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With thanks to the NOGAPS-ALPHA team (ALPHA: Advanced Level Physics High Altitude)

Overall theme: Use high-altitude analysis (3D-Var (NAVDAS), Eckermann et al., 2009; Hoppel et al., 2008) to initialize and compare with a series of short/medium term NOGAPS-ALPHA forecasts

Forecast model resolutions: T79, T239, T479

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1. Case studies of the SSWs of 2008 (minor) and 2006 (major, extended) with combinations of parameterized GWD scheme

orographic (OGWD): Palmer et al (1986) non-orographic ("multiwave") (MGWD): Garcia et al. 2007

What are the relative roles of OGWD and MGWD in the various phases of an SSW?

2. Solstice forecasts at a range of resolutions, including T479, without any parameterized GWD Look at GWD features: warm winter stratopause cold summer mesopause

What spatial resolution is required to model GWD features without parameterizations?



### NAVDAS and the NH winters of 2006 and 2008









#### Label **Dates** Description Period 1 15–20 Jan 2008 dynamically quiet Period 2 20-25 Jan 2008 minor warming Period 3 20-25 Jan 2006 major warming, initial phase Period 4 25-30 Jan 2006 major warming, extended phase Period 5 2-7 Feb 2006 major warming, extended phase

Three forecasts per period:T79O:OGWD onlyT79MO:Both OGWD and MGWDT239O:OGWD only, but higher spatial resolution(note: NAVDAS is always T79)

Compare absolute forecasts and changes over each 5 day period with NAVDAS



#### 3 Simulations of minor 2008 warming (Period 2): Jan 24-Jan 20th





In general, differences are not large, but cooling at 0.01 hPa is better simulated in MO





$$\overline{u}_t + \overline{v}^* \left[ (a \cos \phi)^{-1} (\overline{u} \cos \phi)_{\phi} - f \right] + \overline{w}^* \overline{u}_z = (\rho_o a \cos \phi)^{-1} \nabla \cdot \mathbf{F} + X,$$

#### Two components of F, Eliassen-Palm (EP) Flux

$$\begin{split} \mathbf{F}_{\mathbf{p}}[\phi,z] &= \rho_o a \cos\phi \Big[ (\overline{u}_z \overline{v'\theta'} / \overline{\theta}_z - \overline{u'v'}), \\ & \left( f - \frac{1}{a \cos\phi} [\overline{u} \cos\phi]_\phi \right) \overline{v'\theta'} / \theta_z \Big], \end{split}$$

$$\mathbf{F}_{\mathbf{g}}[z] &= \rho_o a \cos\phi \Big[ \overline{-[u'w']} \Big].$$

$$\begin{array}{l} \mathbf{3} \text{ gravity wave forcing terms} \\ \mathbf{X} \text{ is broken out according to } \mathbf{X}_0 \text{ and } \mathbf{X}_M \end{split}$$















#### Three simulations of the 2006 event















**2008** 

#### 2006



Longitudinally asymmetric GWD (in 2008)
 Instabilities near the stratopause (in 2006) (Thayer et al., 2010)



#### Hovmoller of geopotential height anomaly











**Blue:** divergence, red: convergence











lat/lon	2 km	1 km
1 50°	T79, L74	T79, L115
0.50°	T239, L74	T239, L115
0.25°	T479, L74	ХХ

**Background:** Gravity Wave Resolving (GWR) middle atmosphere models

Hamilton, colleagues (1991, 2001, 2006, 2008): Increased horizontal resolution

Watanabe et al (2008): Increased vertical resolution (L256, only T213) Subsequent studies (Sato et al., 2009; Sato et al., 2012)

Here: 10 day forecasts for June 2007 and January 2008 w/o parameterized GWD. Compare forecast minus analysis (F-A) as measure of model bias and guide to accuracy of the physics (i.e. the resolved gravity waves)





What is the effect of changing horizontal and vertical resolution on gravity wave driven features?

- $\rightarrow$  cold summer mesopause
- $\rightarrow$  warm winter stratopause

Where are the gravity waves in NOGAPS-ALPHA and how do we compare with other published GWR models?

Highlight an important uncertainty in defining a critical resolution for resolving gravity waves







Divergences between forecast and analysis attributed to relative lack of resolution of GWD

#### 2D plot at 240 hours: T79 and T239 with Analysis







#### **T479 at 240 hours**







#### T479, Two other Cases





L115: 25 levels troposphere, ~80 levels to ~0.001 mb, 10 levels to the top vertical resolution increased from 2 km  $\rightarrow$  1 km in strat and mes.







#### Filter the T479 runs







#### Global snapshot of small scale waves (120 hours after June 10)



In USLM, most activity is in the winter hemisphere







Small spatial scales evident in both summer and winter



#### Momentum flux for different model resolutions







#### Momentum Flux, June 10-20<sup>th</sup> average







Models have to dissipate energy at high wave numbers  $(4^{th} \text{ order diffusion} \rightarrow hyperdiffusion})$ How this is done can change the effective resolution (Skamarock, 2004)

**Timescale for dissipation- vary K**<sub>h</sub> (our baseline runs use 2.8 hours for T479)

$$\frac{a^4}{K_h(n_T(n_T+1))^2}.$$

Validate the dissipation by the kinetic energy spectrum/mass Koshyk et al., 1999; Koshyk and Hamilton, 2001

$$E_{n,m} = \frac{a^2}{4n(n+1)} \left( |\xi_{n,m}|^2 + |\delta_{n,m}|^2 \right),$$





# Skamarock, [Mon. Wea. Rev., 2004]



FIG. 10. Schematic depicting the possible behavior of spectral tails derived from model forecasts. Using the methodology outlined in the appendix to compute the spectra, limited-area models (including WRF) usually produce the slightly upturned tail shown at left.





#### **200 mb**



**NOGAPS-ALPHA** seems steeper at high wavenumber than k<sup>-5/3</sup>

# KE spectra for two diffusion time scales





Slow Diffusion (5.6 hours)





## Changing spectral diffusion $\rightarrow$ Zonal mean effects





Weaker damping yields colder summer mesopause and warmer winter stratopause





2008 was a minor warming: Some attenuation of OGWD High altitude eastward forcing (mesospheric cooling), better resolved with "multiwave" GWD but effects are minor

2006 was a major → extended warming heavy attenuation of OGWD <u>Defining characteristic:</u> MGWD swings from eastward → westward

T239 with only OGWD is insufficient to resolve all GWs

Two sources of high altitude planetary waves in 2008: asymmetric GWD in 2006: instabilities at the stratopause





Our approach using DA + 10 day forecasts is viable for exploring resolution effects.

Increased horizontal resolution is more effective than increased vertical resolution for extratropical zonal mean middle atmosphere simulations

**T479 is insufficient** (in our model) to capture winter stratopause or summer mesopause unless spectral diffusion is unrealistically weak. Gravity wave activity significant for *k* > 239.

What resolution is needed? <u>No single answer.</u> Instead of GWD parameterization tuning, must consider spectral dissipation at high wave numbers also, convection scheme, numerical solver







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