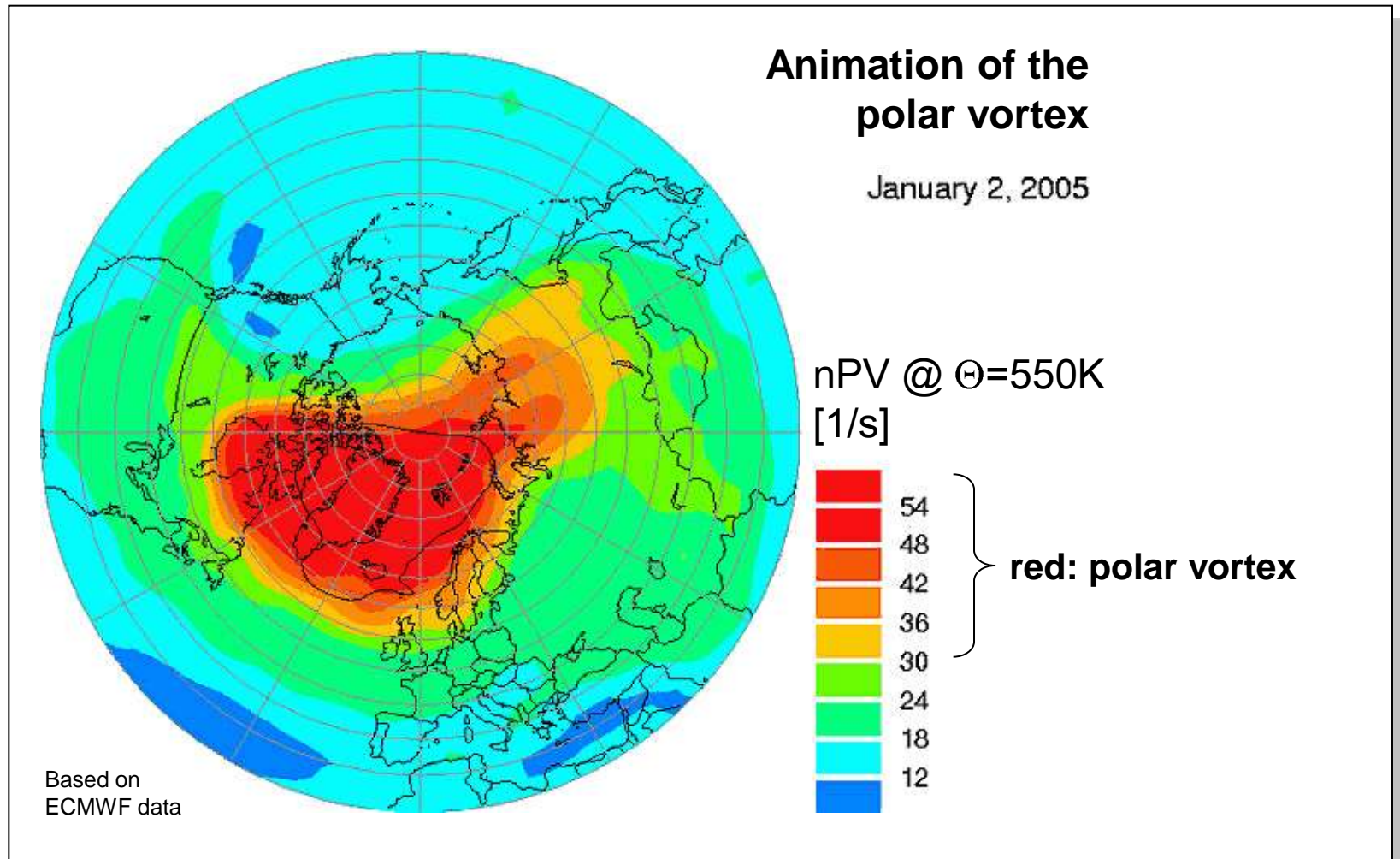
Conclusions

- Arctic ozone loss continues to get worse due to ozone/climate coupling.
- Processes at the tropical tropopause layer, particularly above the West Pacific, play an important role for the global ozone layer.
- A tropospheric ozone and OH hole exists over the tropical West Pacific.
=> Emissions from South East Asia and the tropical oceans there can play a larger role for the stratospheric composition.
- Data assimilation products are the basis for studies of stratospheric processes. These are particularly sensitive on uncertainties in:
 - The temperature fields in the polar lower stratosphere.
 - The vertical wind fields at the tropical tropopause.

Typical movement of Arctic polar vortex



Typical movement of Arctic polar vortex



Impact on mid-latitude UV

PV

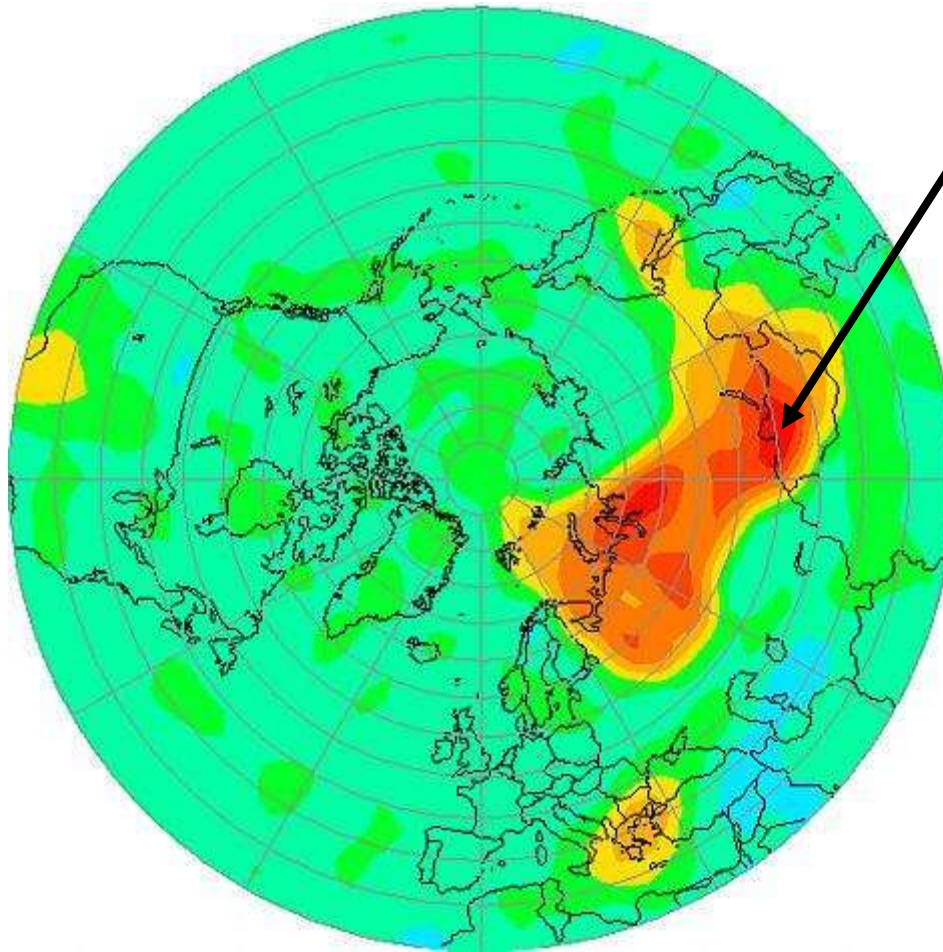
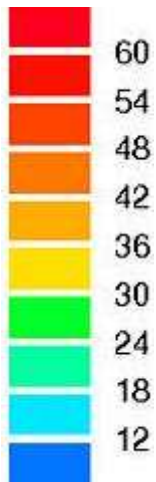
April 22, 2011

Day: 112

Analysis

$PV / 10^6 \text{ K m}^2 \text{ s}^{-1} \text{ kg}^{-1}$

475 K



UV-Index: 8.6

- Sunburn within minutes
- Long term average on April 22: 5.4
- 7 sigma deviation